© by authors. All rights reserved

Revisiting Lake Garba Guracha, high altitude lake in the Bale Mountains, Ethiopia: reconstructing Late Glacial – Holocene lake level history using δ2H/δ18O biomarker

analyses



L. Bittner^{1,2}, M. Bliedtner³, D. Grady⁴, G. Gil-Romera^{5,6}, C. Martin-Jones^{7,8}, B. Lemma², H.F. Lamb⁴, C. De Jonge⁹, H. Meyer¹⁰, B. Glaser², M. Zech¹

1) Heisenberg Chair of Physical Geography with focus on paleoenvironmental research, Institute of Geography, Technical University of Dresden, Germany 2) Institute of Agronomy and Nutritional Sciences, Soil Biogeochemistry, Martin-Luther-Universität Halle-Wittenberg, Halle (Saale), Germany 3) Department of Physical Geography, Institute of Geography, Friedrich-Schiller-University Jena, Jena, Germany 4) Department of Geography and Earth Sciences, Aberystwyth University, Aberystwyth, UK 5) Department of Geo-environmental Processes and Global Change, Pyrenean Institute of Ecology, CSIC, Zaragoza, Spain 6) Department of Ecology, Philipps-Marburg University, Marburg, Germany 7) Department of Geography, University of Cambridge, Cambridge CB2 3EN, UK 8) Limnology Unit, Department of Biology, Ghent University, B-9000 Gent, Belgium 9) Geological Institute, Department of Earth Sciences, ETH Swiss Federal Institute of Technology, 8092 Zurich, Switzerland 10) Alfred Wegener Institute Helmholtz Centre for Polar and Marine Research, Telegrafenberg A45, Potsdam, Germany



MARTIN-LUTHER

UNIVERSITÄT



Introduction

The climate of East Africa is driven by the position of the Intertropical Convergence Zone (ITCZ). The intensity and position of the ITCZ related tropical rain belt changes depending on the interhemispheric temperature gradient (Broccoli et al., 2006). During the early Holocene boreal summer insolation maximum, the mean position of the tropical rain belt shifted north leading to increased precipitation across northern Africa (Gasse, 2000). Additionally, an increased landocean temperature gradient caused a strengthening of the West African Monsoon (WAM) and Indian Summer Monsoon (ISM) generating a water level rise in African lakes (Junginger et al., 2014; Lezine et al., 2014). This shift to a more pluvial early-mid Holocene (12-5 ka), termed African Humid Period (AHP) (DeMenocal et al., 2000), was particularly intense in North Africa and extended south until 10° S in East Africa (Gasse, 2000). While the general mechanisms for the orbitallyforced AHP are well understood the spatial and temporal patterns are highly debated. As part of the Research Unit FOR 2358 'The Mountain Exile Hypothesis', we address the question of 'How humans' benefited from and re-shaped African high-altitude ecosystems during Quaternary climate changes . Therefore, we investigated the high-altitude, small catchment, cirque lake Garba Guracha at 3950 m asl. - an ideal sedimentary archive for reconstructing afro-alpine paleoclimate/-environment. We show geochemical, biomarker and diatom isotope $(\delta^{18}O_{fucose}$ and $\delta^{18}O_{diatom})$ results for the Late Glacial and Holocene.





Specifically, we aim to:

- 1. develop a robust chronology,
- 2. reconstruct the sedimentary history and
- 3. the climate and lake level history since the Last Glacial.

Methods

We analyzed the sediments of Garba Guracha (15.5 m core)

Fig. 2: Summer insolation (Laskar et al. 2011), TOC, TOC/N, Botryococcus braunii (Umer et al. 2007), δ^{13} C, (fuc+xyl)/ara, fuc/(ara+xyl), P_{aq}, titan element and calculated sedimentation rate.



Fig. 4: Ternary diagram illustrating the relative contributions of ara, fuc and xyl in the GG sedimentary record. For comparision, plant, organic layer and topsoil samples from the GG catchment are included (Mekonnen et al., 2019).

Environmental implications

Phase 1 (Fig. 2): High sedimentation rates, high minerogenic input and low TOC values point to fast filling lake bed with a low vegetated catchment. Low TOC/N values and relatively positive δ^{13} C values point to aquatic algae as dominant organic matter source.

Phase 2: The time period of the northern hemisphere Younger Dryas (YD) is marked by a decrease in sedimentation rate by 80% indicating a dry or/and cold phase.

- using the ITRAX[™] core scanner.
- Total carbon (TC), total nitrogen (N) and stable carbon and nitrogen isotopic composition (δ^{13} C and δ^{15} N, respectively) were determined for 110 mixed sediment samples covering roughly 10 cm intervals.
 - TC is equal TOC





Phase 3: With increasing insolation TOC and TOC/N reach the highest values between the Holocene onset and 4.3 ka (ca. African humid period - AHP) pointing to a phase of favorable growth conditions. Most negative δ^{18} O values support high rainfall amounts (amount effect), a different moisture source (source effect) or/and a high P/E value and an overflowing lake. Northhemisphere cold spells (8.2 and 6.5 ka event) are visible in different proxies.

Phase 4: A rapid change in TOC, TOC/N concur with the 4.2 ka event known as a shift to dryer conditions.

Conclusions

- \succ Dating of compound class *n*-alkanes is a valuable tool for lake sediment dating in small catchment areas
- \succ The minimum age of deglaciation is ~16000 cal. yrs BP)
- \succ $\delta^{18}O_{\text{fucose}}$ can be used as proxy for $\delta^{18}O_{\text{lake water}}$
 - $\delta^{18}O_{diatom}$ as aquatic signal supports this finding.
 - Most negative δ^{18} O values agree with other records from the region (Fig. 3) pointing to increased precipitation or/and a different moisture source.
 - Moreover, the range of $\delta^{18}O_{fucose}$ can not be explained by source and amount effect alone. Evaporative enrichment must be considered interpreting the $\delta^{18}O_{sugar}$ record.
- > Most negative δ^{18} O values between 10 and 7 ka concurring with the AHP indicating a high P/E ratio and an overflowing

Fig. 1: Age-depth model for the Garba Guracha sedimentary record. The model was created using Bacon (Blaauw and Christen 2011). The main panel displays the depth (x-axis) and age (y-axis) scale. The probability distribution of the calibrated ages is represented as violin plots (modeled ages (blue), non-modeled ages (pink) and tephra ages (68 % probability) of Lake Tilo tephra TT1 and Lake Chamo tephra CHT2 (red). In the lower-left corner, the sedimentation rate with a mean sedimentation rate (dotted red line) is displayed.

Fig. 3: Comparison of lake level reconstructions in Eastern Africa and δ^{18} O records for the past 12,000 years (Dongge caves - Dykoski et al., 2005; Qunf cave - Fleitmann et al., 2003; Garba Guracha; Lake Abhè -Gasse, 2000; Ziway-Shala - Gillespie et al., 1983; Chew Bahir - Foerster et al., 2012; Turkana - Garcin et al., 2012 (filled curve), Johnson et al., 1991 (dotted curve), Brown and Fuller, 2008 (dashed curve); Paleolake Suguta - Junginger et al., 2014; and insolation variations (Laskar et al., 2004).

lake.

> Northern hemisphere events (8.2, 6.5, and 4.2 ka) are visible in the Garba Guracha record

References

Blaauw M, Christen JA (2011) Flexible paleoclimate age-depth models using an autoregressive gamma process. Bayesian Anal 6:457–474 Broccoli AJ, Dahl KA, Stouffer RJ (2006) Response of the ITCZ to Northern Hemisphere cooling. Geophys Res Lett 33:

Brown FH, Fuller CR (2008) Stratigraphy and tephra of the Kibish Formation, southwestern Ethiopia. J Hum Evol 55:366–403

deMenocal P, Ortiz J, Guilderson T, Adkins J, Sarnthein M, Baker L, Yarusinsky M (2000) Abrupt onset and termination of the African Humid Period: Quat Sci Rev 19:347-361 Dykoski CA, Edwards RL, Cheng H, Yuan D, Cai Y, Zhang M, Lin Y, Qing J, An Z, Revenaugh J (2005) A high-resolution, absolute-dated Holocene and deglacial Asian monsoon record from Dongge Cave, China. Earth Planet Sci Lett 233:71-86

Fleitmann D, Burns SJ, Neff U, Mangini A, Matter A (2003) Changing moisture sources over the last 330,000 years in Northern Oman from fluid-inclusion evidence in speleothems. Quat Res 60:223–232

Foerster V, Junginger A, Langkamp O, Gebru T, Asrat A, Umer M, Lamb HF, Wennrich V, Rethemeyer J, Nowaczyk N, Trauth MH, Schaebitz F (2012) Climatic change recorded in the sediments of the Chew Bahir basin, southern Ethiopia, during the last 45,000 years. Quat Int 274:25-37

Gasse F (2000) Hydrological changes in the African tropics since the Last Glacial Maximum. Quat Sci Rev 19:189-211

Gillespie R, Street-Perrott FA, Switsur R (1983) Post-glacial arid episodes in Ethiopia have implications for climate prediction. Nature 306:680-683

Johnson TC, Halfman JD, Showers WJ (1991) Paleoclimate of the past 4000 years at Lake Turkana, Kenya, based on the isotopic composition of authigenic calcite. Palaeogeogr Palaeoclimatol Palaeoecol 85:189–198

Junginger A, Roller S, Olaka LA, Trauth MH (2014) The effects of solar irradiation changes on the migration of the Congo Air Boundary and water levels of paleo-Lake Suguta, Northern Kenya Rift, during the African Humid Period (15 – 5 ka BP) CA. Palaeogeogr Palaeoclimatol Palaeoecol 396:1–16

Laskar J, Fienga A, Gastineau M, Manche H (2011) La2010: a new orbital solution for the long-term motion of the Earth *. A&A 532:

Laskar J, Robutel P, Joutel F, Gastineau M, Correia ACM, Levrard B (2004) A long-term numerical solution for the insolation quantities of the Earth . A&A 428:261–285

Lezine A-M, Bassinot F, Peterschmitt JYJ-Y, Lézine AM, Bassinot F, Peterschmitt JYJ-Y (2014) Orbitally-induced changes of the Atlantic and Indian monsoons over the past 20,000 years: New insights based on the comparison of continental and marine records. Bull la Soc Geol Fr 185:3–12

Mekonnen B, Zech W, Glaser B, Lemma B, Bromm T, Nemomissa S, Bekele T, Zech M (2019) Chemotaxonomic patterns of vegetation and soils along altitudinal transects of the Bale Mountains, Ethiopia, and implications for paleovegetation reconstructions – Part 1: stable isotopes and sugar biomarkers. E&G Quat Sci J 68:177–188

Umer M, Lamb HF, Bonnefille R, Lézine AM, Tiercelin JJ, Gibert E, Cazet JP, Watrin J (2007) Late Pleistocene and Holocene vegetation history of the Bale Mountains, Ethiopia. Quat Sci Rev 26:2229-2246

Zech M, Tuthorn M, Zech R, Schlütz F, Zech W, Glaser B (2014) A 16-ka ??18O record of lacustrine sugar biomarkers from the High Himalaya reflects Indian Summer Monsoon variability. J Paleolimnol 51:241–251