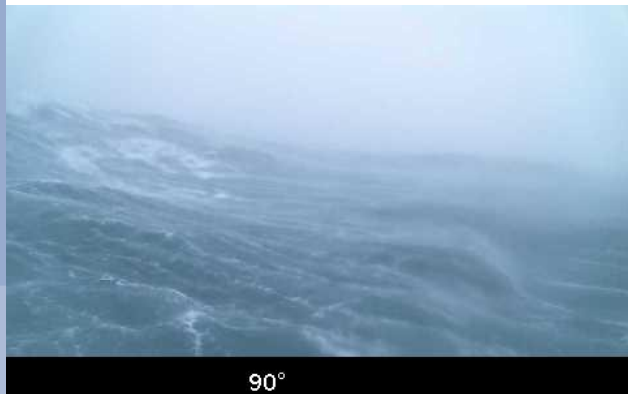


Sea-state dependency of air-sea fluxes for high winds in ECMWF Earth System Model

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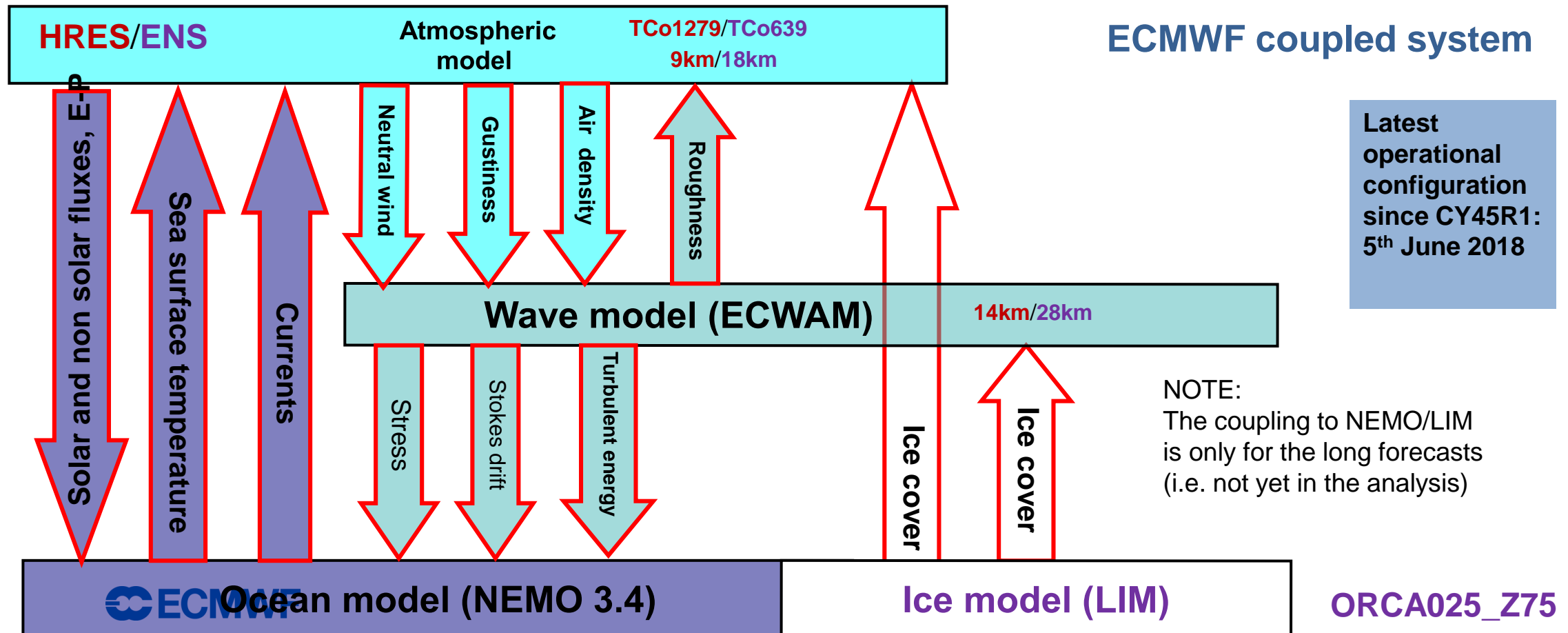


NDBC buoy 41004, Edisto, 76 km Southeast of Charleston, South Carolina, **during** and after Hurricane Dorian

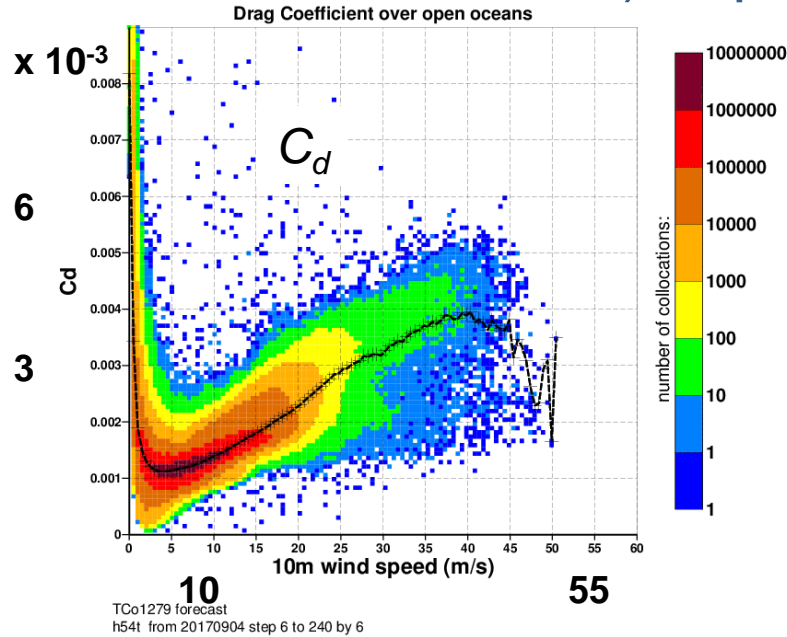
➤ Follow me as I detail what was written in the abstract...



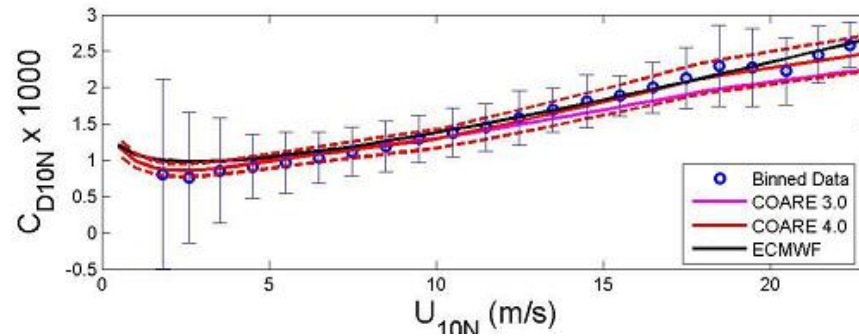
- The global analyses and medium range forecasts from the European Centre for Medium range Weather Forecasts rely on a state-of-the-art Numerical Weather Prediction (NWP) system.
- To best represent the air-sea exchanges, it is tightly coupled to an ocean wave model.
- As part of ECMWF approach to Earth System Model, it is also coupled to a global ocean model for all its forecasting systems from the medium range up to the seasonal time scale.



1) Impact of ocean waves on the surface stress



Drag coefficient (C_d)
with respect to wind speed
Operational model (CY46R1)



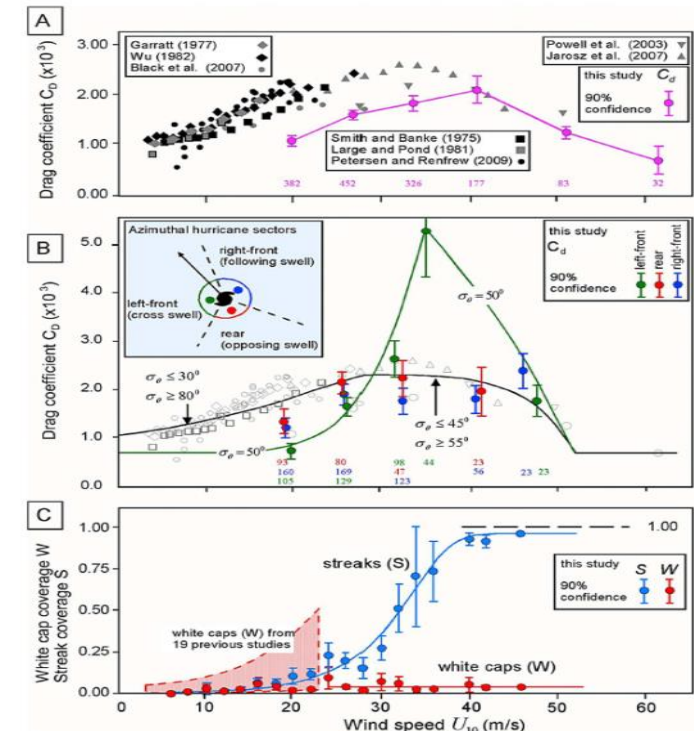
Edson et al., 2013

C_d fits well observations
for winds up to 20m/s
But it might be too high
for larger winds

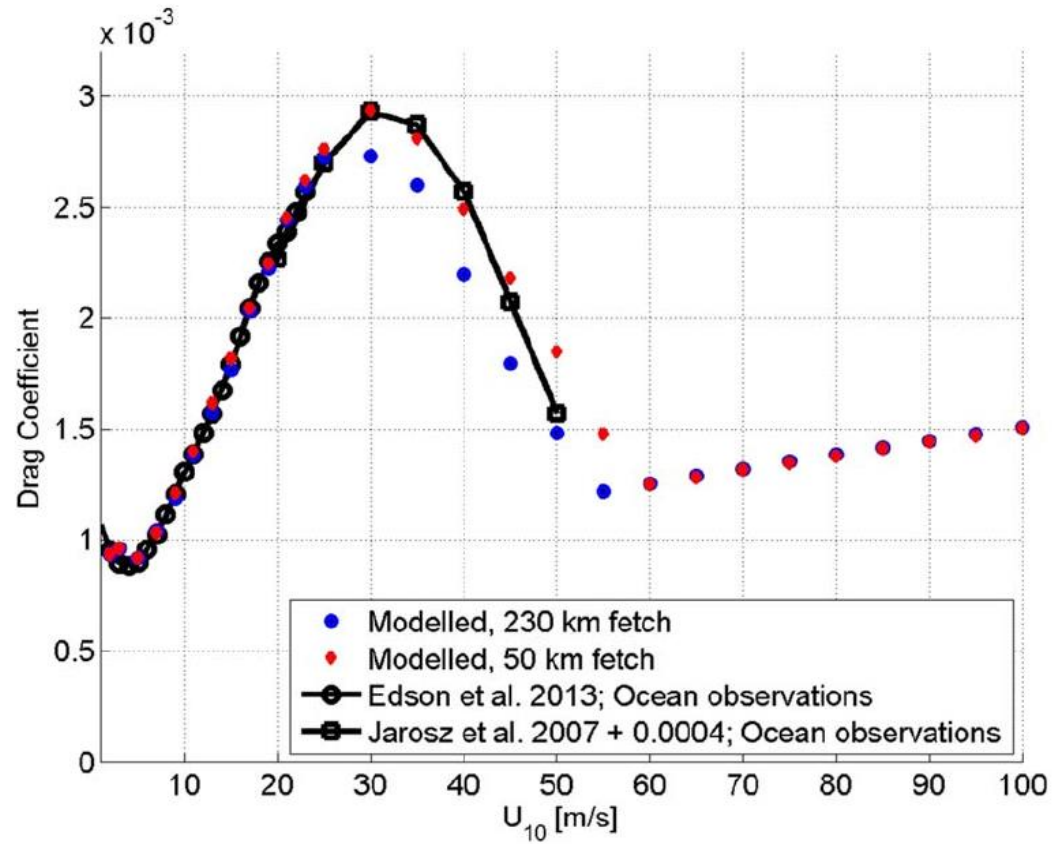
Holthuijsen et al. , 2012

C_d is sea state dependent !

- Because the feedback from and to the ocean can be significant, it is only in the fully coupled system that parameterisation for air-sea processes should be revisited.
- It is now accepted that the drag coefficient should generally attained maximum values for storm winds but should level or even decrease for very strong winds, namely in tropical cyclones or intense mid-latitude wind storms.



Cd(U10)

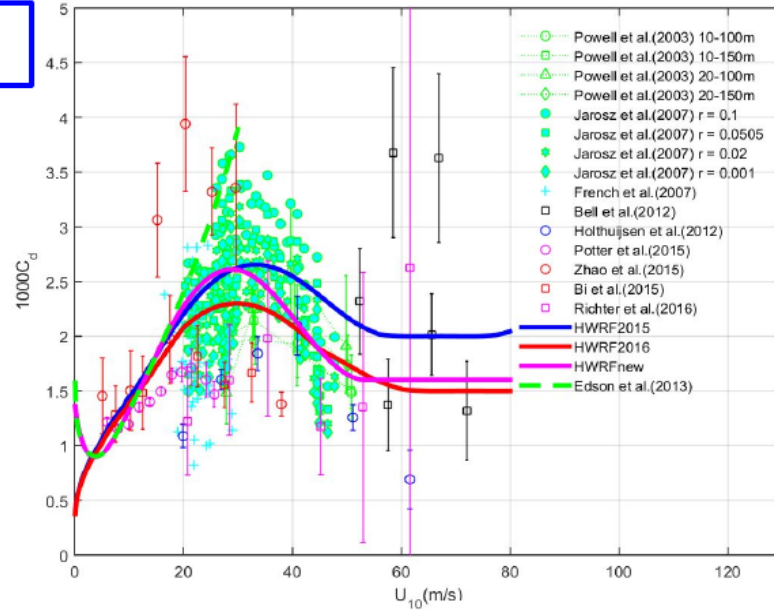


Donelan (2018)

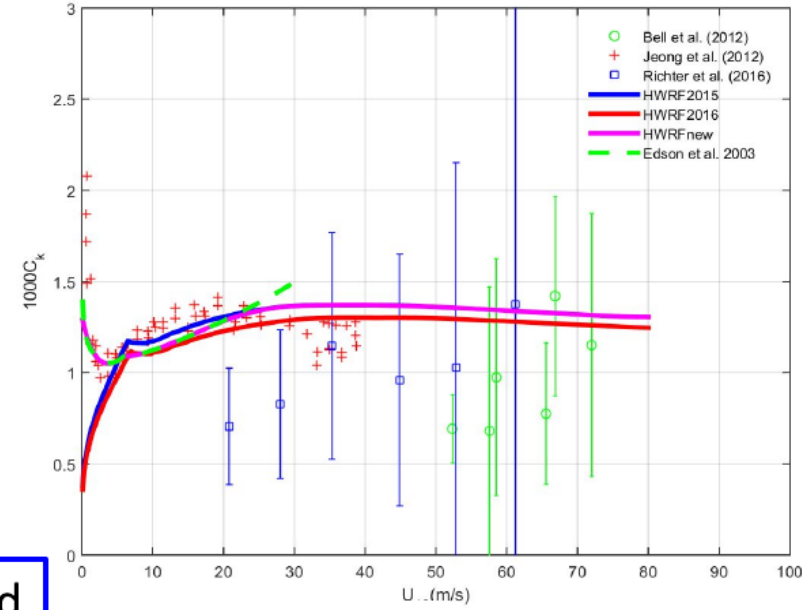
Figure 6. Modeled (with Reynolds number dependent sheltering coefficient, Figure 5, equation (13)) and observed drag coefficients versus wind speed. Note the fetch dependence of the modeled drag coefficient between 30 and 55 m/s.

Cd(U10) : HWRF

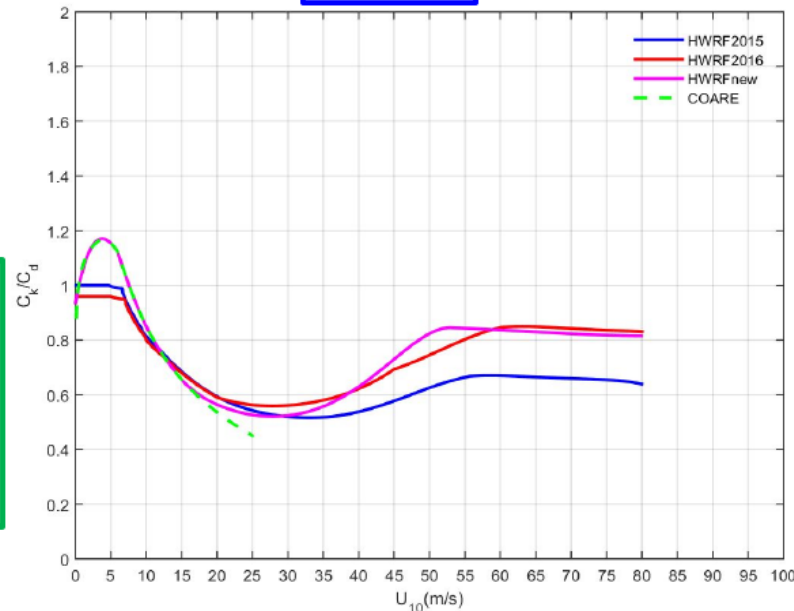
Cd



Ck



Ck/Cd



Magenta: FY2017 HWRF
Red: FY2016 HWRF
Blue: FY2015 HWRF

Low to moderate winds:
Using COARE algorithm
High winds:
Fitting to observed Cd

Emanuel's MPI theory:

$$V_{max}^2 \propto \frac{Ck}{Cd} \times \frac{T_s - T_o}{T_s} \times (k_0^* - k_a)$$

- A modification of the wind input source was tested, whereby the Charnock coefficient estimated by the wave model and therefore the drag coefficient sharply reduce for large winds (> 30 m/s).

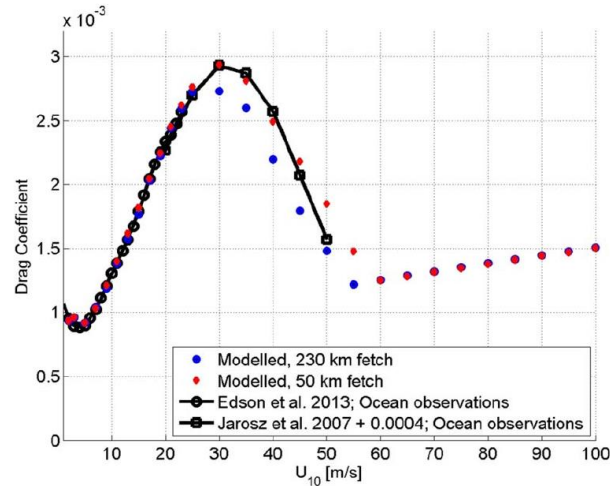


Figure 6. Modeled (with Reynolds number dependent sheltering coefficient, Figure 5, equation (13)) and observed drag coefficients versus wind speed. Note the fetch dependence of the modeled drag coefficient between 30 and 55 m/s.

Donelan (2018)

With the wave model, Charnock is expressed as

$$\alpha = \frac{\tilde{\alpha}}{\sqrt{1 - \frac{\tau_w}{u_*^2}}}$$

with

$$\tilde{\alpha} = \tilde{\alpha}_0 = 0.0065$$

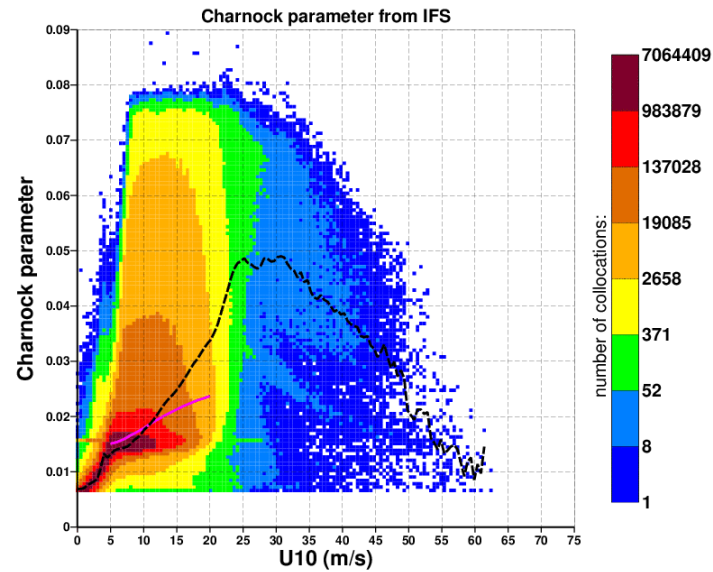
In order to mimic Donelan Cd, Charnock has to reduce quite sharply for winds (U_{10}) above 33m/s and then tails off for very high winds:

$$\tilde{\alpha} = \tilde{\alpha}_{\min} + 0.5(\tilde{\alpha}_0 - \tilde{\alpha}_{\min})[1 - \tanh(U_{10} - 33)]$$

with

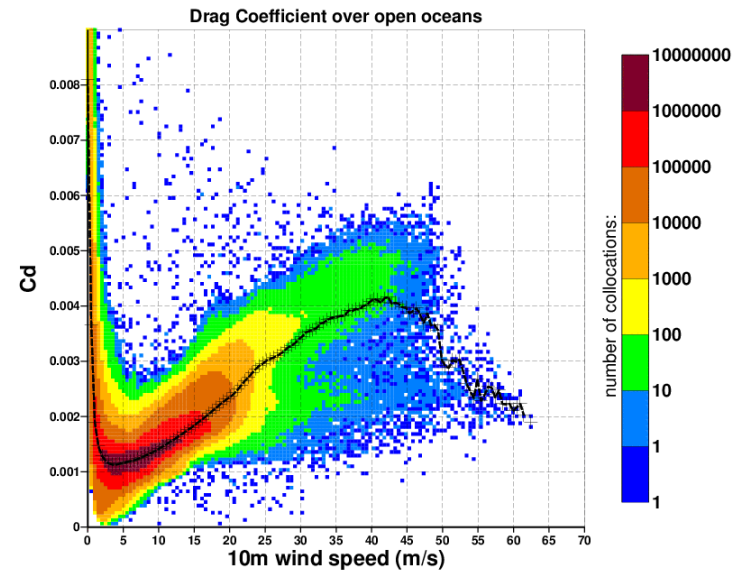
$$\tilde{\alpha}_{\min} = 0.0001$$

Irma, forecast from 20170904, 0 UTC: with CY46R1 at Tco1999 (5km)



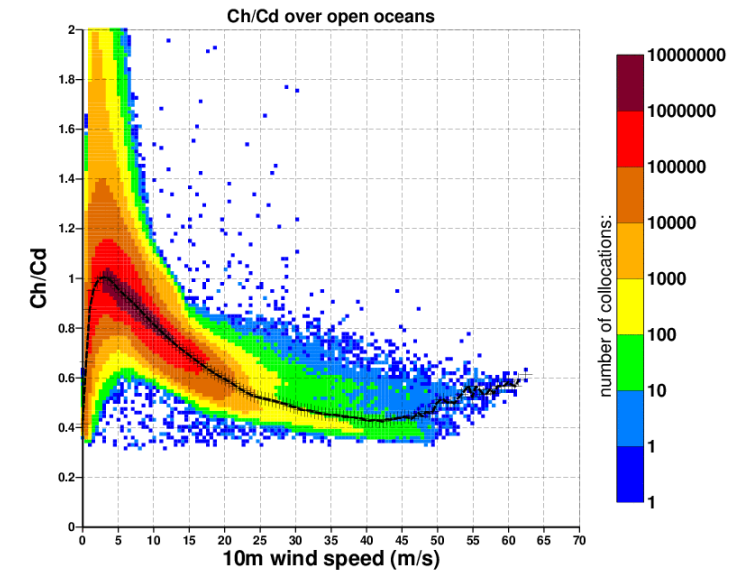
Forecast data from stream da, class rd, expver h5jx, all Sea points with sea ice cover <= 0.3
from 20170904 00UTC, for steps from 6 to 240 by 6

Charnock v U10



TCo1999 forecast
h5jx from 20170904 step 6 to 240 by 6

Cd v U10

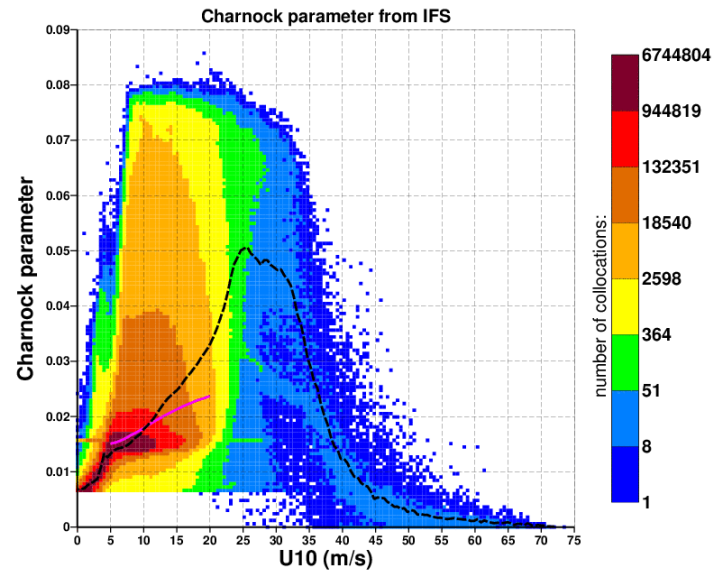


TCo1999 forecast
h5jx from 20170904 step 6 to 240 by 6

Ch/Cd v U10

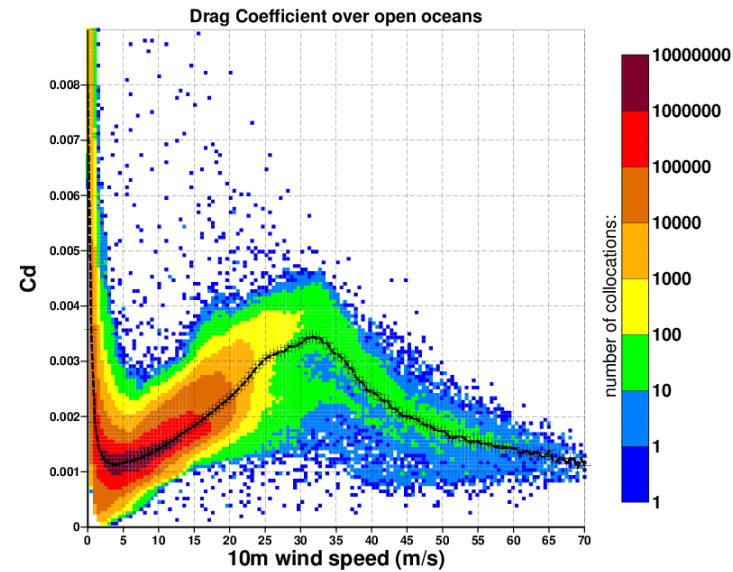
With CY46R1 (2019) wave physics parameterisation

Irma, forecast from 20170904, 0 UTC: CY46R1 Tco1999 + limitation on Charnock



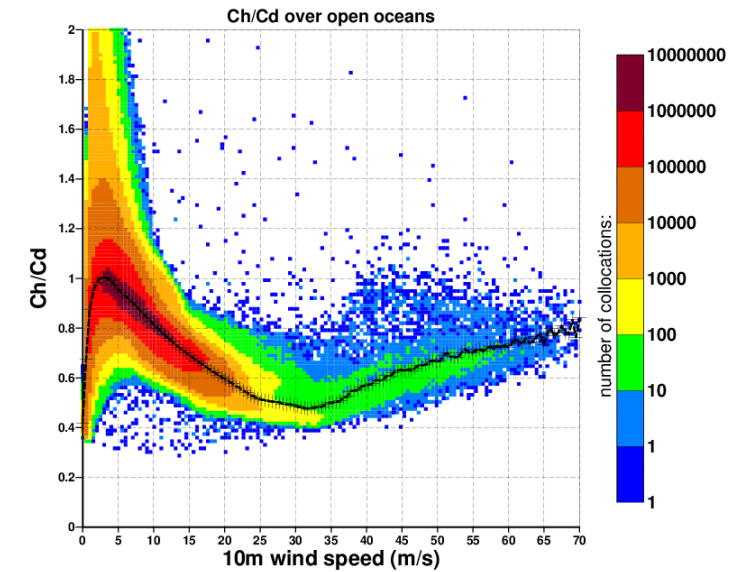
Forecast data from stream da, class rd, expver h5i7, all Sea points with sea ice cover <= 0.3
from 20170904 00UTC, for steps from 6 to 240 by 6

Charnock v U10



TCo1999 forecast
h5i7 from 20170904 step 6 to 240 by 6

Cd v U10

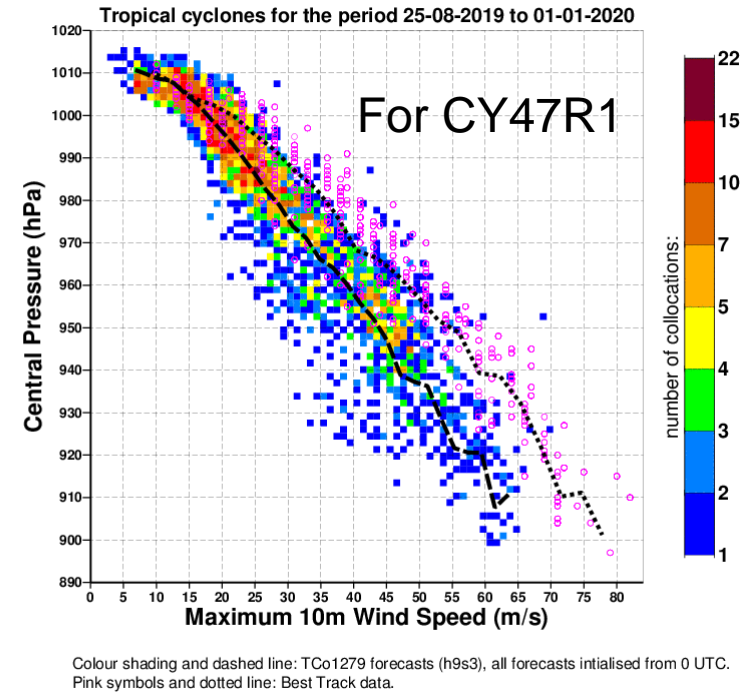
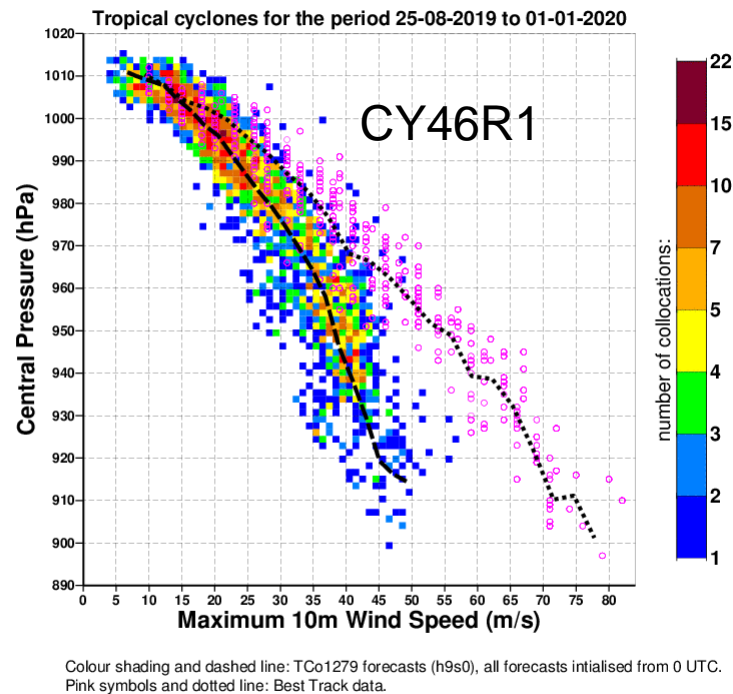


TCo1999 forecast
h5i7 from 20170904 step 6 to 240 by 6

Ch/Cd v U10

With CY46R1 (2019) wave physics parameterisation

Tropical cyclone max wind - min pressure relationship



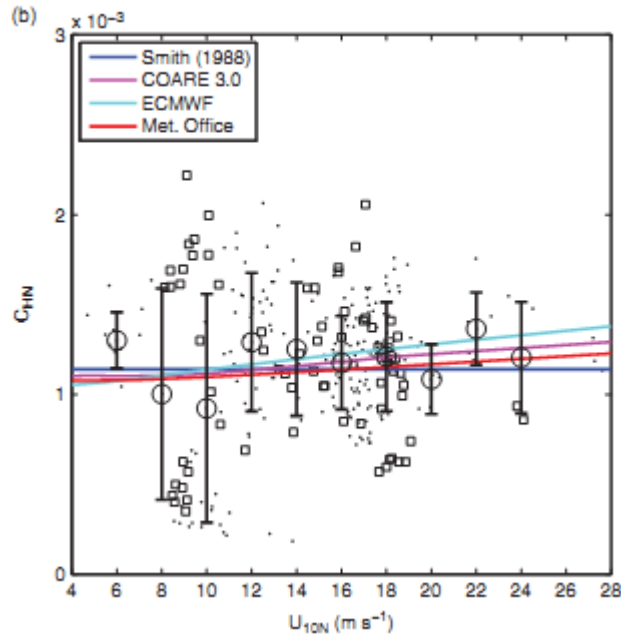
Tco1279 forecasts from 0 UTC for period 25-08-2019 to 01-01-2020 (coloured shading and dotted line). Reported values (pink symbols and dotted line) for tropical cyclones: Ambali, Belna, Bualoi, Calvinia, Dorian, Faxai, Fengshen, Hagibis, Halong, Humberto, Kammuri, Kyarr, Lingling, Lorenzo, Maha, Matmo, Nakri, Phanfone, Sarai, Sebastien.

- As a consequence, ECMWF tendency to under predict strong tropical cyclones was sharply alleviated, in better agreement with observational evidence.
- This change is now planned for operational implementation with the next model cycle (CY47R1, end of June 2020).

Future developments

2) Impact of ocean waves on heat and moisture fluxes

- Experimental evidences also point to a sea state/wind dependency of the heat and moisture fluxes.



Cook and Renfrew 2014

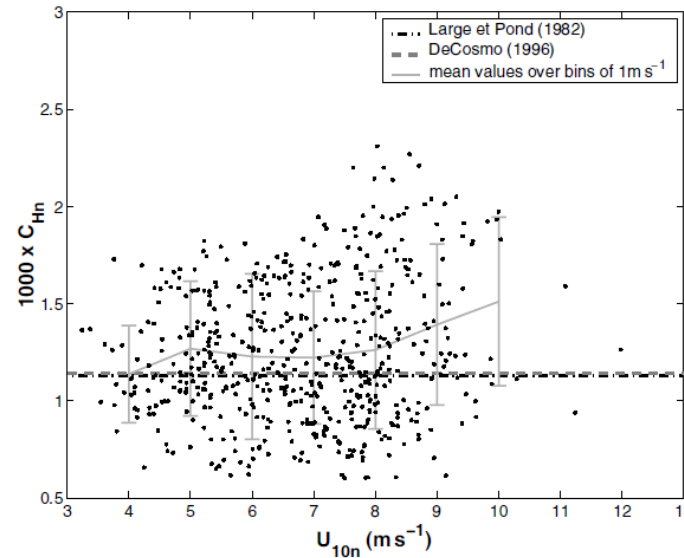
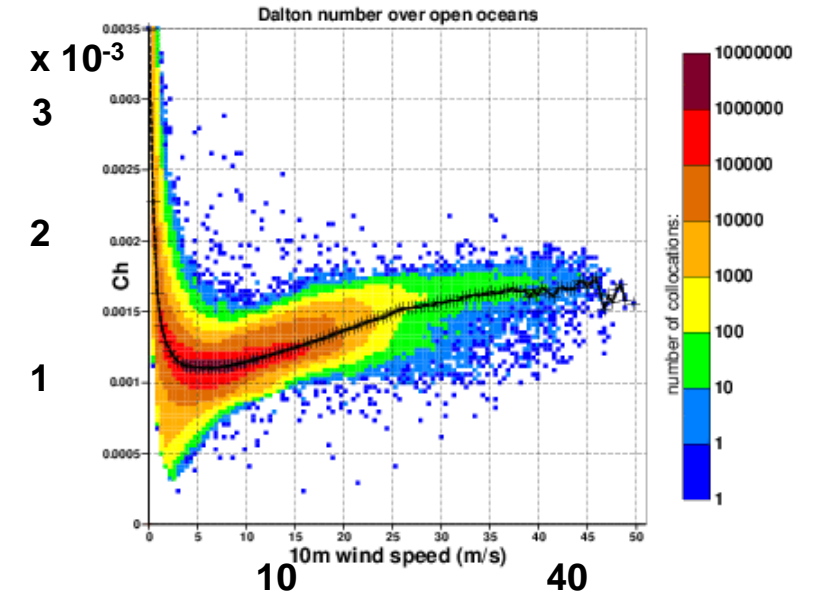


Figure 18. The exchange coefficient for temperature, C_{Hn} , as a function of the neutral wind speed at 10 m, U_{10n} . The dots correspond to 30-minute samples. The solid line with error-bars represents the values averaged over wind speed bins of 1 m s^{-1} . The parametrizations proposed by Large and Pond (1982) and DeCosmo *et al.* (1996) are also plotted.

Brut et al. 2005



Exchange coefficients
dependency on wind speed
Right: for heat (Ch)

The current model
is underestimating a bit
the heat transfer from the
surface .

- Following an extension of the wind wave generation theory, a sea state dependent parameterisation for the roughness length scales for heat and humidity has been tested.

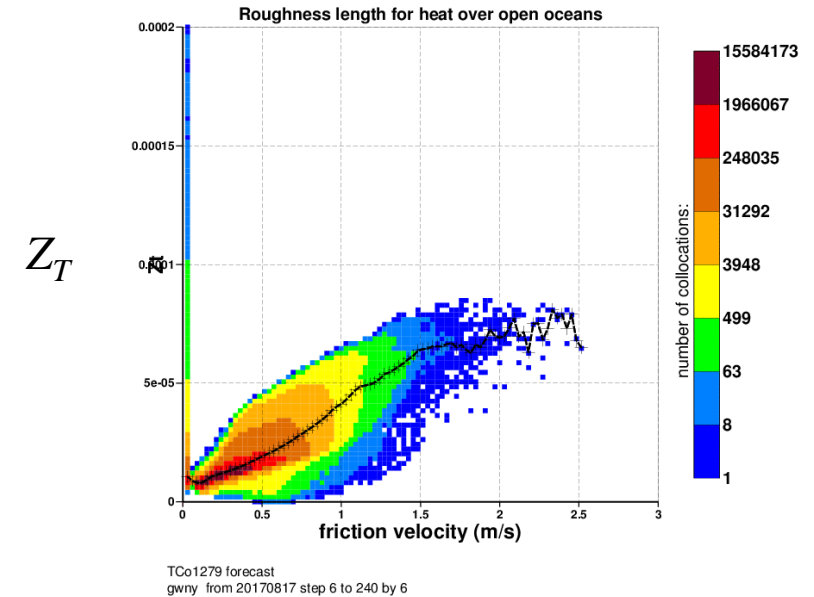
$$z_T = 10 \frac{\left(\frac{10 + x_-}{x_-} \right)^{(z_1 - x_-)}}{\left(\frac{10 + x_+}{x_+} \right)^{(z_1 - x_+)}}$$

$$x_{\pm} = (z_1 + \frac{1}{2} z_v) \mp \{ z_1^2 + (\frac{1}{2} z_v)^2 \}^{1/2}$$

$$z_1 = \frac{u_*^2}{g} (\alpha - \tilde{\alpha}) \quad z_v = \frac{\delta v}{\kappa u_*}$$

α Sea State dependent Charnock

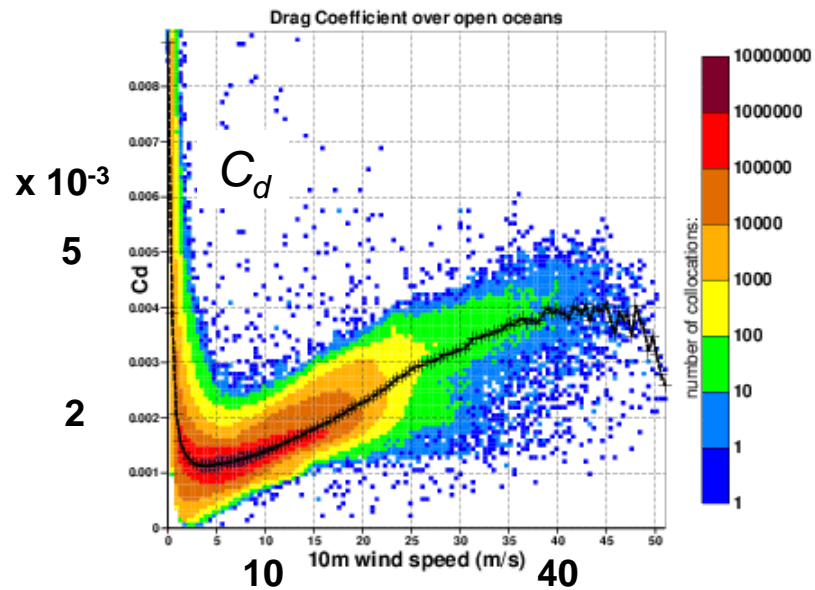
$\tilde{\alpha}$ minimum Charnock



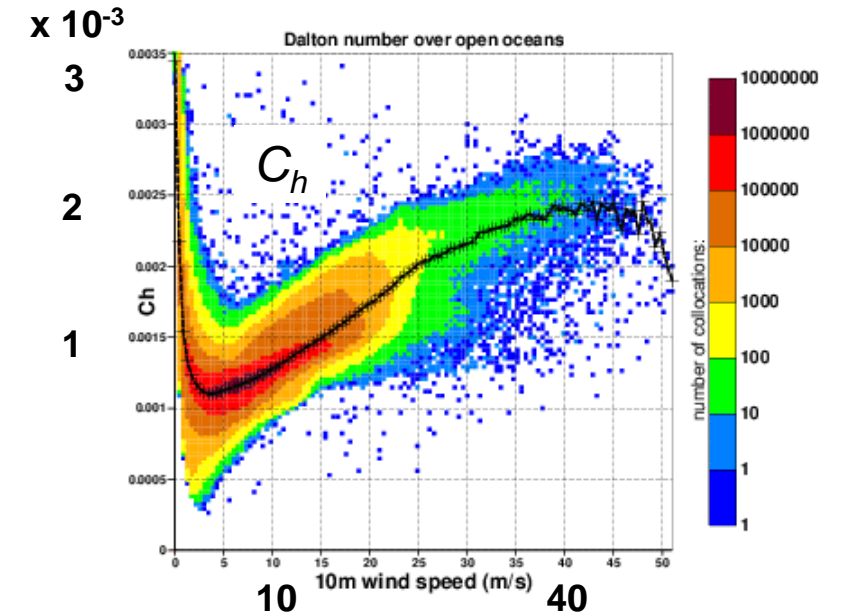
Peter A.E.M.Janssen and Jean-Raymond Bidlot, 2018: Progress in Operational Wave Forecasting, Procedia IUTAM Volume 26, 2018, Pages 14-29.

<https://www.sciencedirect.com/science/article/pii/S2210983818300038>

Impact of Coupling: revisit parameterisations



Heat exchange coefficients
dependency on wind speed

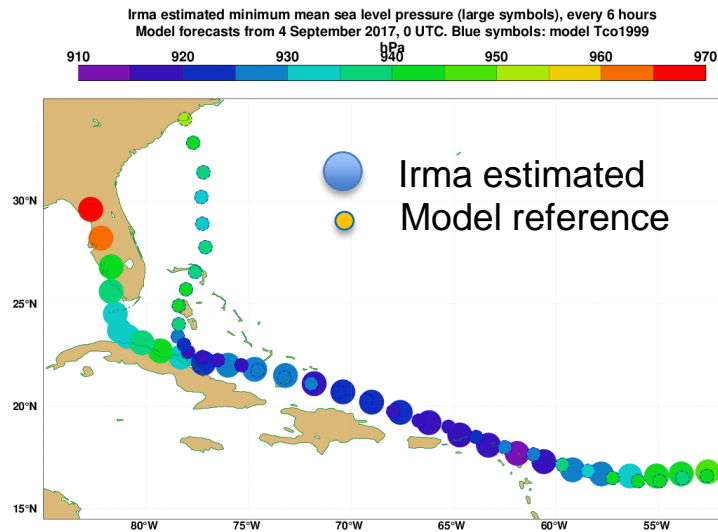


Similarly for moisture flux

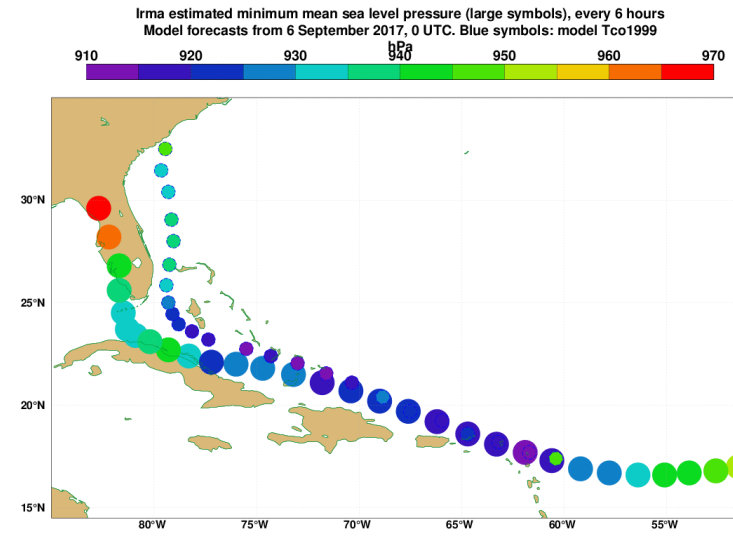
Sensitivity study: Hurricane Irma at Tco1999 resolution (5km)

Mean Sea Level
Pressure (hPa),
every 6 hours.

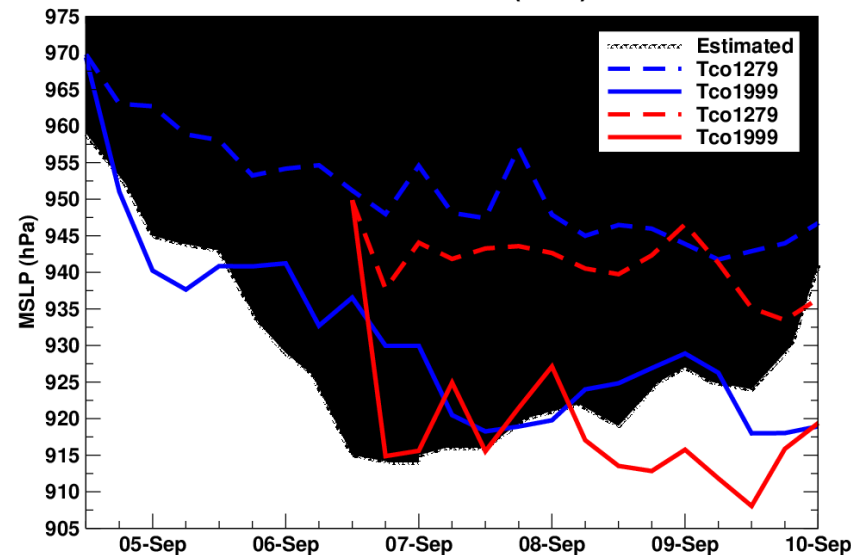
Forecast
from 4 September,
2017, 0 UTC.



Forecast
from 6 September,
2017, 0 UTC.



Mean Sea Level Pressure minimum (hPa)
Hurricane Irma (2017)



Blue: reference
Red: new parameterisation

NB:
ECMWF, 10 year
strategic goal is to run a
global ensemble at 5km
resolution

Conclusions

- The limitation of Charnock is based on wind speed only.
- Ongoing work with Peter Janssen to extend his theory on wind input to make this limitation dependent on waves parameters by introducing a model for the role of gravity-capillary waves on the overall surface levels.
- Implementing direct wave effect on the exchange coefficients for heat and moisture should be considered in future model cycle.

Thank you for your attention ...

Øyvind Breivik, Kristian Mogensen, Jean-Raymond Bidlot, Magdalena Alonso Balmaseda, and Peter A.E.M. Janssen, 2015: Surface Wave Effects in the NEMO Ocean Model: Forced and Coupled Experiments. JGR, doi: 10.1002/2014JC010565

Janssen, P.A.E.M., 1997: Effect of surface gravity waves on the heat flux. ECMWF Technical Memorandum 239.
<http://www.ecmwf.int/en/elibrary/technical-memoranda>

Peter A.E.M.Janssen and Jean-Raymond Bidlot, 2018: Progress in Operational Wave Forecasting, Procedia IUTAM Volume 26, 2018, Pages 14-29.
<https://www.sciencedirect.com/science/article/pii/S2210983818300038>

K.S. Mogensen, L. Magnusson and J-R. Bidlot., 2017: Tropical cyclone sensitivity to ocean coupling in the ECMWF coupled model J.Geophys. Res. Oceans, 122, 4392–4412, DOI: 10.1002/2017JC012753