

Thin Clouds Producing Graupel in the Arctic



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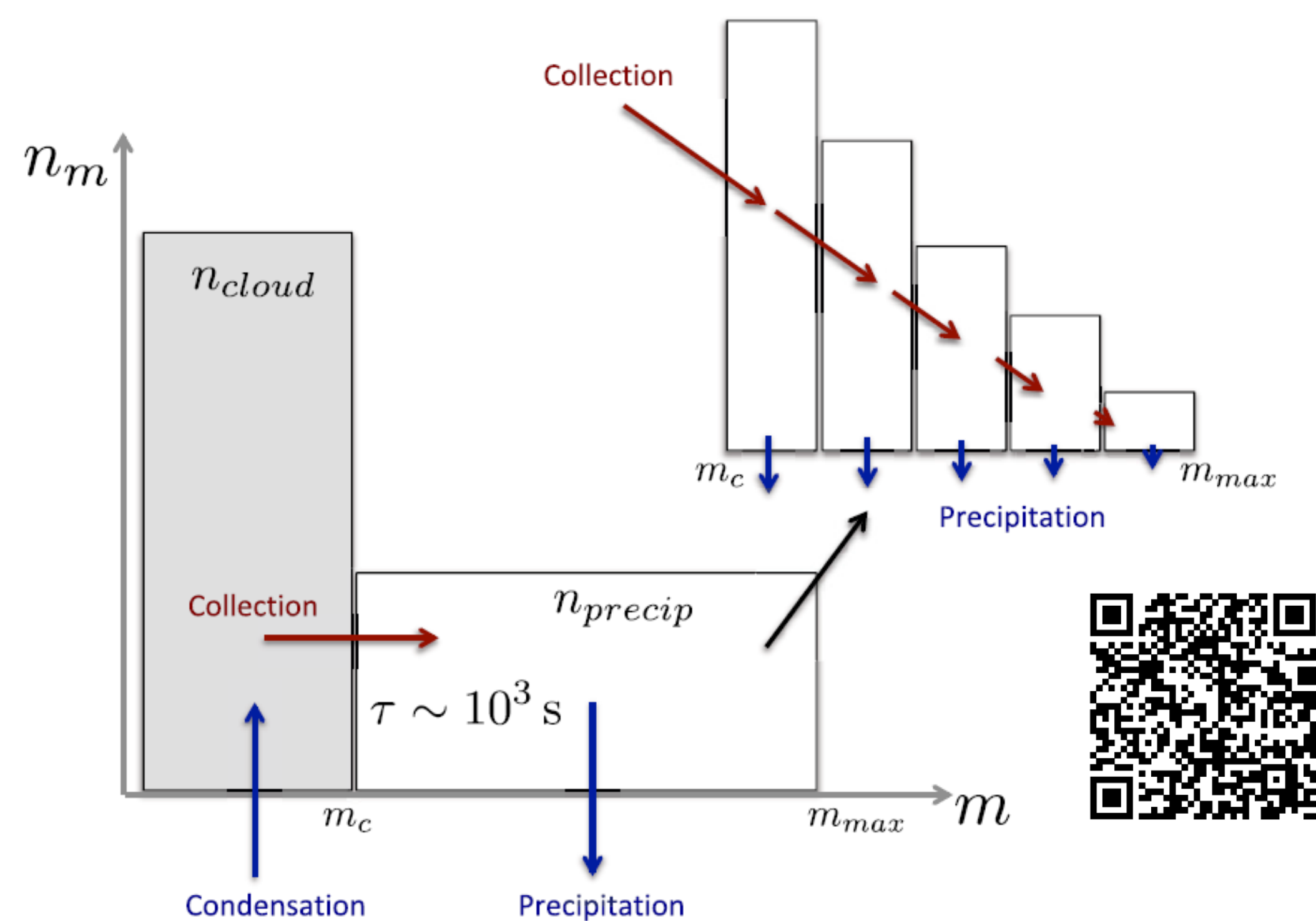
INTRODUCTION

- A common continuous collection approach for riming growth is to assume a relatively massive ice crystal with cross-section $\sigma \propto D^2$ falls at terminal velocity v_T through a homogeneous, suspended field of supercooled liquid cloud droplets with concentration ρ_c , collecting with efficiency E_c :

$$\frac{dm}{dt} = E_c \sigma v_T \rho_c \quad (1)$$
- In reality, these hydrometeors are exposed to turbulent motions that enhance the role of inertia (Naso et al., 2018) and lead to changes in localized convective precipitation intensities (Lee et al., 2014; Lee and Baik, 2016). A similar effect may be occurring in the Arctic.

UPDRAFTS DECREASE PSD SLOPE

- Analytical solutions for precipitation size distributions derived by assuming timescales of collection and precipitation are approximately the same – i.e., $\tau_{coll} \sim \tau_{precip} \sim 10^3$ s (Garrett, 2019)



- General solution for exponential tail of resulting gamma size distribution is

$$n_D = n_{D_0} \exp \left\{ -\lambda \left[1 - \frac{w}{aD^b(1-b)} \right] D \right\} = \exp \{ -\lambda' D \} \quad (2)$$

where the slope of the tail λ is modified based on the magnitude of the updraft speed w relative to the fall speed $v_p = aD^b$

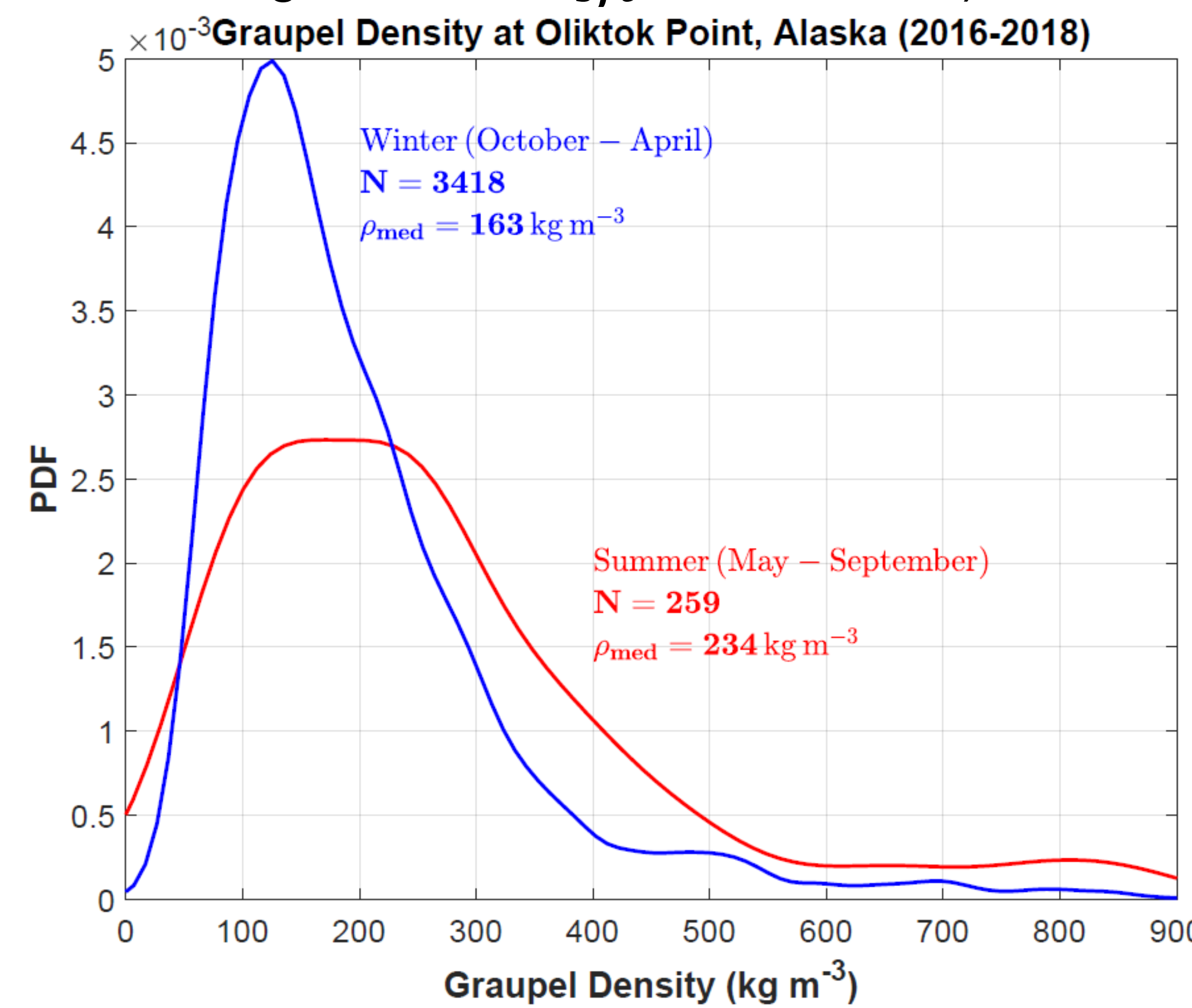
- If $w \ll v_p(1-b)$ ($\ll \sim 0.4 \text{ m s}^{-1}$ for lump graupel)

$$\exp \{ -\lambda D \} \quad (3)$$
- & slope is related to bulk density ρ_e & liquid water path LWP

$$\lambda = \frac{2\rho_e}{LWP} \quad (4)$$

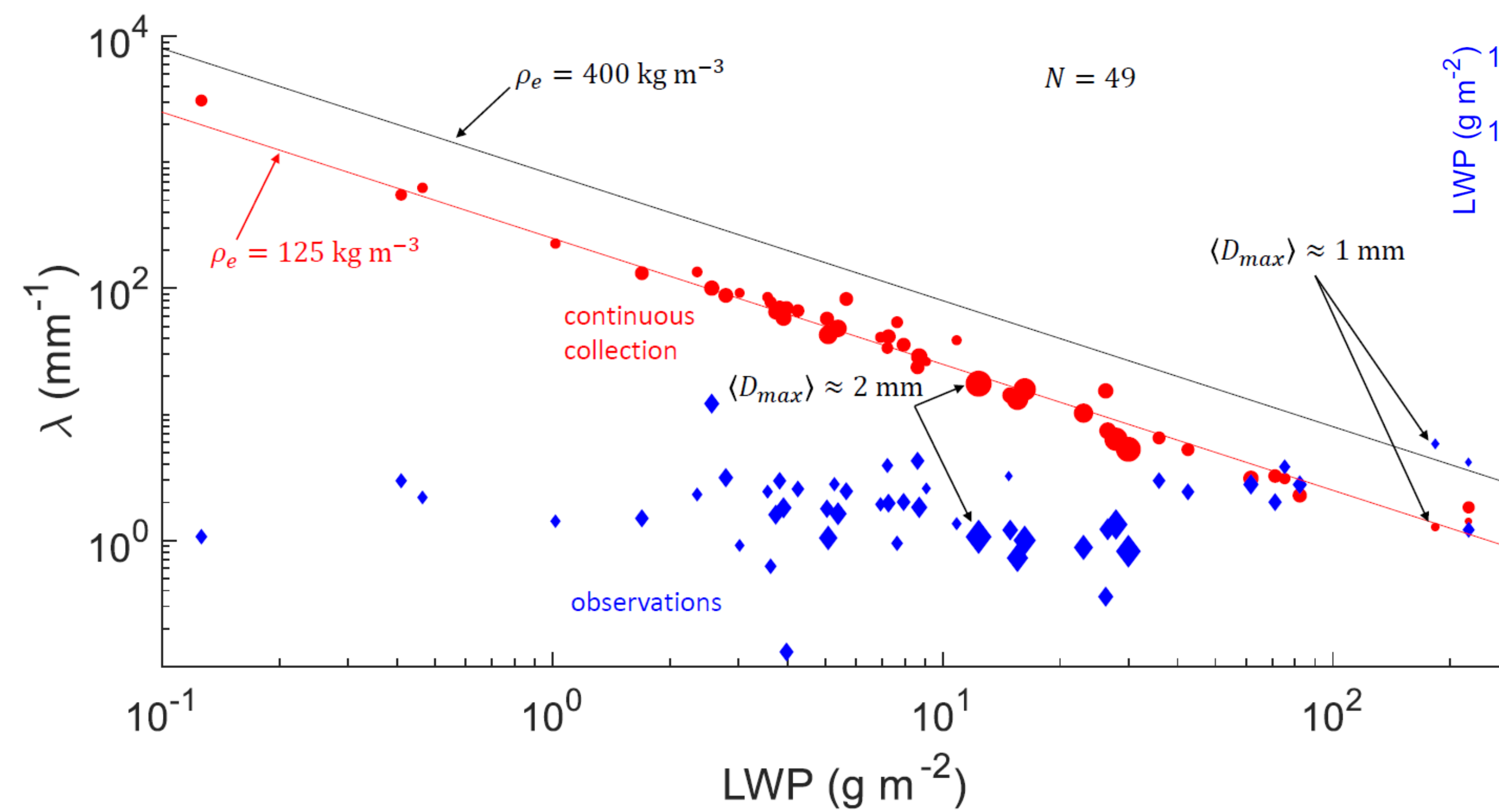
SUMMERTIME GRAUPEL MORE DENSE THAN WINTERTIME GRAUPEL

- Assumes spherical graupel & $v_p \approx v_T$
- Measurements made in still air (shielding + light winds, $U_{sfc} < 4.5 \text{ m s}^{-1}$)



CONTINUOUS COLLECTION DOESN'T ALWAYS EXPLAIN RIMING

- Observed λ' is the slope of the least-squares, best-fit line passing through the tail of a 10-minute size distribution
- Continuous collection λ is calculated from Eqn. (4)
- $\lambda' \ll \lambda$ for $LWP < 50 \text{ g m}^{-2}$; to satisfy Eqn. (2) with $b = 0.66$, $w \approx v_p/3$ is implied – or more generally $w/v_p \approx 1 - b$

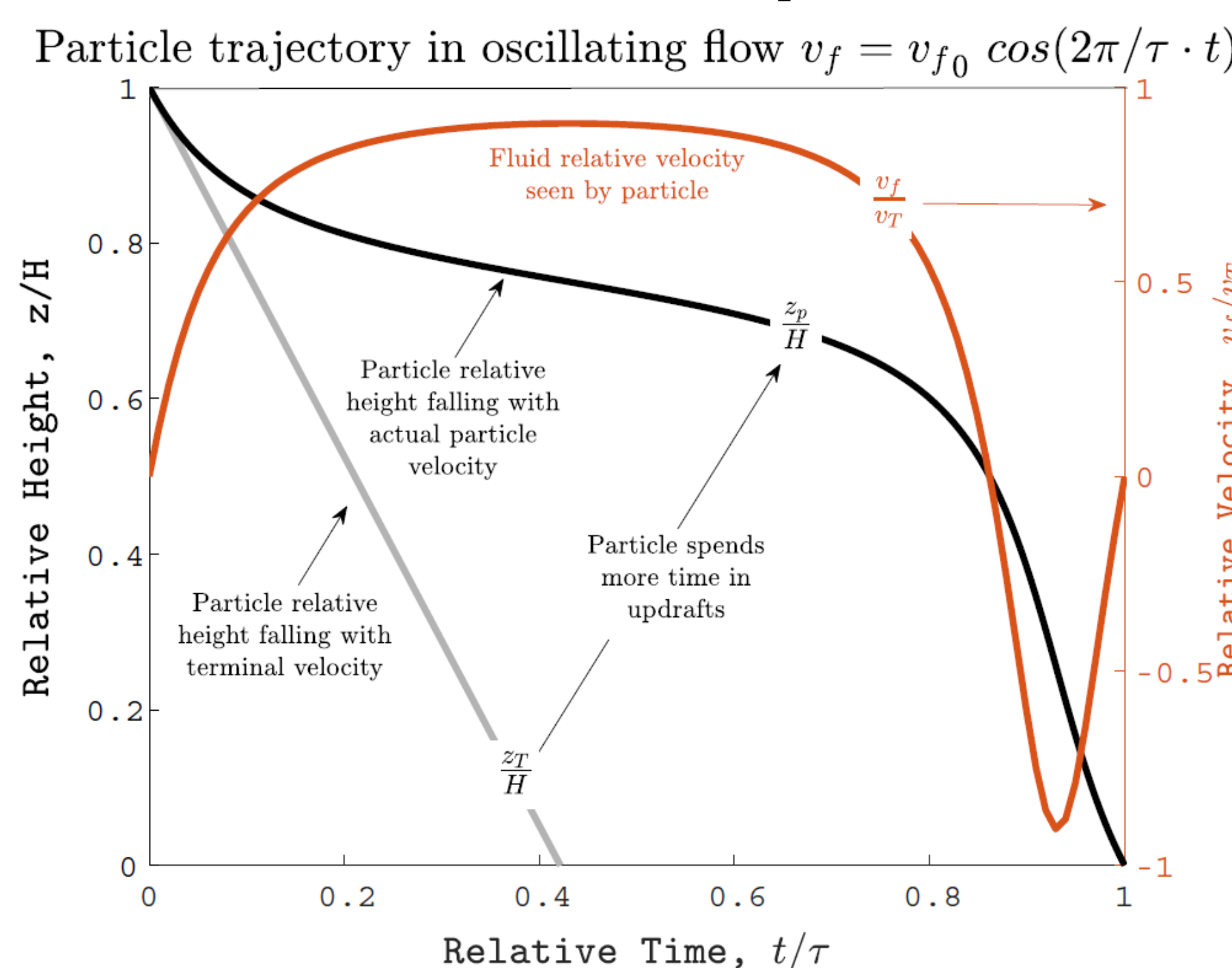


Turbulence and updrafts combining to enhance riming in thin Arctic clouds

OSCILLATING FLOW EXTENDS RIMING TIME

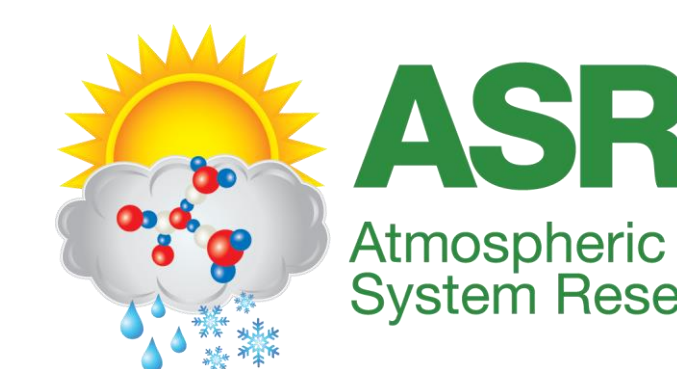
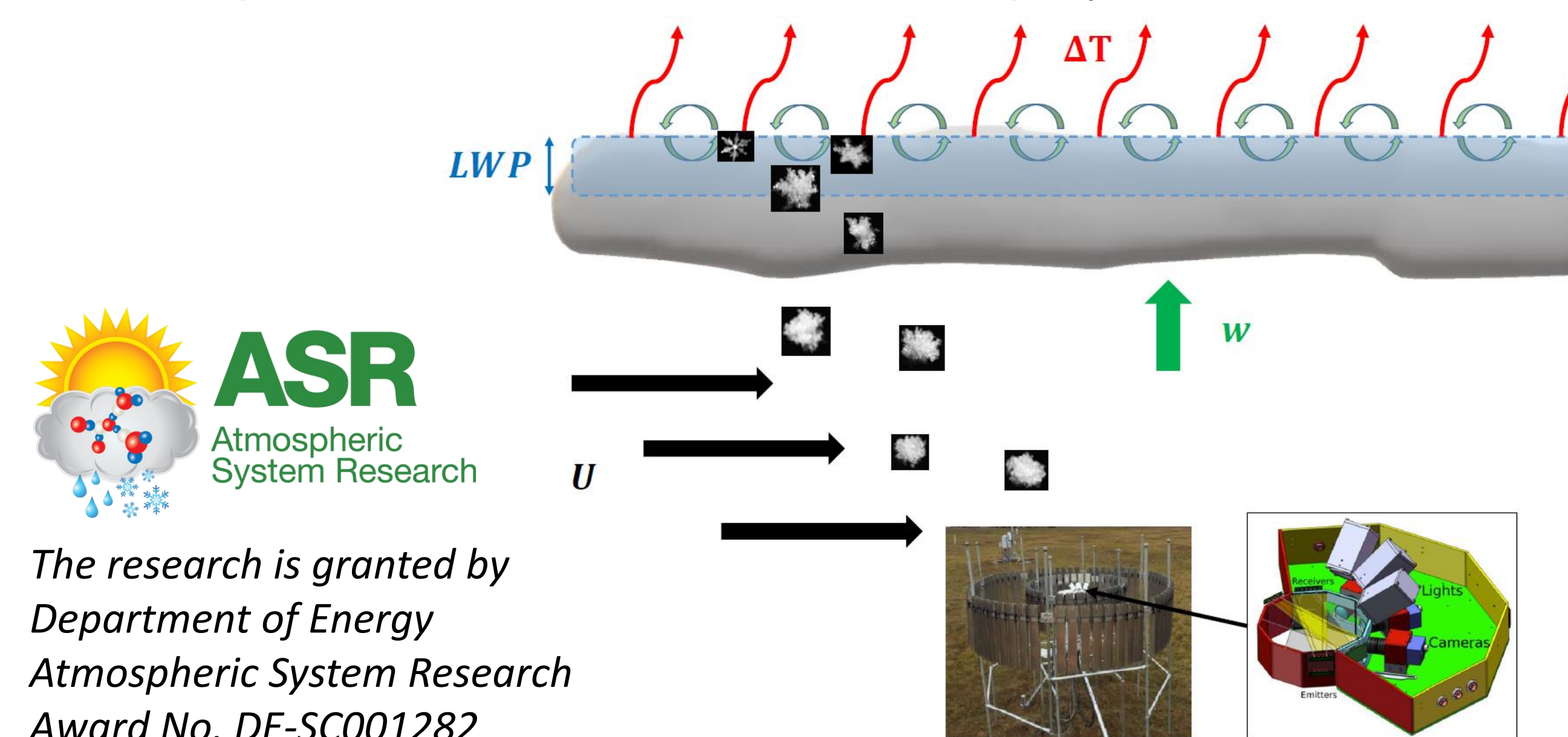
- Oscillating, turbulent fluid velocity (v'_f) is “frozen” with respect to particle ($v'_f \ll v_T$)
- Lagrangian frame of reference
- Relative particle velocity parameterized as

$$\frac{\langle v_p \rangle}{v_T} \approx 1 - \frac{\langle v_f'^2 \rangle}{v_T^2} \quad (5)$$



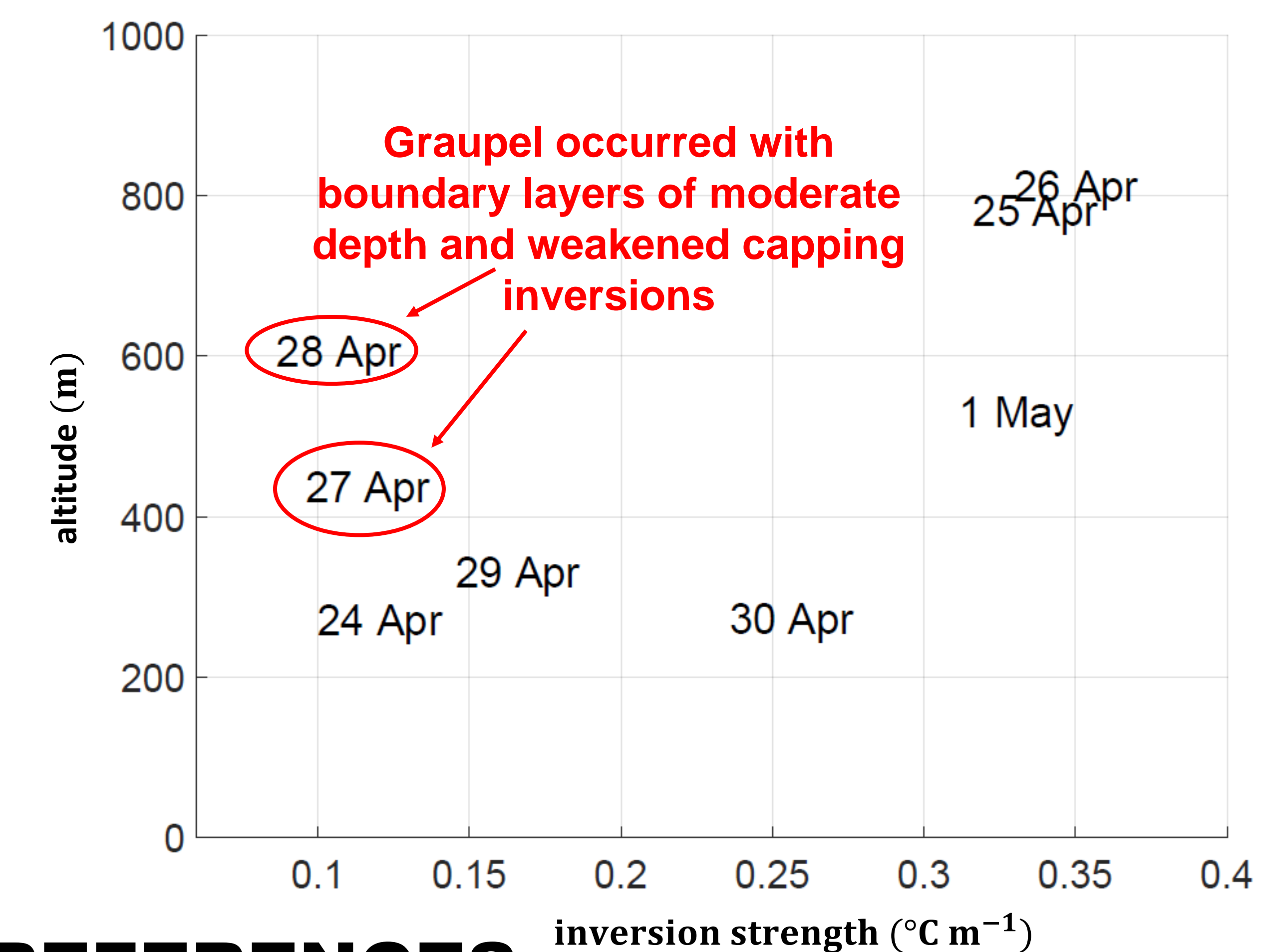
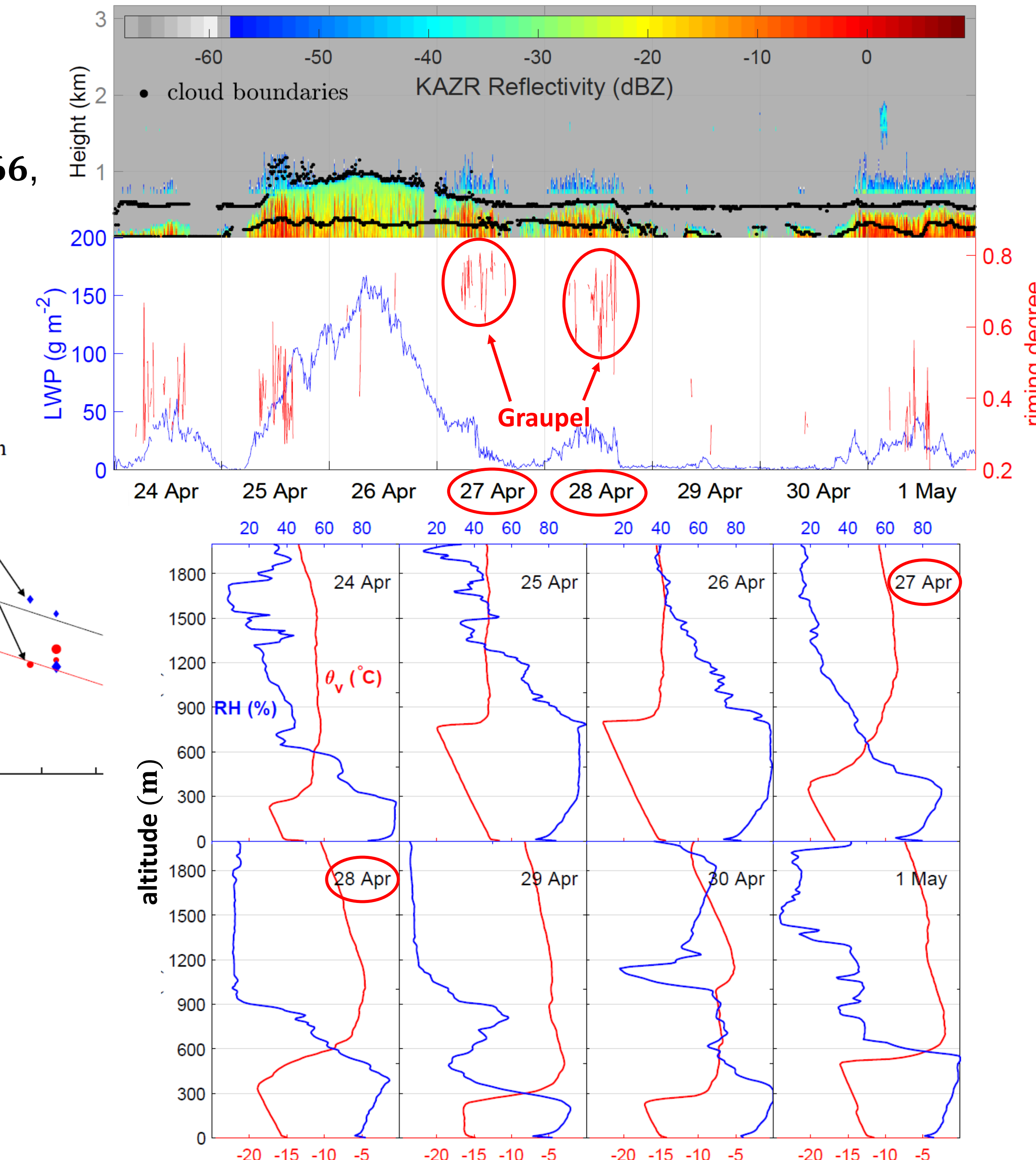
CONCLUSIONS

- In simulated turbulent (oscillating) flow, the particle spends more time in the updraft than in the downdraft, thereby increasing the time spent riming in the cloud.
- Continuous collection theory and observations only agree on λ for cases where $LWP > 50 \text{ g m}^{-2}$. Below that, updraft strength and turbulence may help to explain enhanced riming.
- Thin cloud case showed graupel occurring with weak inversions. The weakening of the boundary layer top suggests cloud top entrainment and turbulence played a role.



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THIN-CLOUD GRAUPEL WITH WEAK CAPPING INVERSION



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*Dr. Garrett is a co-owner of Particle Flux Analytics, Inc., the manufacturer of the Multi-Angle Snowflake Camera (MASC)