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# **Investigations of the Oligocene-Miocene opening of the** Ligurian Basin using refraction seismic data - LOBSTER

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Mountain Building Processes in 4D





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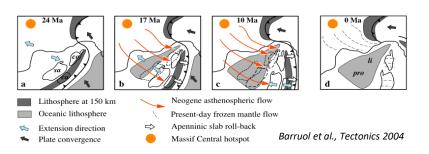
Introduction

## LOBSTER - Ligurian Ocean Bottom Seismology and Tectonics Research

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## Introduction

- The Ligurian Basin is a back-arc basin generated by SE trench retreat of the Apennines-Calabrian subduction zone.
  - ~30 Ma rifting was initiated associated with magmatism on land along the western margin.
  - ~21 Ma end of rifting and start of anticlockwise rotation of the Corsica-Sardinia block and oceanic spreading was <sup>43.00°</sup> proposed.
  - ~16-15 Ma end of opening associated with a second volcanic phase along the Corsican margin



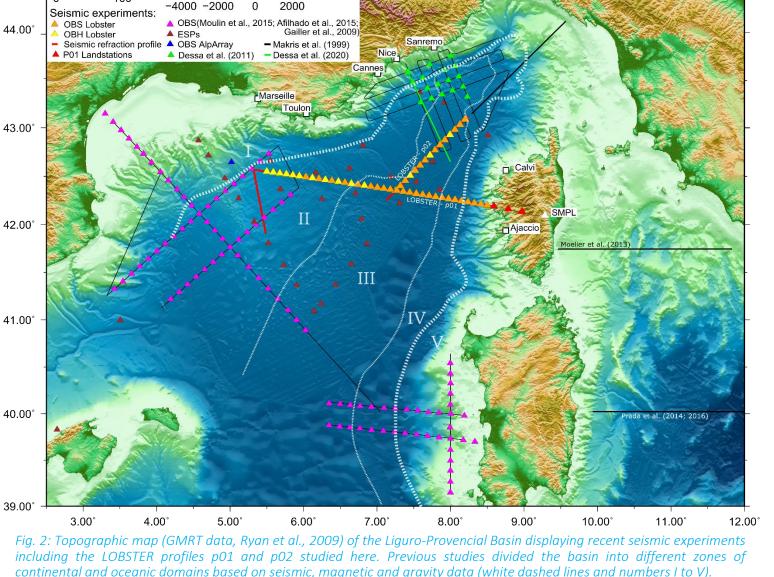
## *Fig. 1: Geodynamic evolution of the Liguro-Provencial Basin after Barruol et al. (2004)*

Key questions of our study:

What is the nature of the Ligurian Sea basement?

Where is the transition from continental to oceanic crust?

How is the geometry of the continent-ocean transition (COT) zone?





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## LOBSTER - Ligurian Ocean Bottom Seismology and Tectonics Research



**Results and interpretation of LOBSTER profile p02** 

- Stacked picks along p02 show a nearly 1D structure (Fig. 4a)
- Clear arrivals from sediments (Psed < 4.6 km/s) and a small velocity gradient</li>
- Abrupt change to mantle phases (Pn ~ 8 km/s) (Fig. 4b)

Fig. 3: In total 15 OBS/H were deployed along profile p02.

Results from travel time tomography (Fig. 5):

Uppermost sedimentary layer: 2.2 – 3.5 km/s with a high velocity gradient (Plio-Quaternary)

43.00

42.50

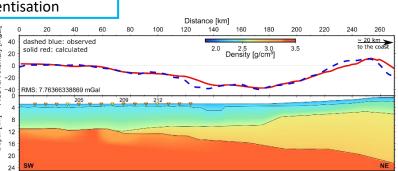
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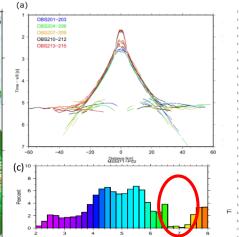
- Deeper layers: 3.5 ~5.7 km/s (Messinian salt and Pre-Messinian and Tortonian sediments)
- Acoustic Basement: at ~9 km (CB-MCS) or ~10 11.5 km depth (CB)
- Mantle: at ~11-13 km depth and velocities >7.3 km/s
- Results from the Free-Air gravity modelling support the findings (Fig. 6).

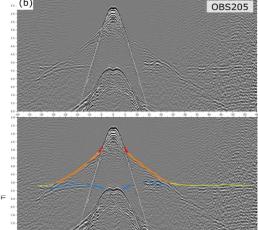
 $\rightarrow$  Continental crust along the entire profile p02

 $\rightarrow$  Low degree of mantle serpentisation

Fig. 6: Density model along the LOBSTER profile pO2 and extended northwards along the MAKRIS profile towards the NE Ligurian coast (Dannowski et al., 2020). The upper panel shows the fit between the model response and the observed Free-Air gravity data from (Smith and Sandwell (2014).

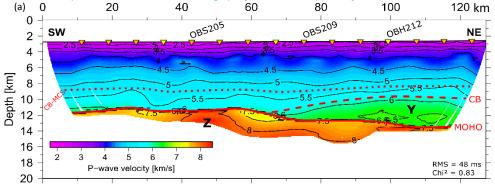


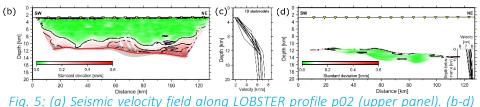




Distance [km]

Fig. 4: (a and b) Data and travel time picks of LOBSTER profile p02 (modified, Dannowski et al., 2020). (c) The shows the velocity distribution of the final velocity model and indicates a gap of velocities typical for oceanic crust.





*Fig. 5: (a) Seismic velocity field along LOBSTER profile p02 (upper panel). (b-d) Statistics of the modelling procedure (modified, Dannowski et al., 2020).* 

Data, results and interpretation

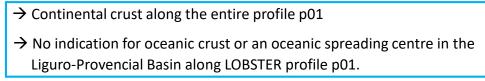


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### Preliminary results and interpretation of LOBSTER profile p01

First results along p01 are similar to profile p02 in the centre of the basin:

- Uppermost sedimentary layer: ~2.2-3.5 km/s with a high velocity gradient (Plio-Quaternary)
- Deeper layers: ~3.5-5.6 km/s (Messinian salt and Pre-Messinian and Tortonian sediments); deeper sediment layers pinch out towards Corsica.
- Crystalline Basement (CB): at ~9-10 km in the central basin shallowing towards the Corsica coast with seismic velocities between ~5.6 km/s to 6.8 km/s and thickening in the Necking Zone (~30 km distance) from 6 km to ~24 km. The lower crust pinches out in the Central Basin.
- Mantle: at ~12 km depth and velocities >7.5 km/s in the central basin. It slightly shallows in the distal margin to ~11 km before it extremely deepens over a short distance to a depth of ~24 km.



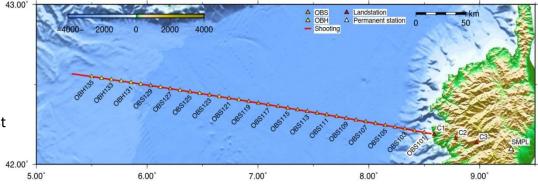
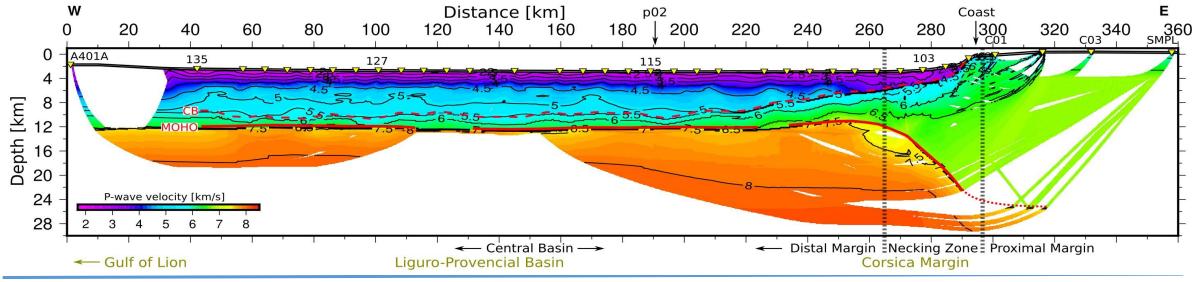


Fig. 7: In total 35 OBS/H and 3 stations on land (C01-C03) were deployed along profile p01. The red line illustrates the position of the airgun shots that were given every ~123 m.

Fig. 8: Preliminary P-wave velocity model. The red dashed line marks the crystalline basement (CB), while the solid red line marks the seismic crust-mantle boundary (MOHO). Between profile KM 110 and KM 160 we observe a poor signal penetration of the deeper sediment phases and below.

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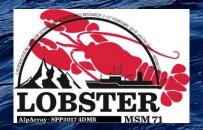


Preliminary results and interpretation

## Thank you!

### References

Afilhado et al., 2015, https://doi.org/10.2113/gssgfbull.186.4-5.331 Barruol et al., 2004, https://doi.org/10.1016/j.tecto.2004.08.002 Dannowski et al., 2020 (modified accepted), https://doi.org/10.5194/se-2019-187 Gailler et al., 2009, https://doi.org/10.1016/j.epsl.2009.07.001 Makris et al., 1999, https://doi.org/10.1016/S0040-1951(98)00225-X Moulin et al., 2015, https://doi.org/10.2113/gssgfbull.186.4-5.309 Ryan et al., 2009, https://doi.org/10.1029/2008GC002332















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