Styles and Scales of Structural Inheritance throughout Continental Rifting

Examples from the Great South Basin, New Zealand

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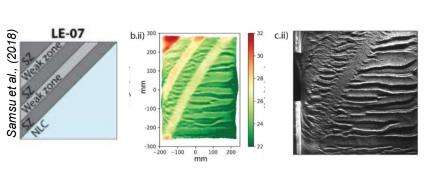




Rationale



Continental crust comprises distinct crustal units and intruded magmatic material brought together throughout multiple tectonic events



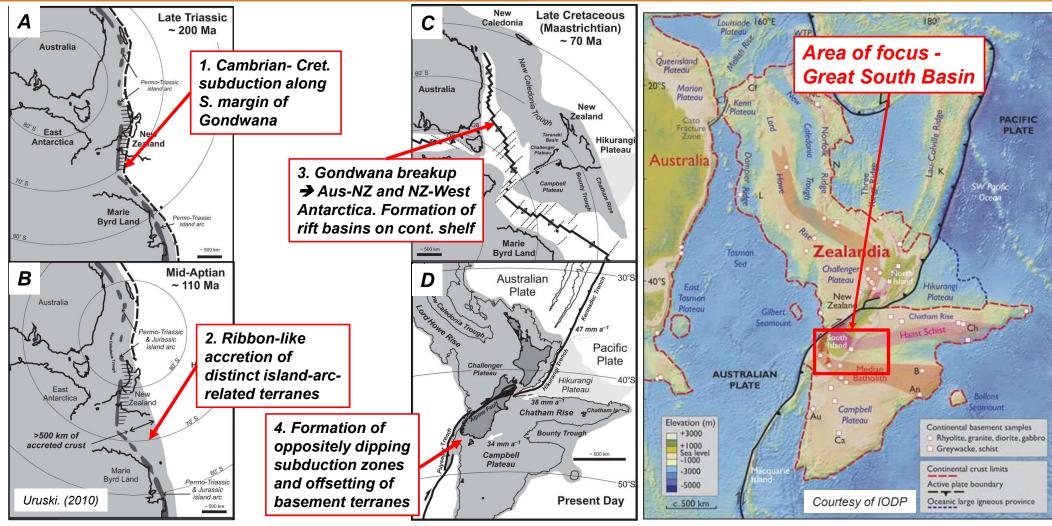
- A Model 3 GI - G3 (Pa) GI - G3 (Pa)G
- Crustal/lithospheric strength may influence the rift structural style and physiography
- Strain may initially localise in weaker areas of lithosphere, rather than at the boundaries between different domains

How do lateral crustal strength contrasts, along with prominent crustal boundaries, influence rift structural style and physiography?

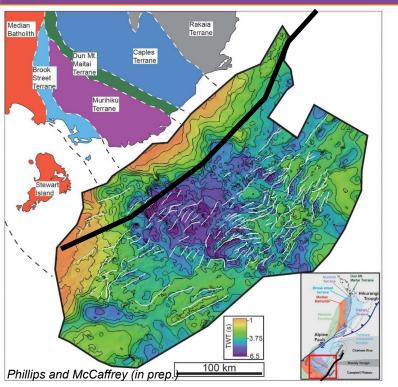
- The Great South Basin, New Zealand forms atop basement comprising multiple distinct terranes and magmatic intrusions.
- The extension direction during rifting is parallel to the terrane boundaries, such that all terranes experience extensional strain

Geological evolution of Zealandia



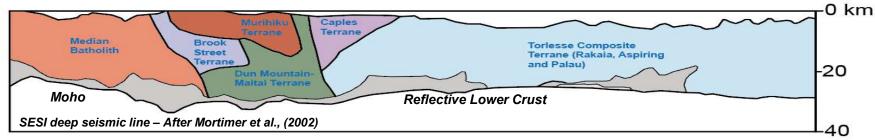


Basement beneath the Great South Basin



Distinct basement terranes of varying strength related to Island Arc system accreted to Gondwana margin

- **Median Batholith** Composite batholith consisting of multiple generations of granitic plutons
- Brook Street Terrane Volcanics and volcanogenic sediments
- *Murihiku* Sediments (relict forearc basin)
- **Dun Mountain-Maitai Terrane** Ophiolite complex (Dun Mt.) and mudstone (Maitai)
- Caples Terrane Volcanogenic sediments, becoming more schistose to NE



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Part I – Influence of crustal-scale terrane boundaries

Terrane boundaries controlling volcanism



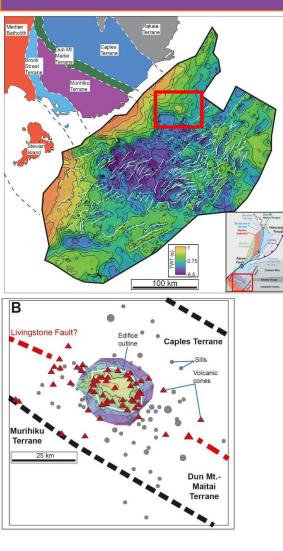
Shield Volcano

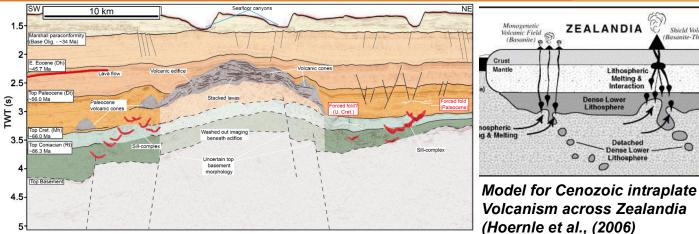
Asthenospher

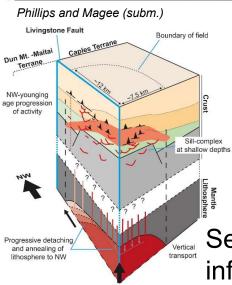
(gt. pyroxenite

0

+ peridotite)







Expression of terrane boundary at base of lithosphere focusses detaching of lithospheric material, controlling the location and lifespan of the volcanic field

Monogenetic Volcanic Field 🚱

ZEALANDIA

Dense Lowe

0

Lithospheric Melting &

Interaction

Dense Lowe 0 Lithosphere

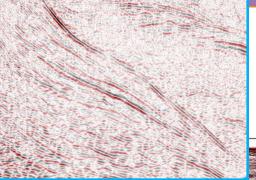
Terrane boundary acts as conduit for magma transiting the crust.

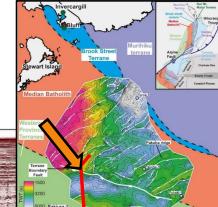
See *Phillips and Magee (in review)* for more information - https://eartharxiv.org/b94ds/

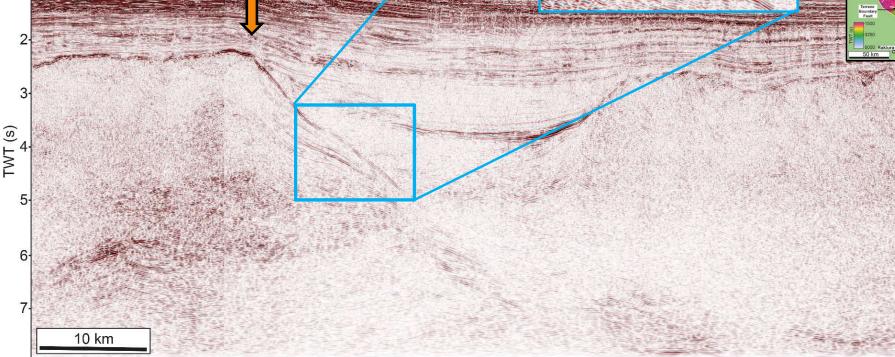
Reactivation of terrane boundaries

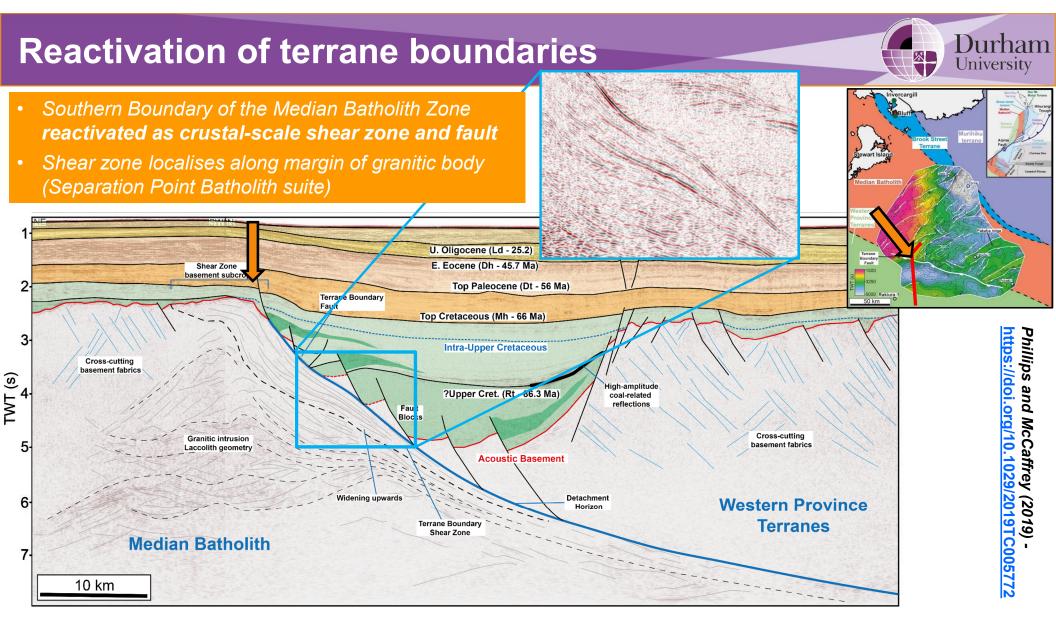


- Southern Boundary of the Median Batholith Zone
 reactivated as crustal-scale shear zone and fault
- Shear zone localises along margin of granitic body (Separation Point Batholith suite)



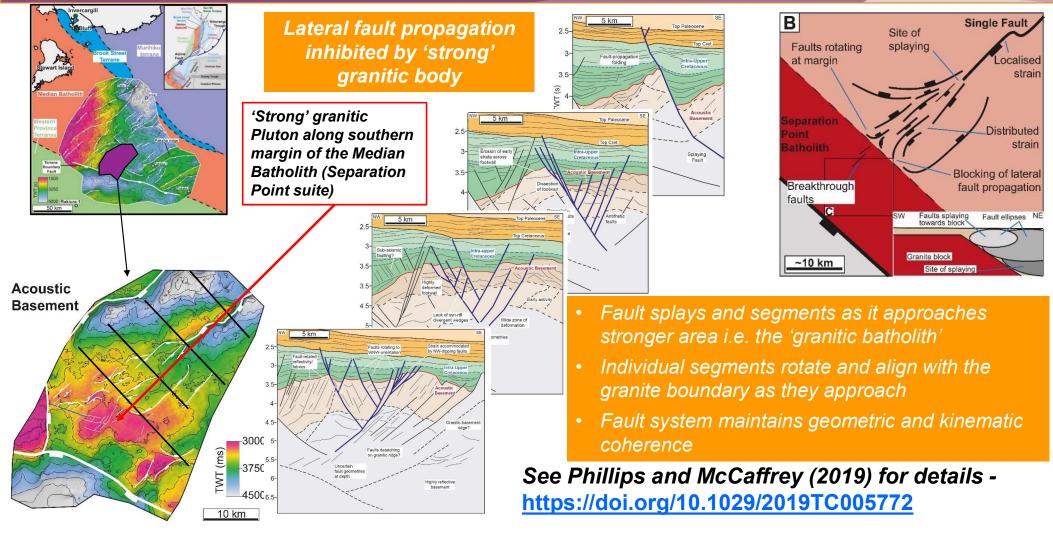






Barriers to lateral fault propagation





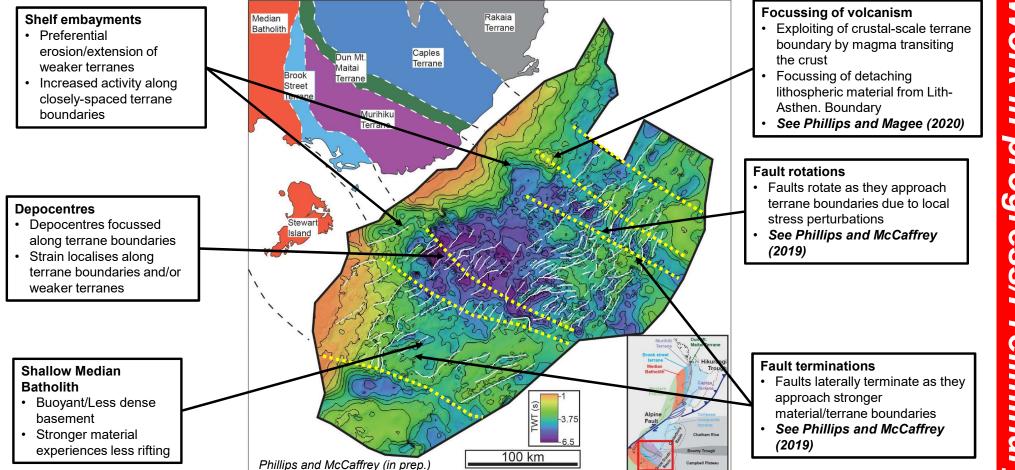


Part II – Rift structural styles across basement terranes of varying strength

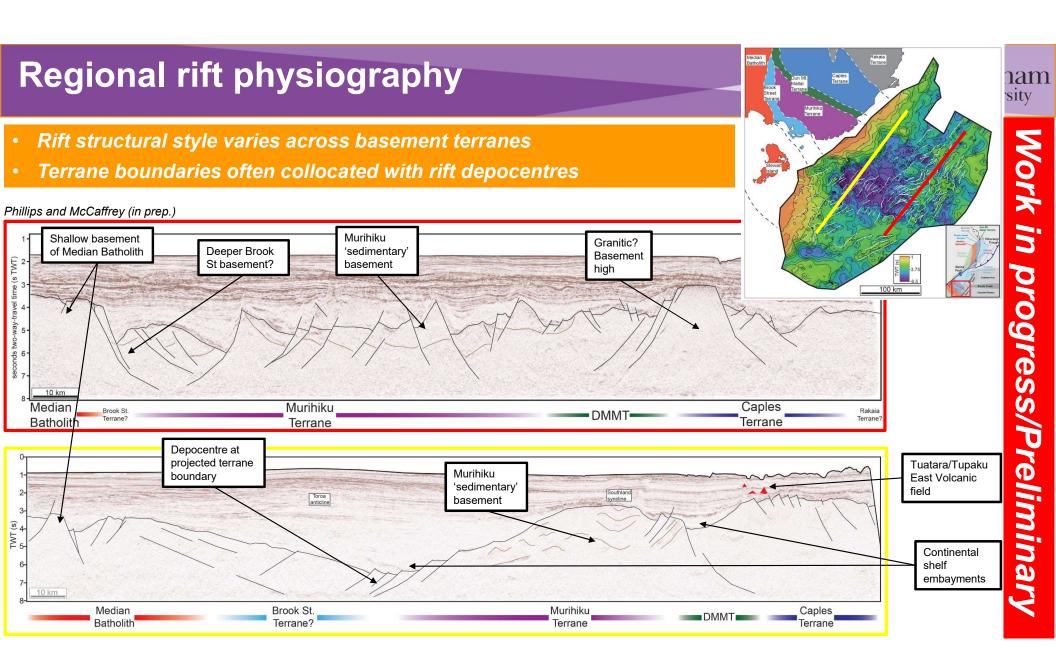
Work in progress – Quantitative analyses unable to be completed

Regional rift physiography





Work in progress/Preliminary





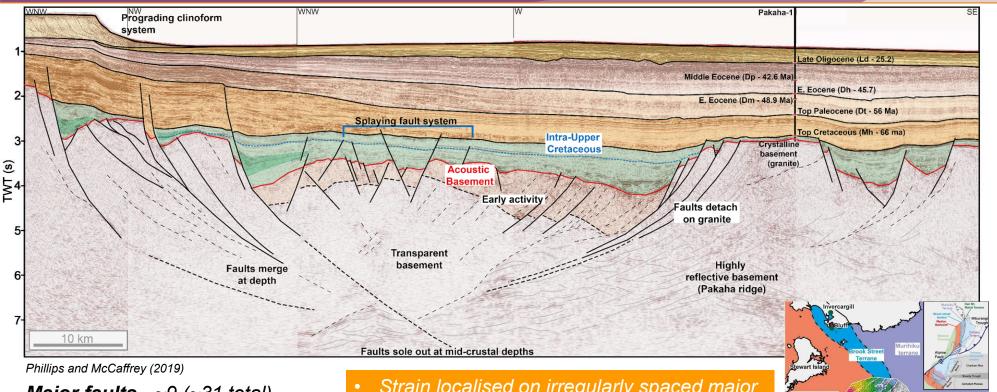
What are the prevailing structural styles associated with different basement terranes?

'Strong' case - Median Batholith



Work in

progress/Preliminary



Major faults - ~9 (~31 total) Total heave - ~10 km, ~1 km per fault (~0.33 km on all faults) Average throw - ~0.3-0.4 s TWT Beta Factor - ~1.11

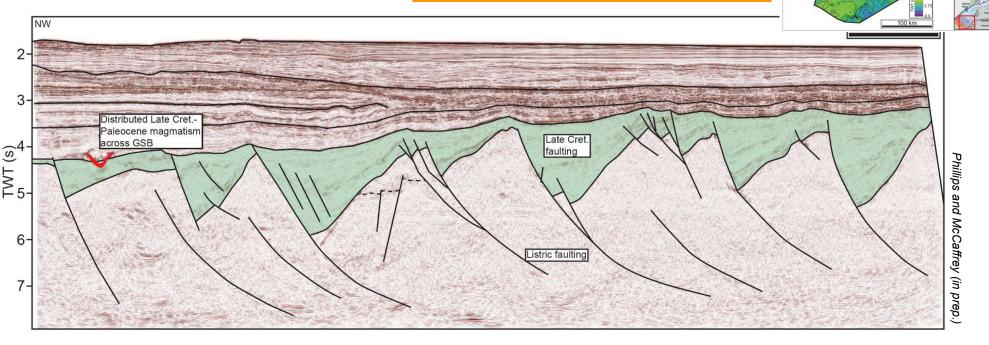
- Strain localised on irregularly spaced major faults
- Low β, dominated by minor faults?
- Location of major faults governed by granite-cored basement highs?

'Normal' case - Dun Mt.-Maitai terrane



Faults - ~8 (~0.079 faults per km) Total heave - ~25 km, ~3 km per fault Average throw - ~1-1.5 s TWT Beta Factor - ~1.32

- *Regularly spaced, SE-dipping faults*
- Listric fault geometry, indicating mid-crustal detachment (Uruski, 2010)
- High displacement faults
- Relatively high β factor



Work in progress/Preliminary

'Weak' case - Murihiku Terrane

NW-dipping

faults dominate

Mid-crustal detachment?

Murihiku

sedimentary

basement

Major faults - ~12 (28 total) (~0.08 (0.19)) faults per km)

Total heave - ~30 km, ~2.5 km per fault

Average throw - ~0.6 s TWT

Beta Factor - ~1.26

1-NW

2-

3-

(s) TWT

5-

6-

7-

- NW-dipping faults west of the basement high
- Basement characterised by sedimentary reflectivity
- Detaching faults *>* exploiting mid-crustal detachment/internal Murihiku horizon
- Unfaulted granite- and Murihiku-cored basement highs

block



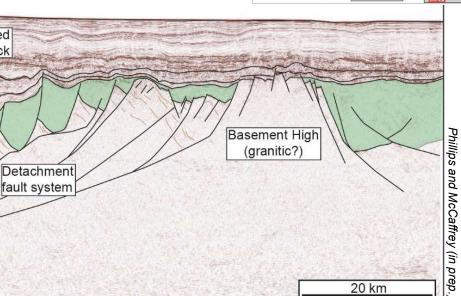
Murihiku-cored

basement high

Pre-rift

Murihiku

faulting

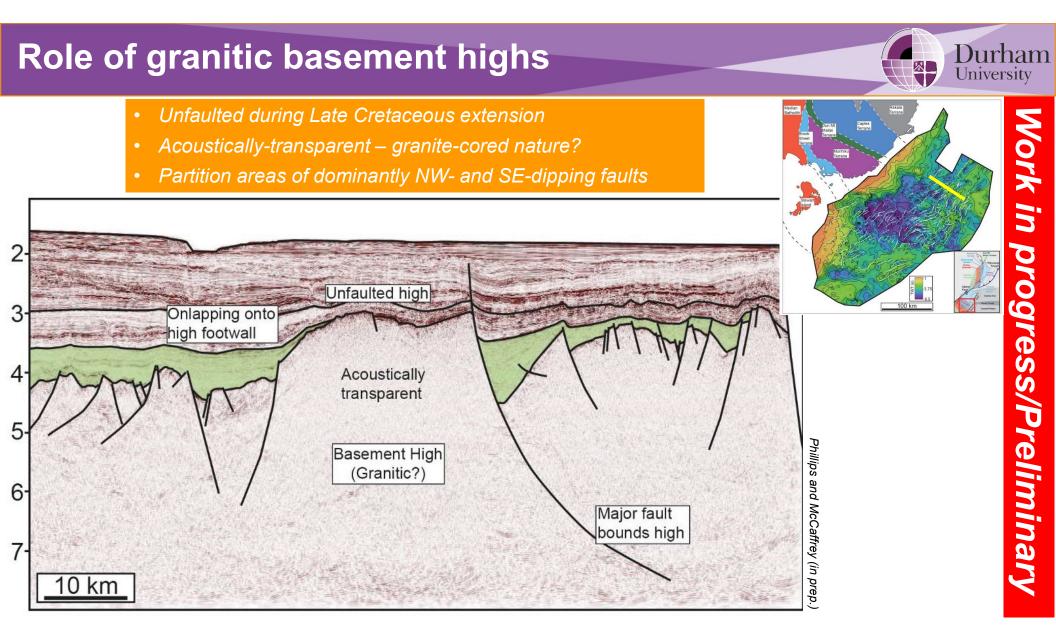


Work in progress/Preliminary

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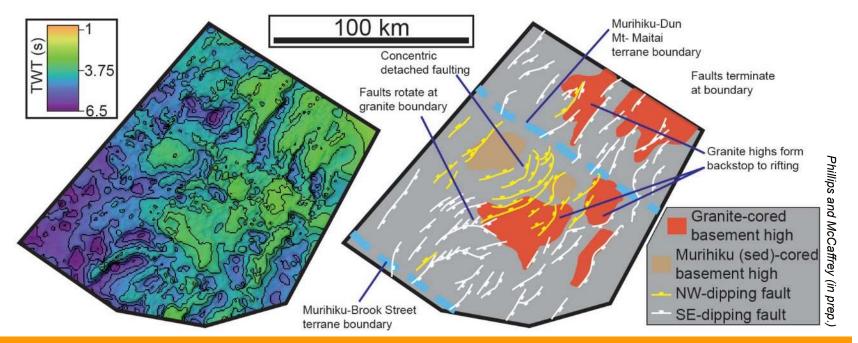
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Basement highs – Rift anchors?



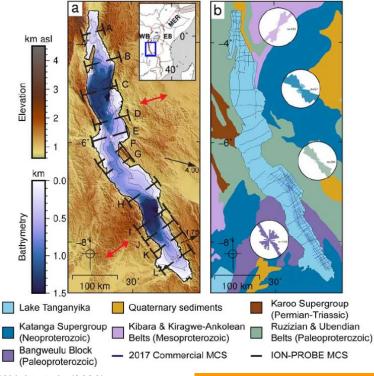


- Terrane boundaries delineated by fault terminations, fault rotations, and areas of changing fault polarity
- Relatively 'strong' granitic bodies are resistant to extension and form fixed anchors during rifting
- Faults localise along both or one of the margins of these granite-cored **anchors**, governing the dip polarity of faults on that side
- Sedimentary-cored (Murihiku) highs are transported along detaching fault systems

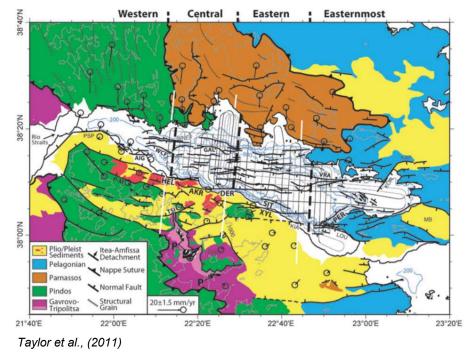
Global comparisons



Tanganyika Rift



Gulf of Corinth



Wright et al., (2020)

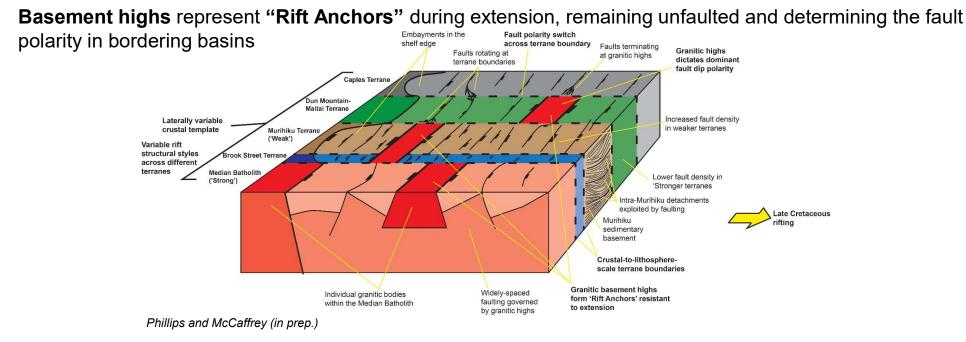
- Rifting across a laterally variable and heterogeneous basement
- Extension direction often (sub-) parallel to the boundaries between basement terranes.

Summary

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- Terrane boundaries may **reactivate** during earlier phases of extension, but **partition and segment** the rift system during the main rift phase
- Stronger areas of crustal material and terranes(i.e. granitic batholiths) are typically **resistant to faulting** and undergo seemingly less extension
- Weaker terranes experience larger strains and extension, characterised by the development of highdisplacement, listric faults



References



Phillips and Magee (in Review) - Structural controls on the location, geometry, and longevity of an intraplate volcanic system - The Tuatara Volcanic Field, Great South Basin, New Zealand - https://eartharxiv.org/b94ds/

Phillips and McCaffrey (2019) - Terrane Boundary Reactivation, Barriers to Lateral Fault Propagation and Reactivated Fabrics: Rifting Across the Median Batholith Zone, Great South Basin, New Zealand - https://doi.org/10.1029/2019TC005772

Samsu et al., <u>https://eartharxiv.org/afd7w/</u> Beniest et al., (2018) - <u>https://doi.org/10.3389/feart.2018.00148</u> Uruski, 2010 - <u>https://doi.org/10.1016/j.marpetgeo.2010.05.010</u> Mortimer et al., (2002) - <u>https://doi.org/10.1080/00288306.2002.9514978</u> Wright et al., (2020) - <u>https://doi.org/10.1029/2019TC006019</u> Taylor et al., (2011) - <u>https://doi.org/10.1111/j.1365-246X.2011.05014.x</u> Wright et al., (2016) - <u>https://doi.org/10.1016/j.earscirev.2015.11.015</u>