

# Convective self-aggregation in a mean wind

## Role of thermodynamic and dynamic processes

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# KEY POINTS

- Simulated organized convection in a mean wind using a convection-permitting model in radiative convective equilibrium
- In a mean wind, a convective cluster initially propagates, subsequently slows down and finally becomes stationary
- The surface momentum flux mainly modulates the propagation speed acting as a drag on the near-surface wind, while the surface enthalpy flux marginally contributes to the propagation speed

# Motivation

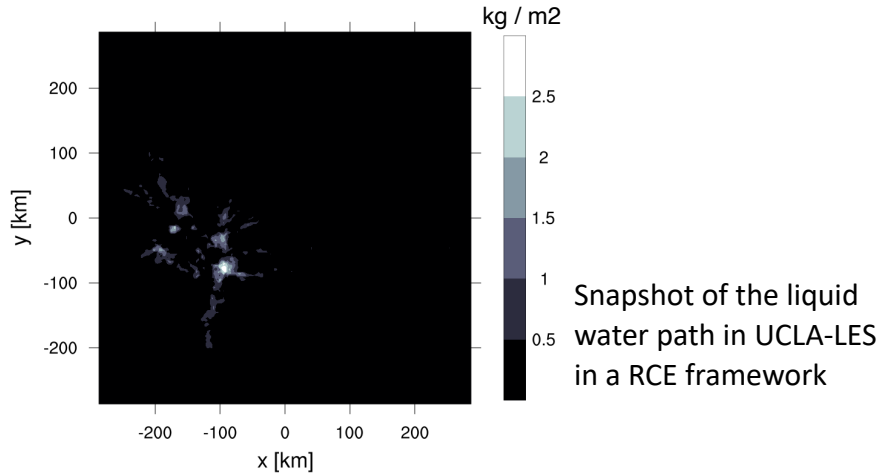


Image courtesy of the Earth Science and Remote Sensing Unit (STS005-44-1739), NASA Johnson Space Center

- Similarities between convective self-aggregation and tropical deep convection
- No large-scale background flow considered in the self-aggregation studies

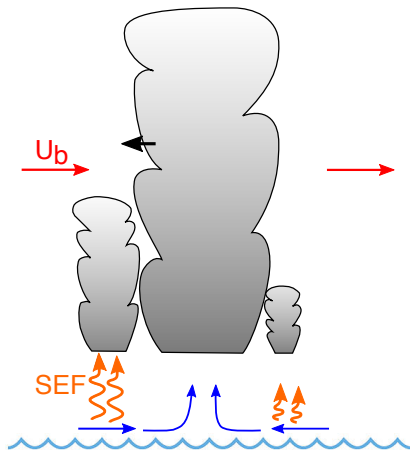
*Introduce **a large-scale background flow** to a convection-permitting simulation to analyze the behavior of convective self-aggregation in the mean flow*

# Hypothesis

Propagation of a convective cluster through the wind-induced surface heat exchange (WISHE) feedback (Emanuel, 1987)

Upwind side:

- *Intensified* wind through the low-level circulation in a mean wind  $u_b$
- *Enhanced* surface enthalpy fluxes (SEF)



Downwind side:

- *Reduced* wind through the low-level circulation in a mean wind  $u_b$
- *Suppressed* surface enthalpy fluxes (SEF)

Asymmetry of SEF → Preference of the development of convection on the **upwind** side  
→ *The convective cluster would propagate upwind!*

# Simulation setup

- UCLA-LES in radiative convective equilibrium
  - $\Delta x = 3$  km and a stretched vertical grid levels ( $\Delta z_{\min} = 75$  m)
  - on a domain size of  $576 \times 576 \times 27$  km<sup>3</sup>
  - following Hohenegger and Stevens (2016)
- No convective parameterization
- Sub-grid fluxes with Smagorinsky model
- Homogeneous initial conditions
- Cyclic boundary conditions
- No Coriolis force, no diurnal cycle
- SST = 301 K

Research focus:

How a convective cluster behaves in a mean wind.

→ Simulation is run for **26 days** until convection is self-aggregated.

*After day 26 ..*

# Impose a mean wind

- A mean wind  $u_b$  is imposed through the surface flux calculation to minimize the vertical shear

Momentum flux  $F_m = \rho (\overline{w'u'^2} + \overline{w'v'^2})^{\frac{1}{2}}|_{\text{sfc}}$

Surface enthalpy flux  $\text{SEF} = \rho (c_p \overline{w'\theta'} + l_v \overline{w'q'})|_{\text{sfc}}$

$$\overline{w'u'^2} + \overline{w'v'^2} = u_*^2$$

$$\overline{w'\theta'} = -u_*\theta_*$$

$$\overline{w'q'} = -u_*q_*$$

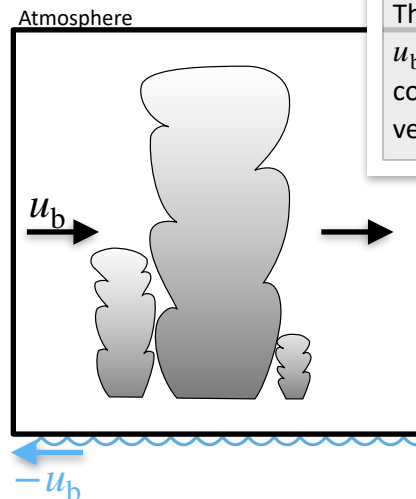
The turbulence scale of velocity  $u_*$  is proportional to the near-surface horizontal velocity  $u_h$ .

$$u_h = \sqrt{(u + u_b)^2 + v^2}$$



Each experiment is run for 20 days.

Experiment	$u_b [m s^{-1}]$
UB0	0
UB2	2
UB4	4
UB2_unius	2

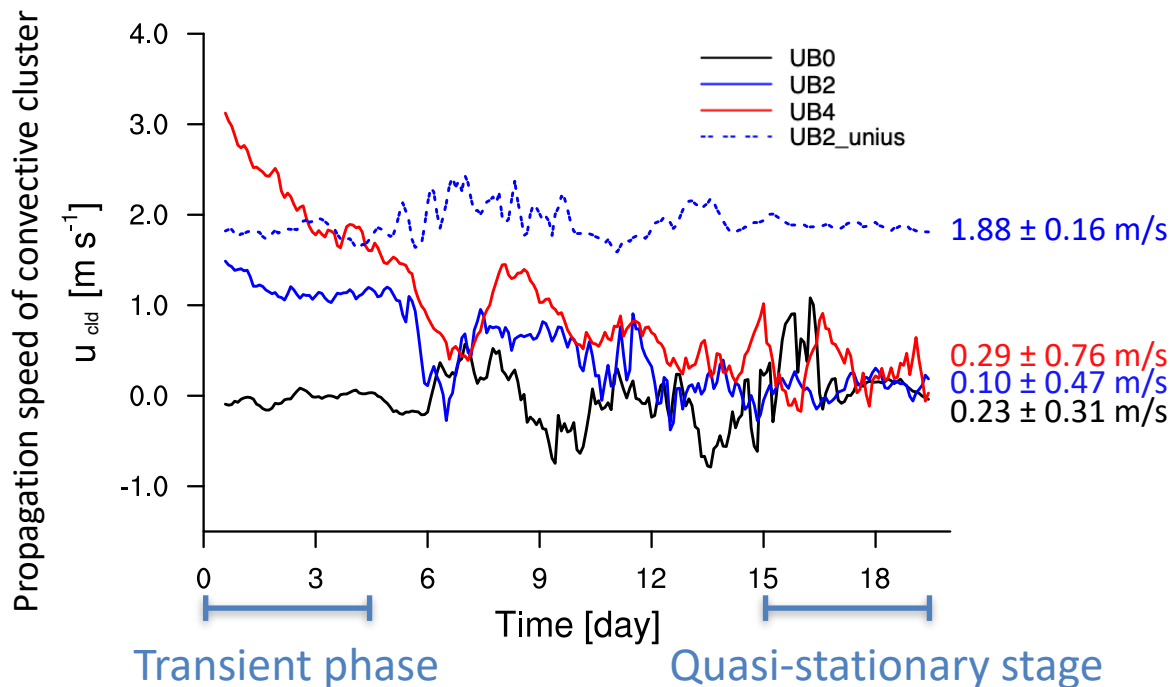


The modeled flow  $u$  feels  $u_b$  as if  $u$  is on a moving conveyor belt with a velocity of  $-u_b$

UB2\_unius:

$u_* = 0.09 \text{ m/s}$  in  $F_m$  (not in SEF!)

# How does a convective cluster propagate in the mean wind?



\*Note that here day 0 indicates the time when the mean wind begins to be imposed.

- Transient phase:

$$u_{\text{cld}} < u_b$$

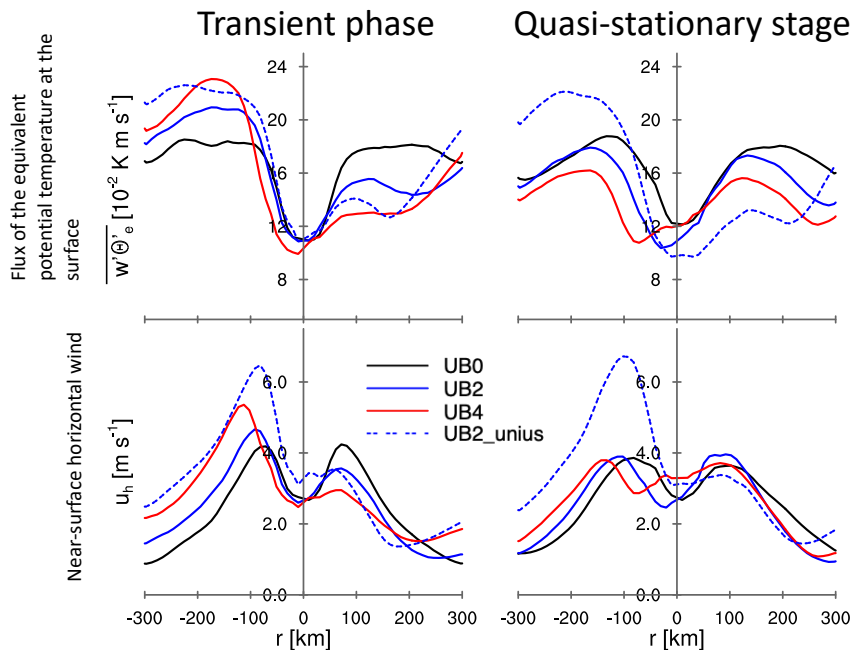
→ Convection propagates upwind!

- Quasi-stationary stage:

$$u_{\text{cld}} = 0 \text{ m/s}$$

*Why does propagation speed arrive at 0 m/s regardless of the mean wind speed???*

# Thermodynamic response to the mean wind



$r=0$ , center of a convective cluster  
 $r<0$ , upwind side  
 $r>0$ , downwind side

\*Note that  $\overline{w'\theta'_e}$  is a indicator of *SEF*

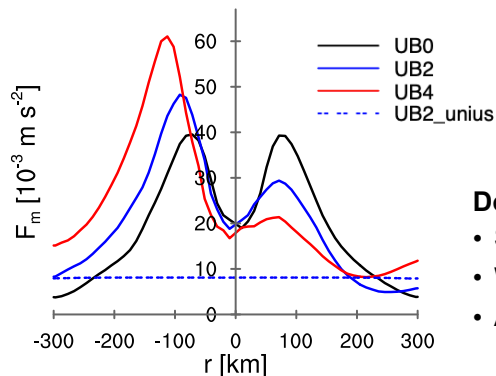
$$\overline{w'\theta'_e} \approx \left[ \overline{w'\theta'} + \frac{l}{c_p} \left( \frac{p_0}{p} \right)^{\frac{R_{cp}}{c_p}} \overline{w'q'} \right]$$

- **Asymmetry** w.r.t. the center of the convective cluster in the transient phase
  - Asymmetry in  $\overline{w'\theta'_e}$  agrees with the hypothesis.
  - Same spatial distribution in  $u_h$  but not in  $\Delta q$  or  $\Delta\theta$
- **Symmetry** in the quasi-stationary stage
  - How do the asymmetries become symmetric?
  - Why is a convective cluster NOT advected in a mean wind, instead stationary?



# Role of the surface momentum flux $F_m$ acting as a drag

Transient phase



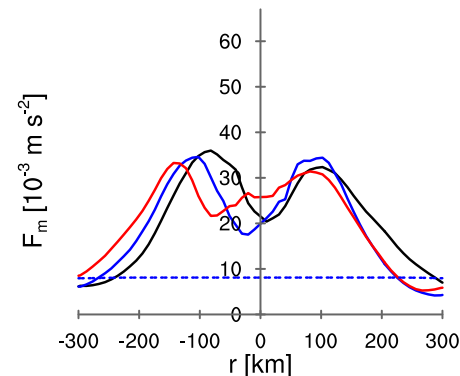
**Upwind side:**

- Enhanced  $F_m$
- Strong drag on  $u_h$
- Decrease in  $u_h$

**Downwind side:**

- Suppressed  $F_m$
- Weak drag on  $u_h$
- Attaining  $u_h$

Quasi-stationary stage



→ the near-surface horizontal wind  $u_h$  has the comparable maximum on the upwind and downwind sides.

**Symmetrized!**

→ Symmetry in  $u_h$  results in the surface turbulent fluxes ( $F_m$ ,  $SEF$ ) in the quasi-stationary stage

# Does the momentum flux dominantly modulate the near-surface wind?

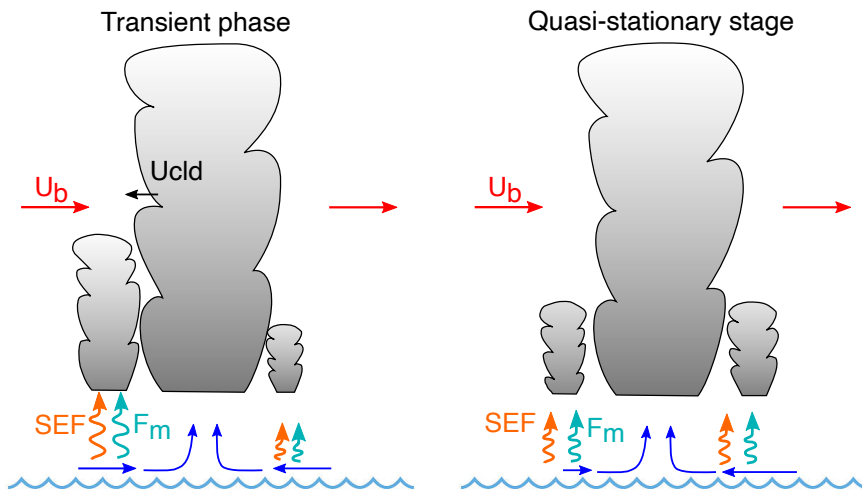
UB2\_unius

- Experiment to suppress the dynamic feedback modulating the interaction between the surface momentum flux  $F_m$  and the near-surface horizontal wind  $u_h$
- Propagation speed:  $u_{\text{cld}} = 1.88 \text{ m/s}$ 
  - smaller than the mean wind speed  $u_b = 2 \text{ m/s}$
  - convective cluster indeed propagates upwind
- Asymmetry of the thermodynamic flux  $\overline{w'\theta'_e}$  in the transient phase and the quasi-stationary stage

*Yes, the thermodynamic response to the mean wind contributes to the propagation speed to a small extend.*

# Conclusion

- In a mean wind, the convective cluster propagates upwind, slows down, and eventually becomes stationary.
- The dynamic response to the mean wind dominates the propagation as the surface momentum flux acts as a drag on the near-surface wind, which consequently feeds back moderating surface turbulent fluxes.
- The thermodynamic feedback, through WISHE, is very weak and is regulated by the dynamic feedback.



# References

- Emanuel, K. A.: An air–sea interaction model of intraseasonal oscillations in the tropics, *J. Atmos. Sci.*, 44, 2324–2340, 1987. DOI: 10.1175/1520-0469(1987)044<2324:AASIMO>2.0.CO;2
- Hohenegger, C. and Stevens, B.: Coupled radiative convective equilibrium simulations with explicit and parameterized convection, *J. Adv. Model. Earth Syst.*, 8, 1468–1482, 2016. DOI: 10.1002/2016MS000666