

nc

Impact of using different prior flux products on NEE estimates derived from the Jena Carboscope regional inversion system

S. Munassar^[1], T. Koch^[1,2], C. Rödenbeck^[1], and C. Gerbig^[1]

For independent verification of CO₂ flux budgets at annual and national scales over Europe, within the research project VERIFY, flux estimates of CO₂ have been calculated for the period 2006-2018 using the Jena CarboScope Regional inversion system (CSR). Based on prior knowledge of CO₂ fluxes, the NEE is optimized against observational datasets of CO₂ dry mole fractions collected through the station network of the Integrated Carbon Observation System (ICOS) across the European domain. To distinguish the impact of using different terrestrial biosphere models on posterior NEE, the Vegetation Model (VPRM), the Simple Biosphere/Carnegie-Ames Standford Approach (SiBCASA), and FLUXCOM model are assimilated in ensemble inversions in the CSR. Moreover, Mikaloff-Fletcher et al. (2007) and Jena CarboScope pCO₂-based ocean fluxes are used as various ocean flux models in the ensemble inversions. CO₂ national emission inventories are provided from EDGAR_v4.3 and updated based on BP statistics. Results from the ensemble inversion runs utilizing different terrestrial biosphere models show agreement in the estimated interannual variability (IAV) of NEE, despite a large

difference on the prior NEE annual budgets of biosphere models. A much smaller impact is observed when applying different ocean flux models, in particular for regions far inland.

Innovation fluxes for 2018 (VPRM)

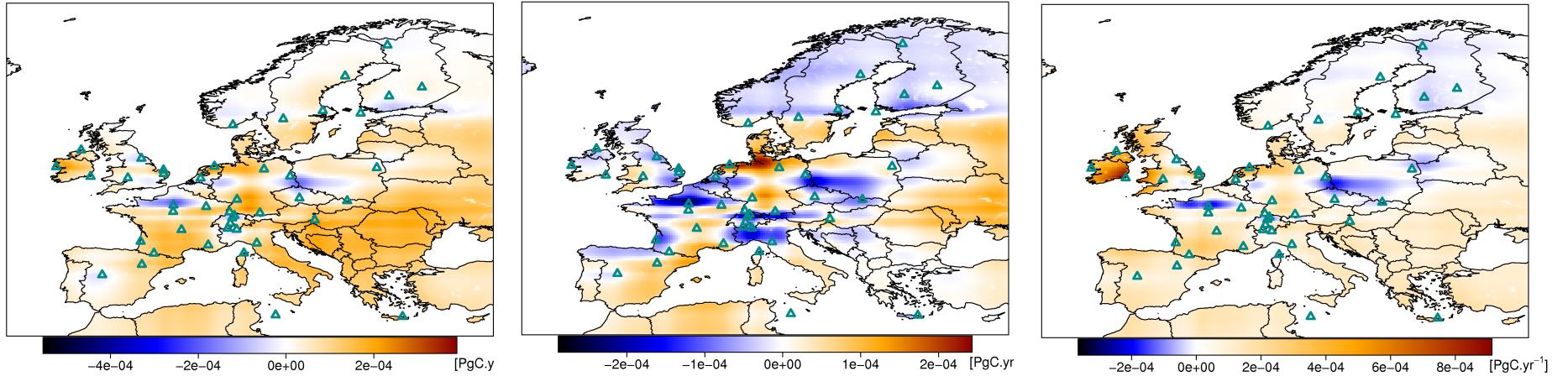
Innovation fluxes for 2018 (SIBCASA)

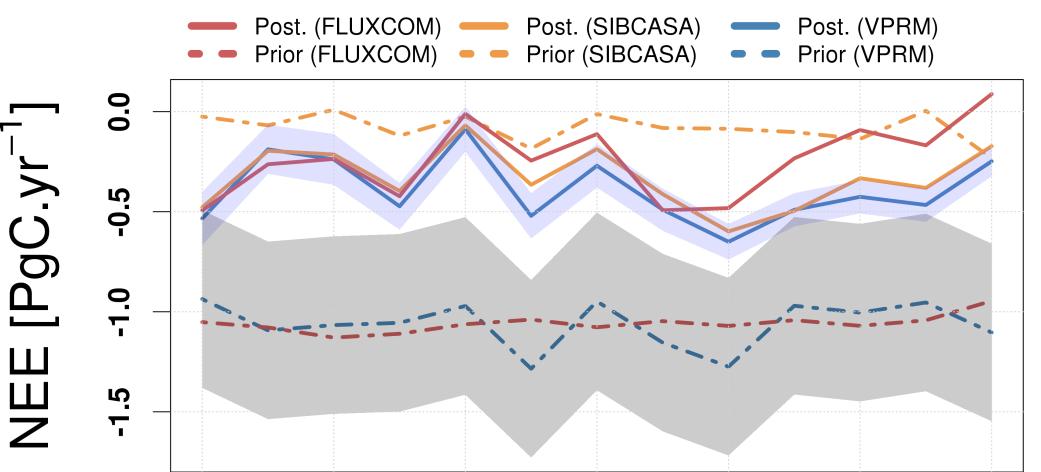
Innovation fluxes for 2018 (FLUXCOM)

all Europe

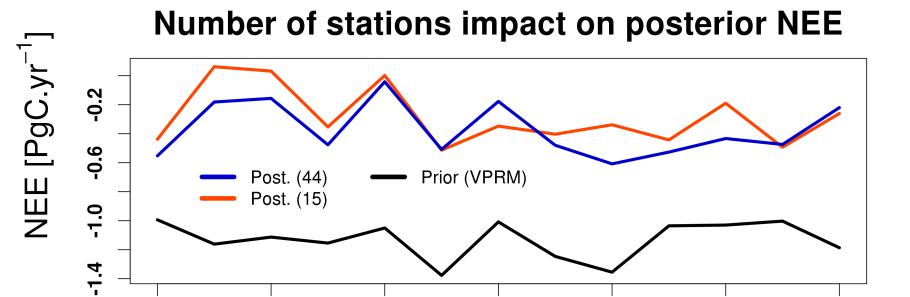
Max Planck Institute

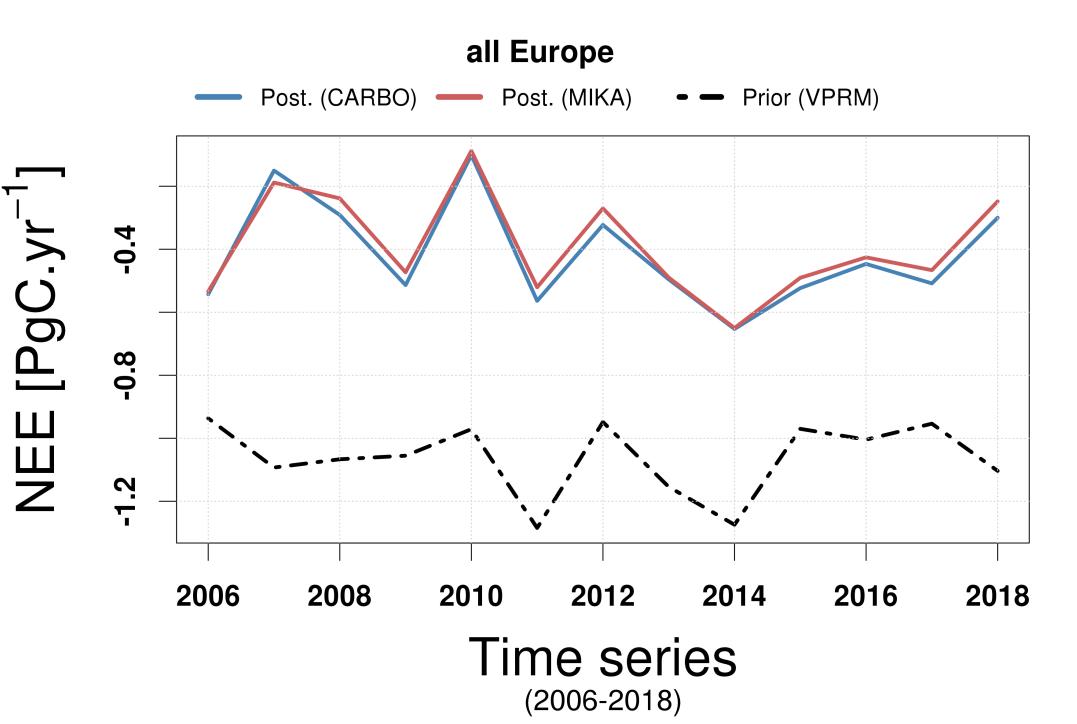
for Biogeochemistry





- Innovation of fluxes for 2018 depicted on maps outlines the difference between posterior fluxes and the apriori calculated from VPRM, SIBCASA, and FLUXCOM models at spatial resolution of 0.5 degree
- IAV of posterior NEE is consistent for all biosphere models despite inconsistency over prior flux models (Figure 1, above); a quite uncertainty reduction realized over the a-posteriori for VPRM (shadow of lines)
- Quite a small impact appears on posterior NEE when using different ocean flux model (Figure 1, below)
- As the main constraint of prior fluxes, observational data have a major impact in estimating posterior fluxes.
- Corresponding figure shows estimated fluxes from two inversion runs assimilating 44 stations (blue line) and 15





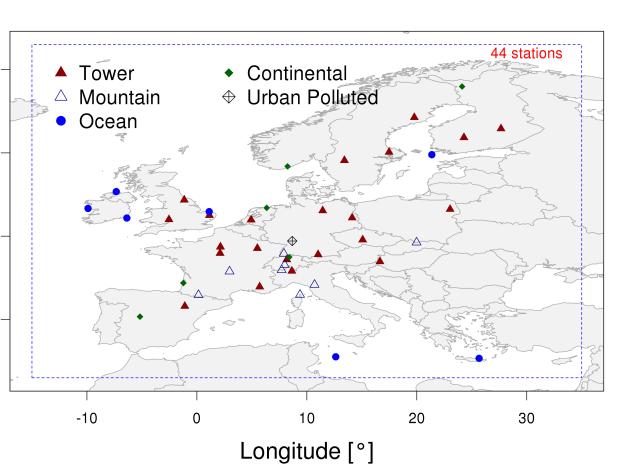
stations (red line) across Europe

2018 2006 2008 2010 2012 2014 2016

Time series

Figure 1: Impact of applying various prior flux models of biosphere (above) and ocean (below)

Station network Measurements of CO₂ dry ▲ Tower Continental Mountair mole fraction • Ocean Latitude [°] 44 stations Coverage over the domain of Europe Longitude [°] **Station classifications** [mdd] က \sim UP Classificatior station location



Station classifications

The representation uncertainty of atmospheric transport model defined weekly based on

T: Tall Tower M: Mountain C: Surface S: Ocean UP: Urban Polluted

observations In contrast to SIBCASA and FLUXCOM, VPRM has a

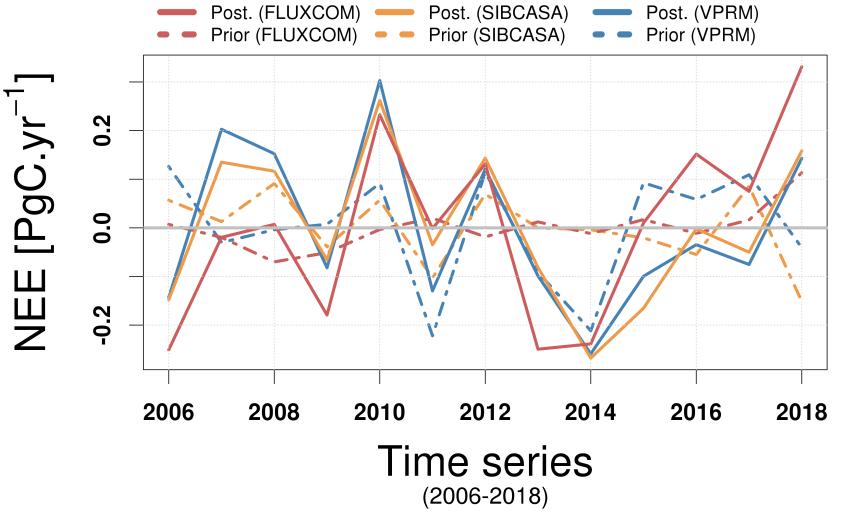
better representation of IAV



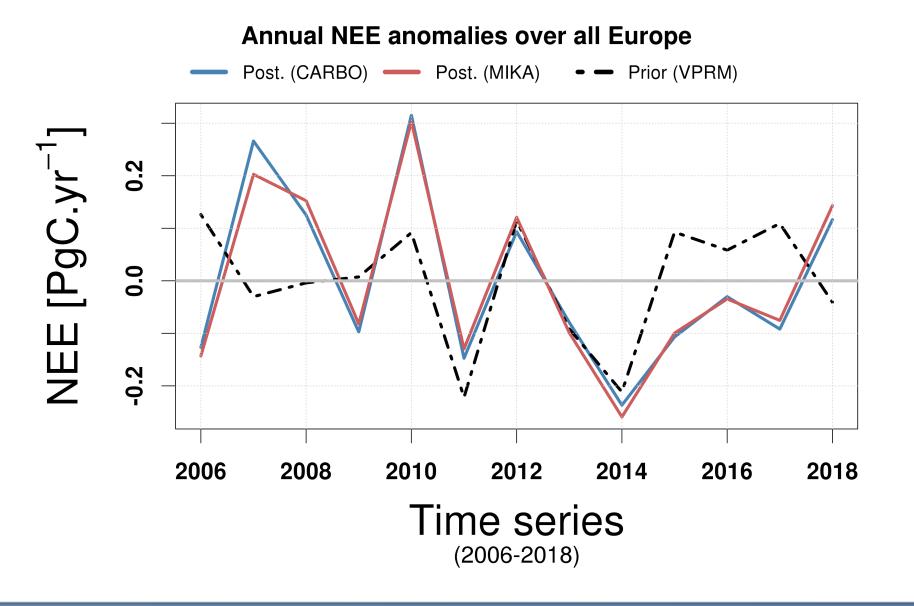
Posterior fluxes of biosphere models show a

stronger anomaly signal, likely captured from

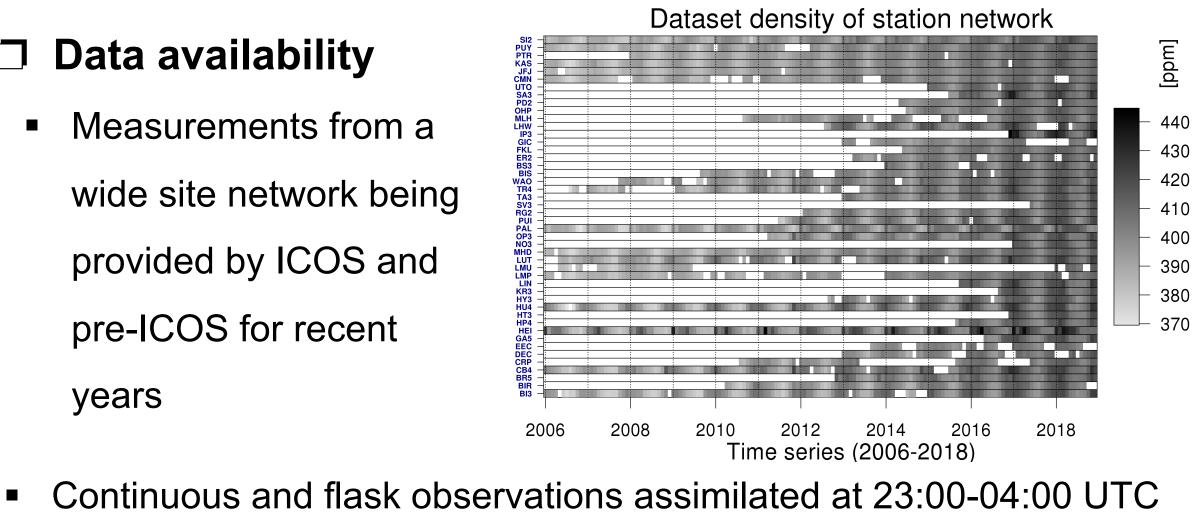
□ NEE anomalies relative to long-term mean of 2006-2018



 Using different ocean flux models has a smaller impact on total CO2 estimate over domain-wide in comparison with biosphere models



uo ati Obser



for mountain sites and at 11:00-16:00 for the rest

- Hourly STILT footprints calculated over receptors (stations) Prior flux models Diagnostic biogenic terrestrial models (VBRM, SiBCASA, FLUXCOM) \diamond
 - Ocean flux model (Mikaloff-Fletcher et al. (2007), Jena CarboScope pCO₂-based ocean fluxes) \diamond
 - Fossil fuel emission statistics (EDGARv4.3 updated according to BP) \diamond
- Lateral Boundary Conditions (LBCs)
 - Global TM3 model \diamond

20

0

VERIFY

- Two-step scheme inversion
- 100 km spatial correlation length of prior error, hyperbolic decay 0
 - 30 day temporal correlation length of prior error







[1] ATM group, Biogeochemical Systems Department, Max-Planck Institute for Biogeochemistry, Jena [2] Meteorological Observatory Hohenpeissenberg, Deutscher Wetterdienst, Germany

*Email: smunas@bgc-jena.mpg.de

Affiliations:

MAX-PLANCK-GESELLSCHAFT