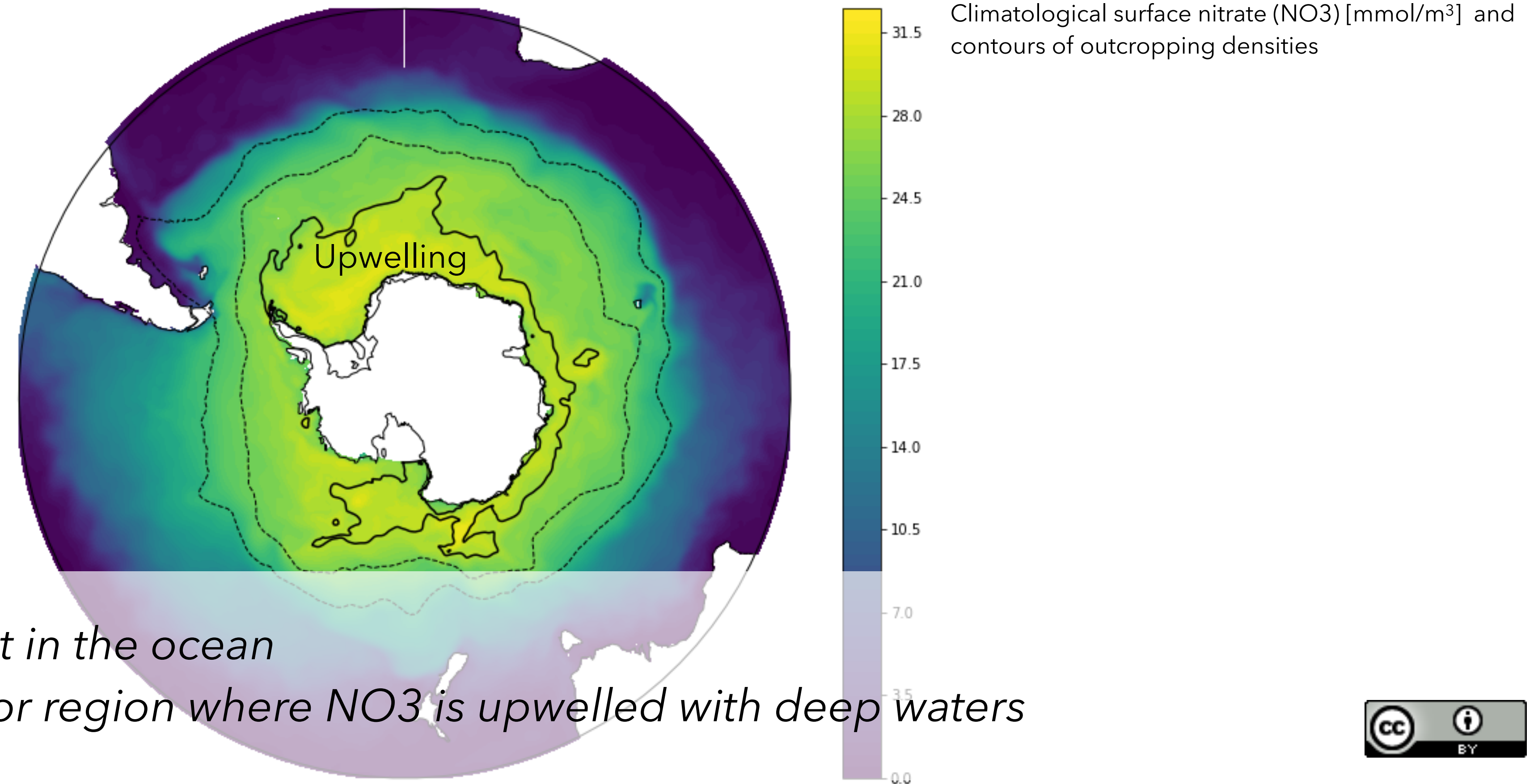
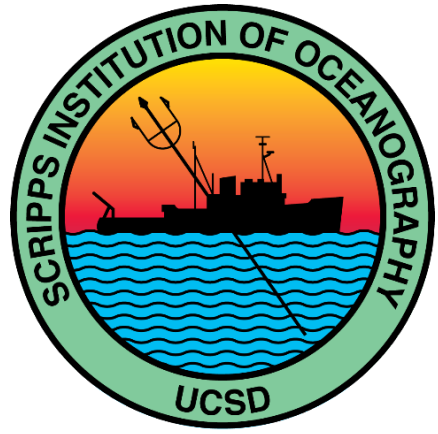


Reshuffling of Nitrate (NO_3) in the Southern Ocean

Ivy Frenger, Ivana Cerovecki, Matt Mazloff

With input from Wolfgang Koeve, Andreas Oschlies, Ariane Verdy

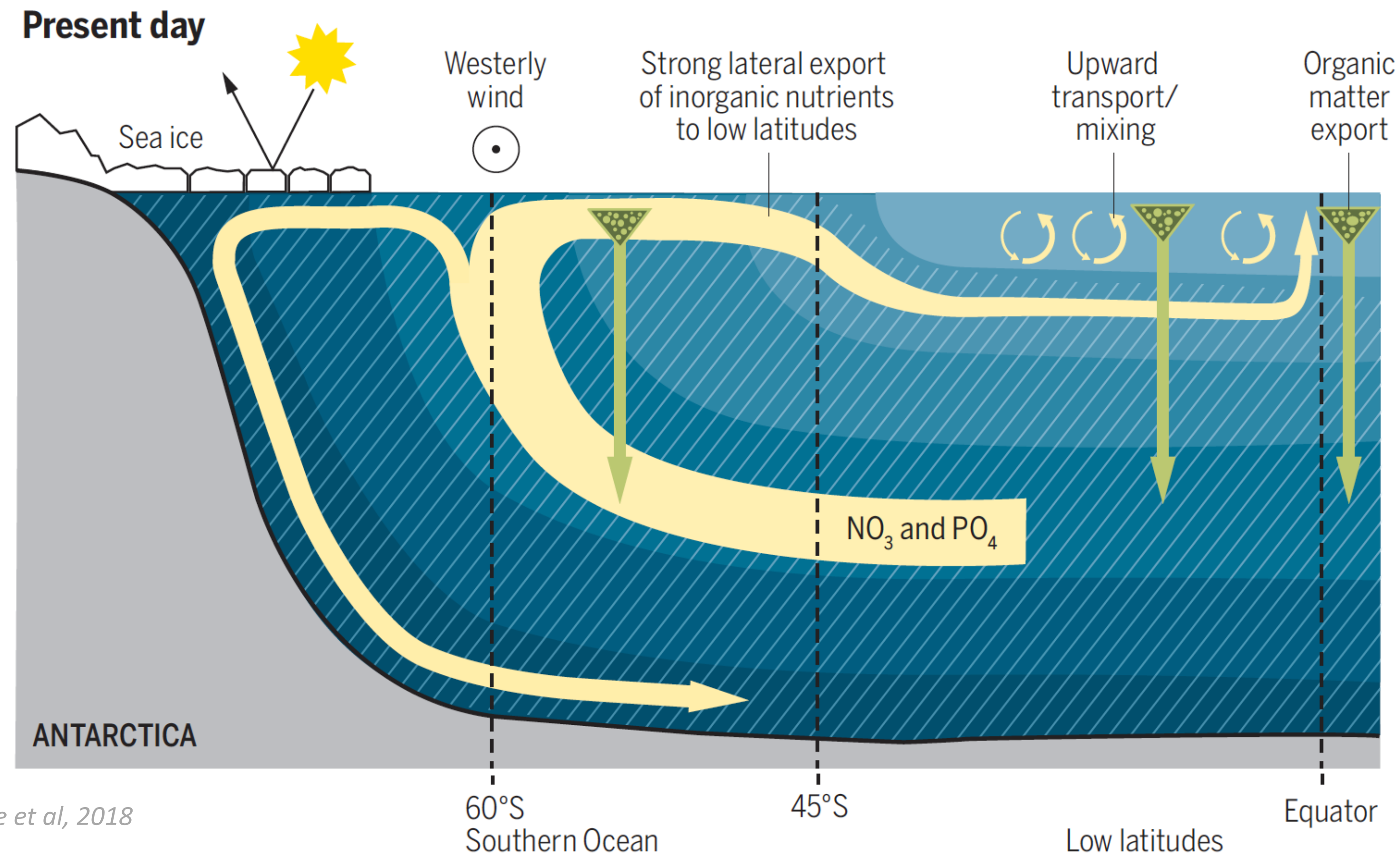
Sharing Geoscience Online, May 2020



- NO_3 is a major limiting nutrient in the ocean
- The Southern Ocean is the major region where NO_3 is upwelled with deep waters

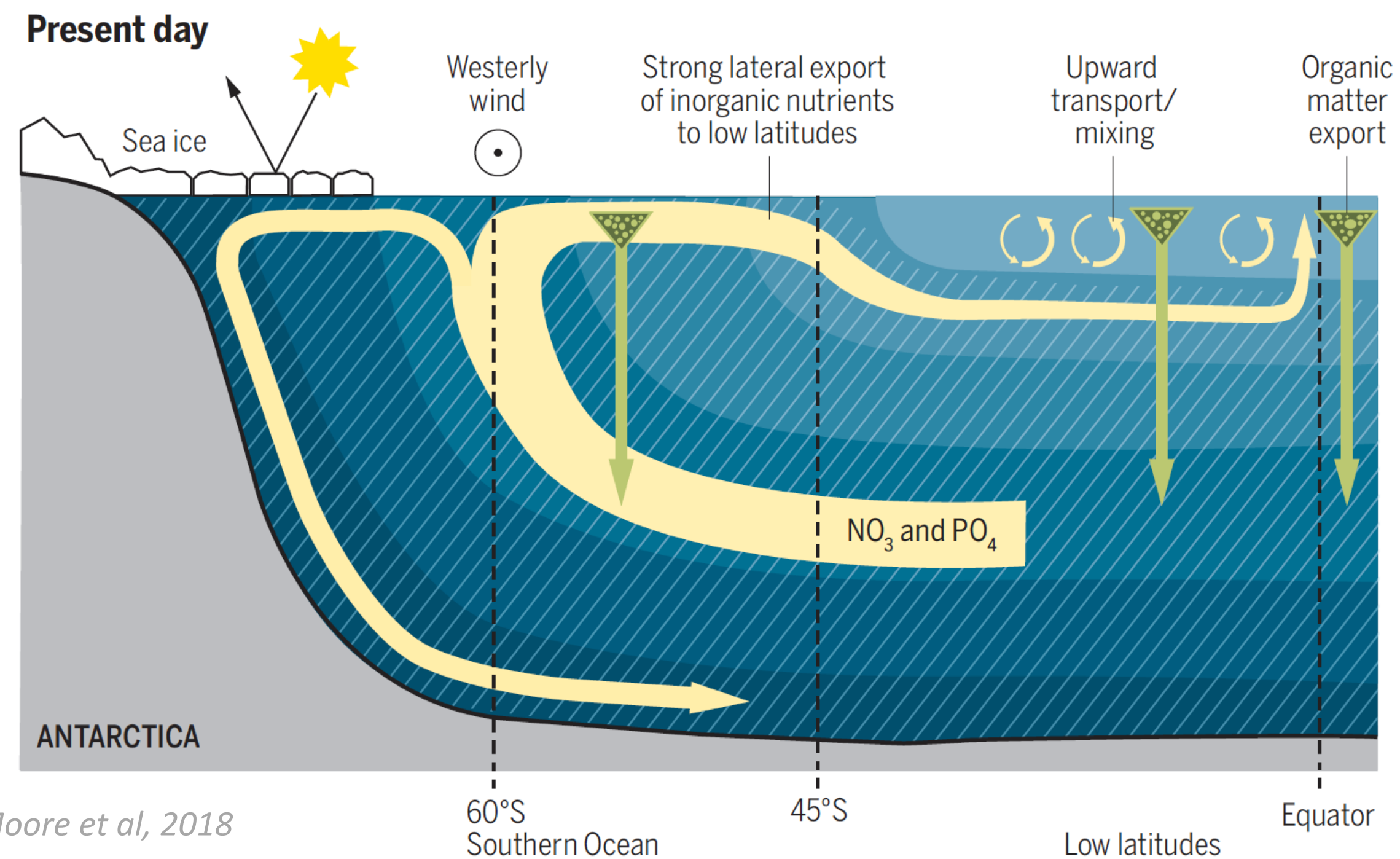
... MOTIVATION: WHY SHOULD I CARE ABOUT HOW NO₃ IS RESHUFFLED IN THE SOUTHERN OCEAN?

- ▶ E.g., Sarmiento et al, 2004, Moore et al, 2018:
Southern Ocean fuels lower latitude primary productivity -> **fisheries**
- ▶ E.g., Gruber & Galloway, 2008:
The nitrogen cycle is coupled with the carbon-cycle -> **climate**



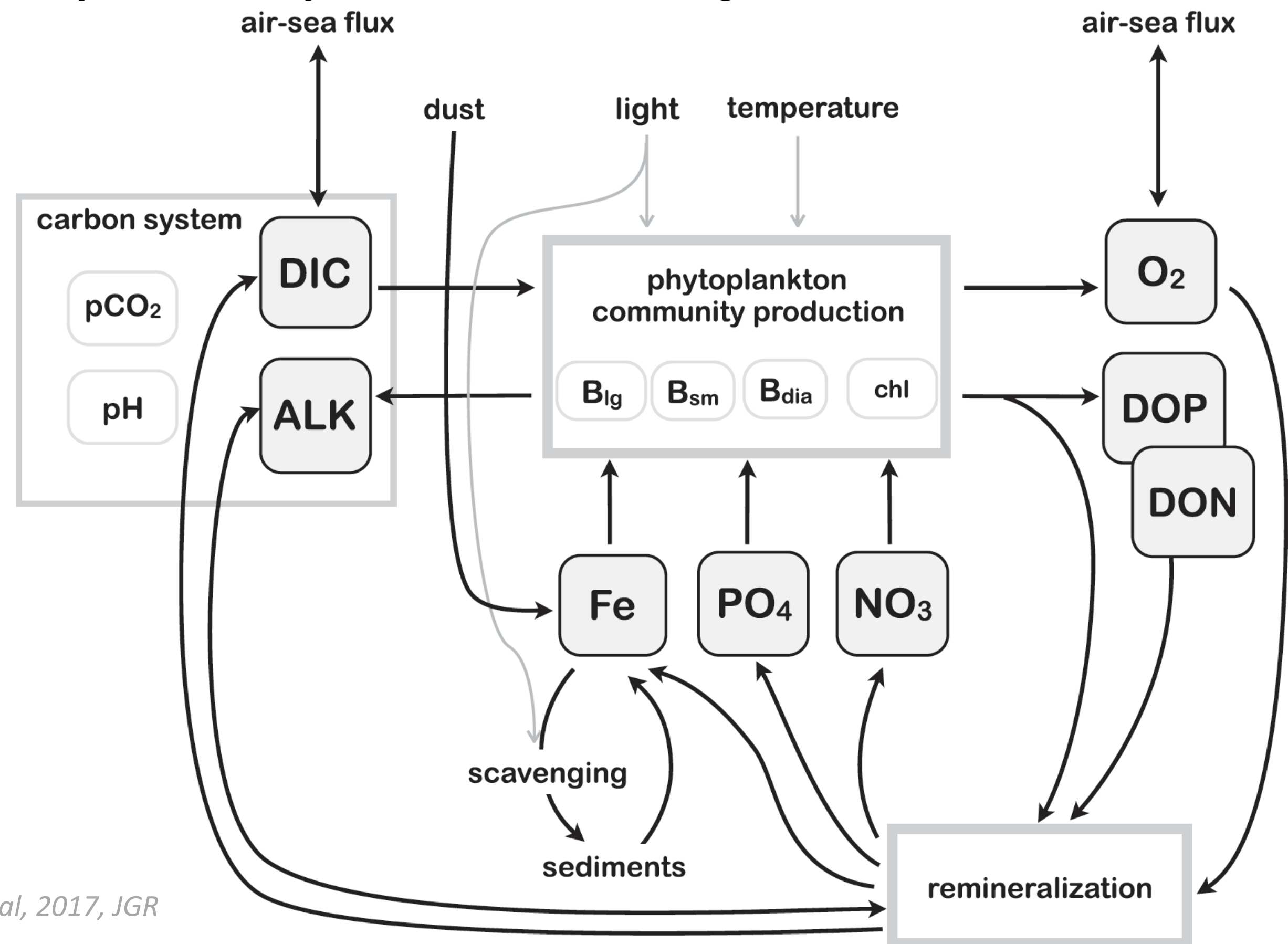
(1) Quantify nitrate (NO_3) reshuffling across Southern Ocean water masses

(2) Assess bottleneck processes



B-SOSE (Biogeochemical Southern Ocean State Estimate) Mazloff et al, 2010, JPO; Rosso et al, 2017, Verdy et al, 2017, JGR

- Ocean circulation model based on MITgcm
biogeochemical (BGC) model: BLING with added N-cycle complexity
- Assimilates observations (including BGC Argo floats), dynamically consistent, budgets close
- Spatial resolution: $1/3^\circ$, 52 levels
- 5-day averages from 2008 to 2017

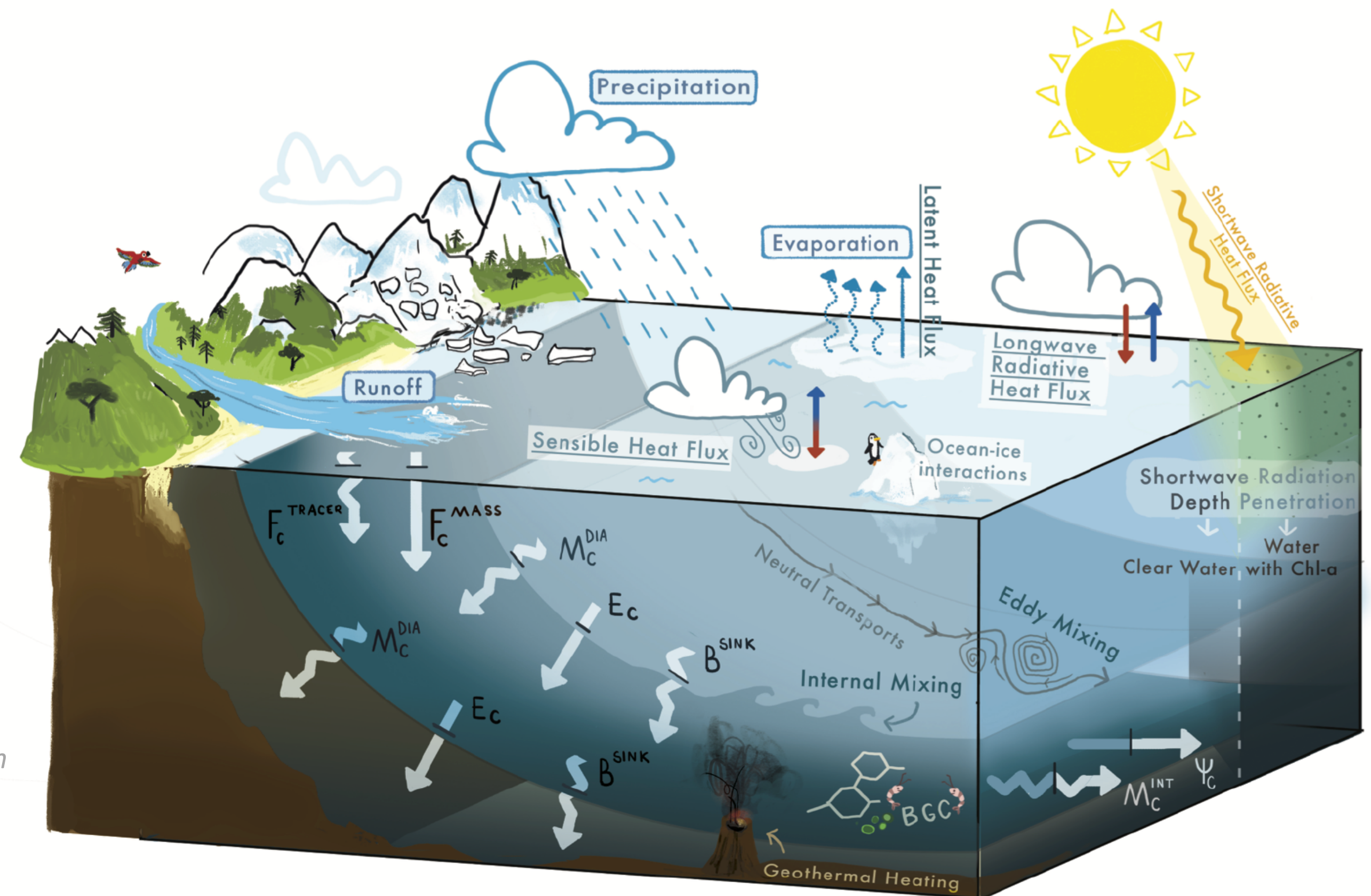


Schematic of BLING from Verdy et al, 2017, JGR

Combination of

- What drives the nitrate (NO_3) water mass inventories and their changes? -> **Based on NO_3 budget**
 NO_3 is driven by advection (overturning), mixing and biological uptake and remineralization Rosso et al, 2017, JGR
- What drives NO_3 fluxes associated with the overturning? -> **Based on water mass transformation analysis**

Iudicone et al, 2008, JPO; Iudicone et al, 2011, Biogeosci.; Walin, 1981, Tellus; Marshall et al, 1999



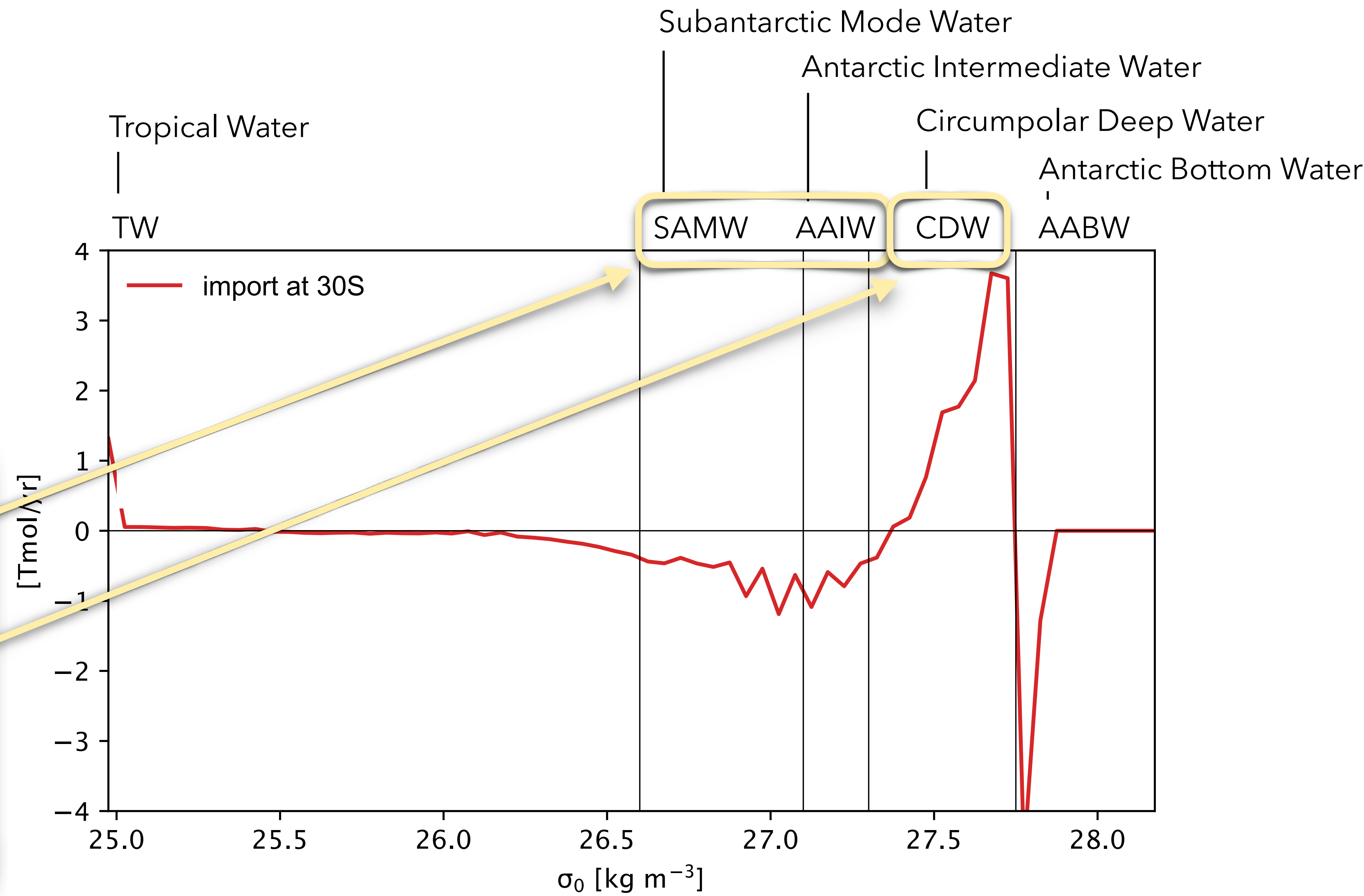
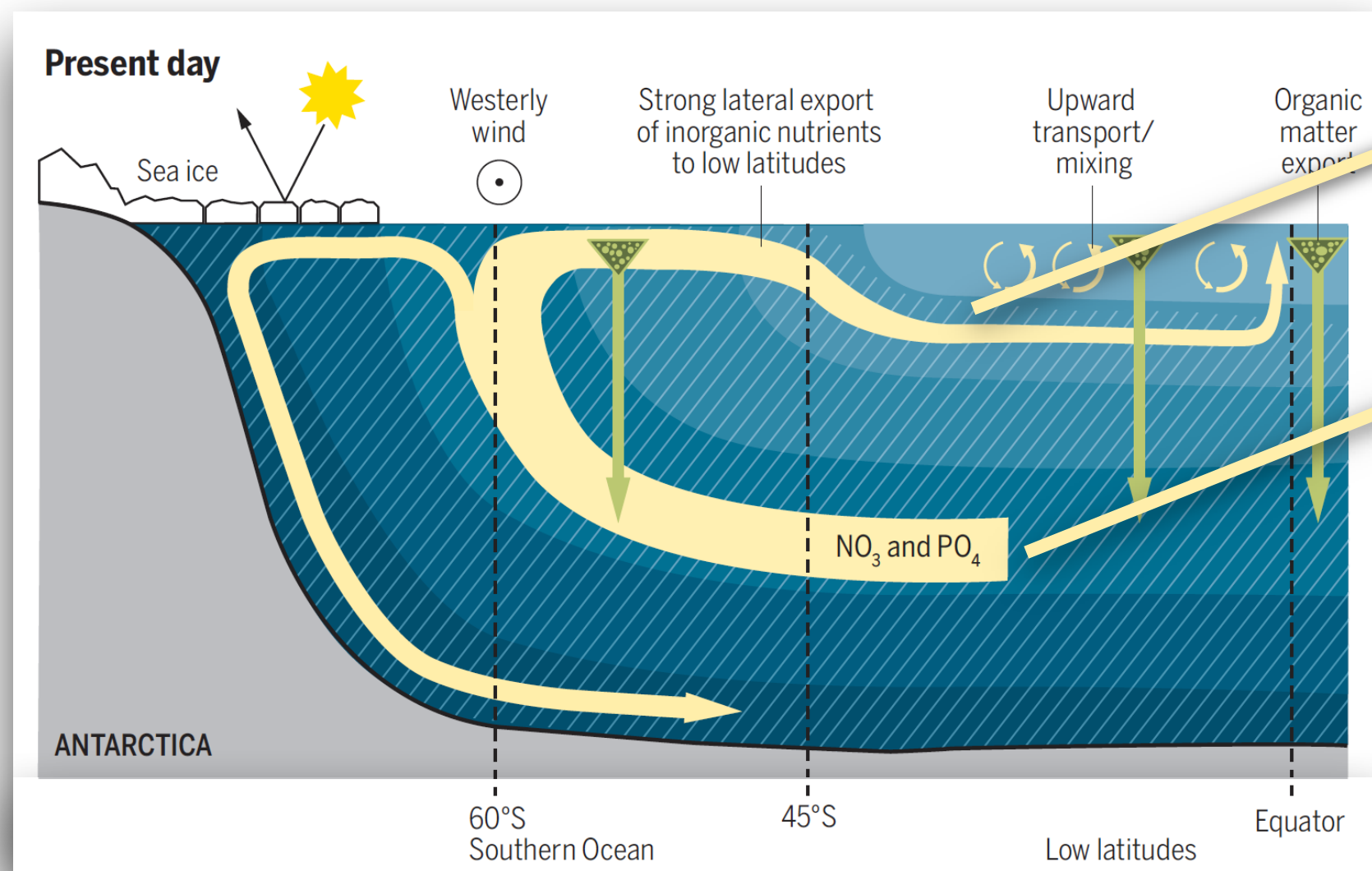
For the mathematical derivation of the combination of water mass transformation analysis with tracers (NO_3) see, e.g., the review by Groeskamp et al, 2019, Ann. Rev. of Mar. Sci. (see also Figure to the right)

... RESULTS: NO₃ import at 30S

Example quantification: **Import in CDW** and negative import, i.e., **export in SAMW/AAIW**

- ▶ CDW imports 14 Tmol/yr
- ▶ SAMW + AAIW export 9 Tmol N/yr, sufficient to support the carbon sequestration flux below 2000m (0.66 Gt C/yr, i.e., ~8 Tmol N/yr)

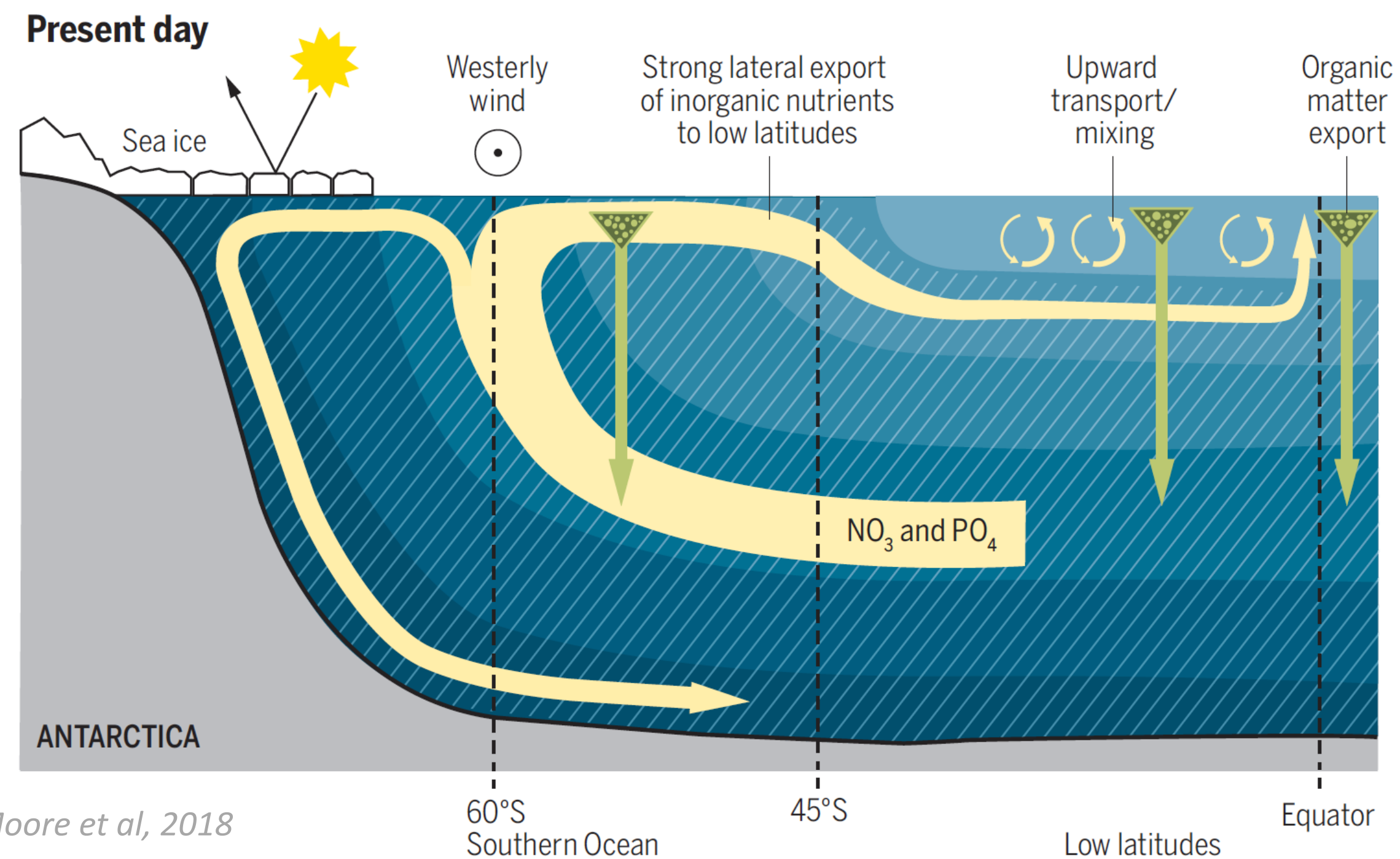
Henson et al, 2012, GBC



(1) Quantify nitrate (NO_3) reshuffling across Southern Ocean water masses

Example quantification: 9 Tmol N/yr exported in SAMW (2/3) + AAIW (1/3)

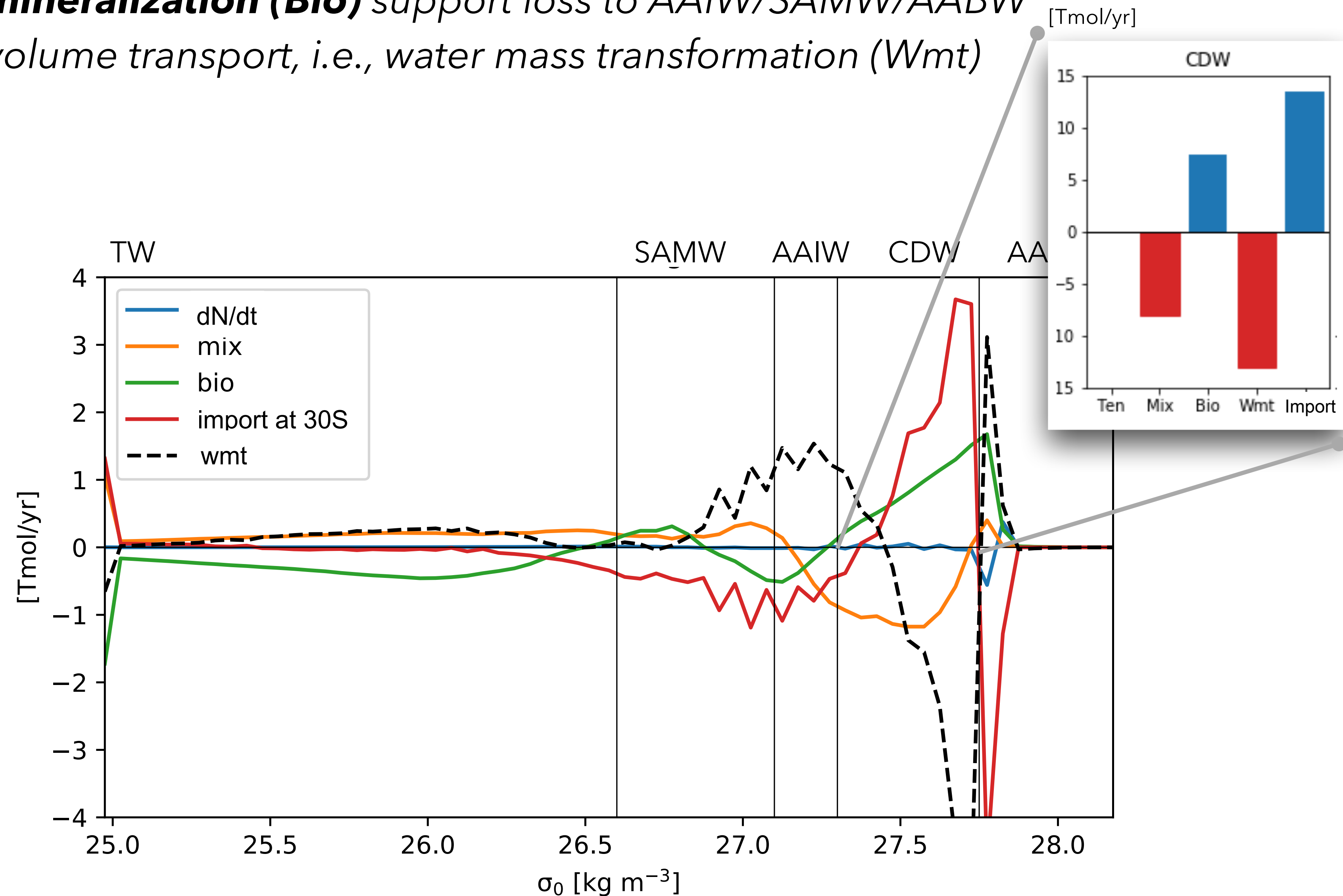
(2) Assess bottleneck processes



... RESULTS: bottleneck processes

CDW

- **60% import** of NO_3 & **30% remineralization (Bio)** support loss to AAIW/SAMW/AABW via mixing (Mix) and diapycnal volume transport, i.e., water mass transformation (Wmt)



... RESULTS: bottleneck processes

CDW

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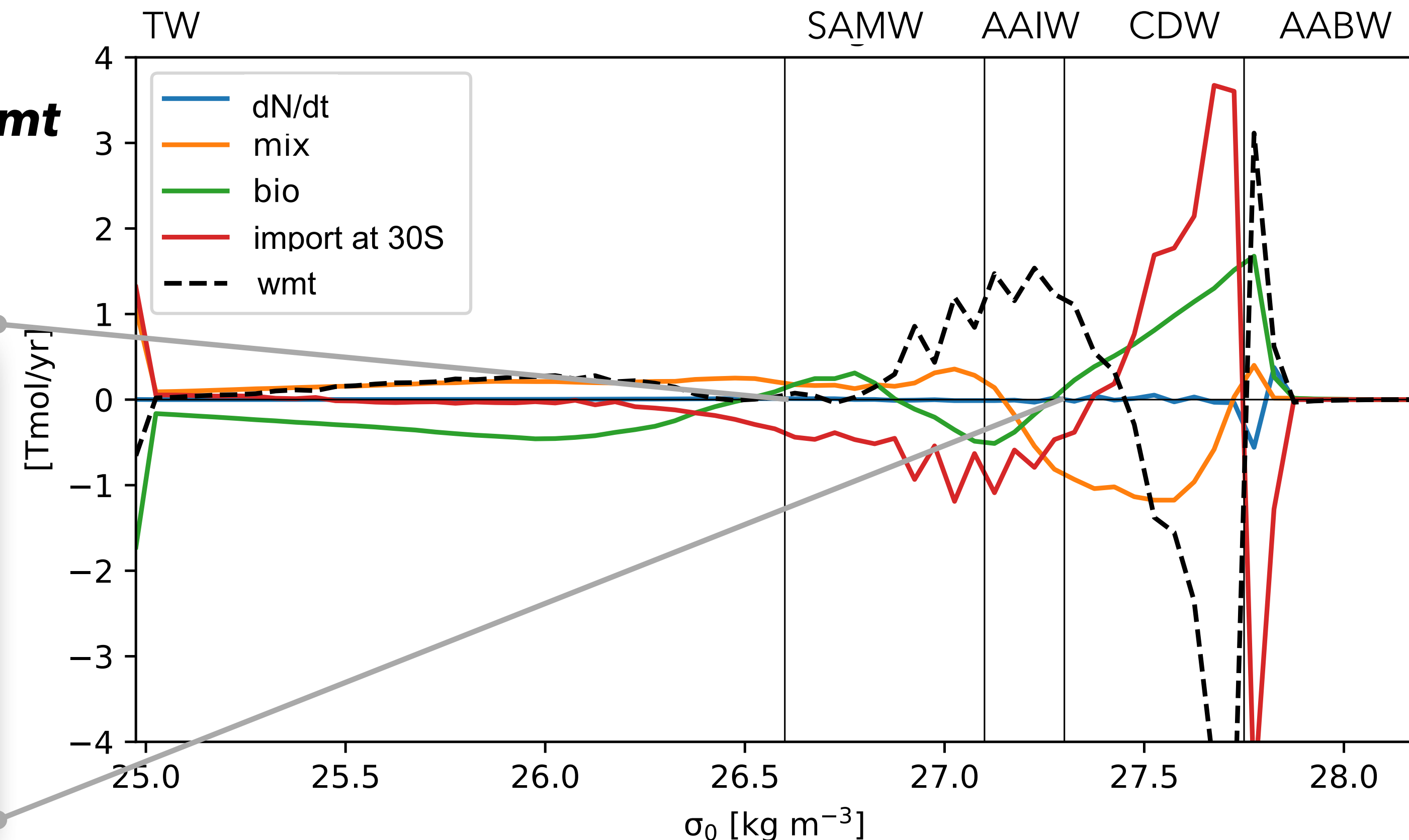
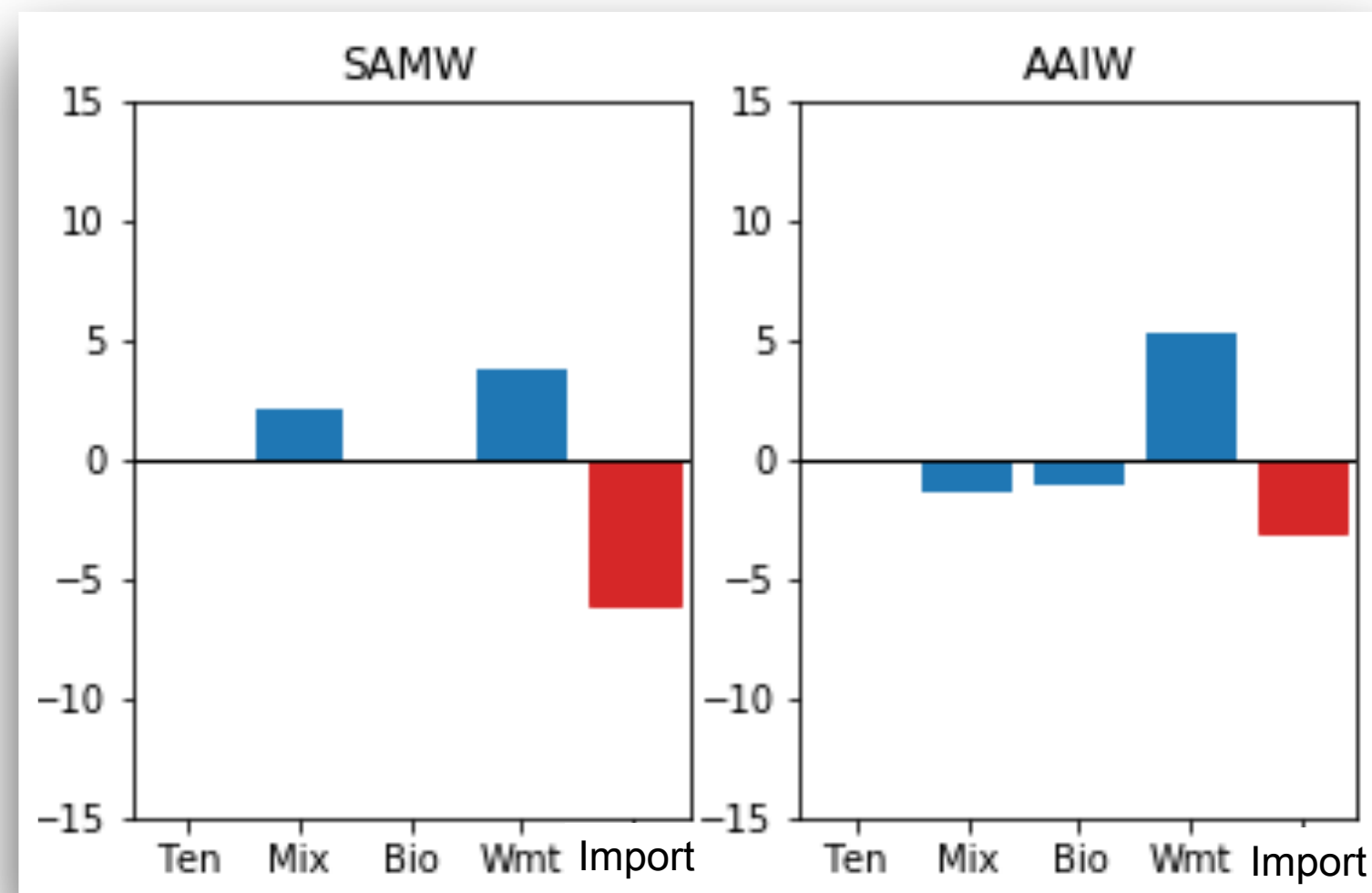
SAMW

- Export of NO_3 supported by **35% Mix** & **65% Wmt**
(driven by surface fluxes)

AAIW

- Export of NO_3 supported by **Wmt**
(driven by density mixing)

[Tmol/yr]



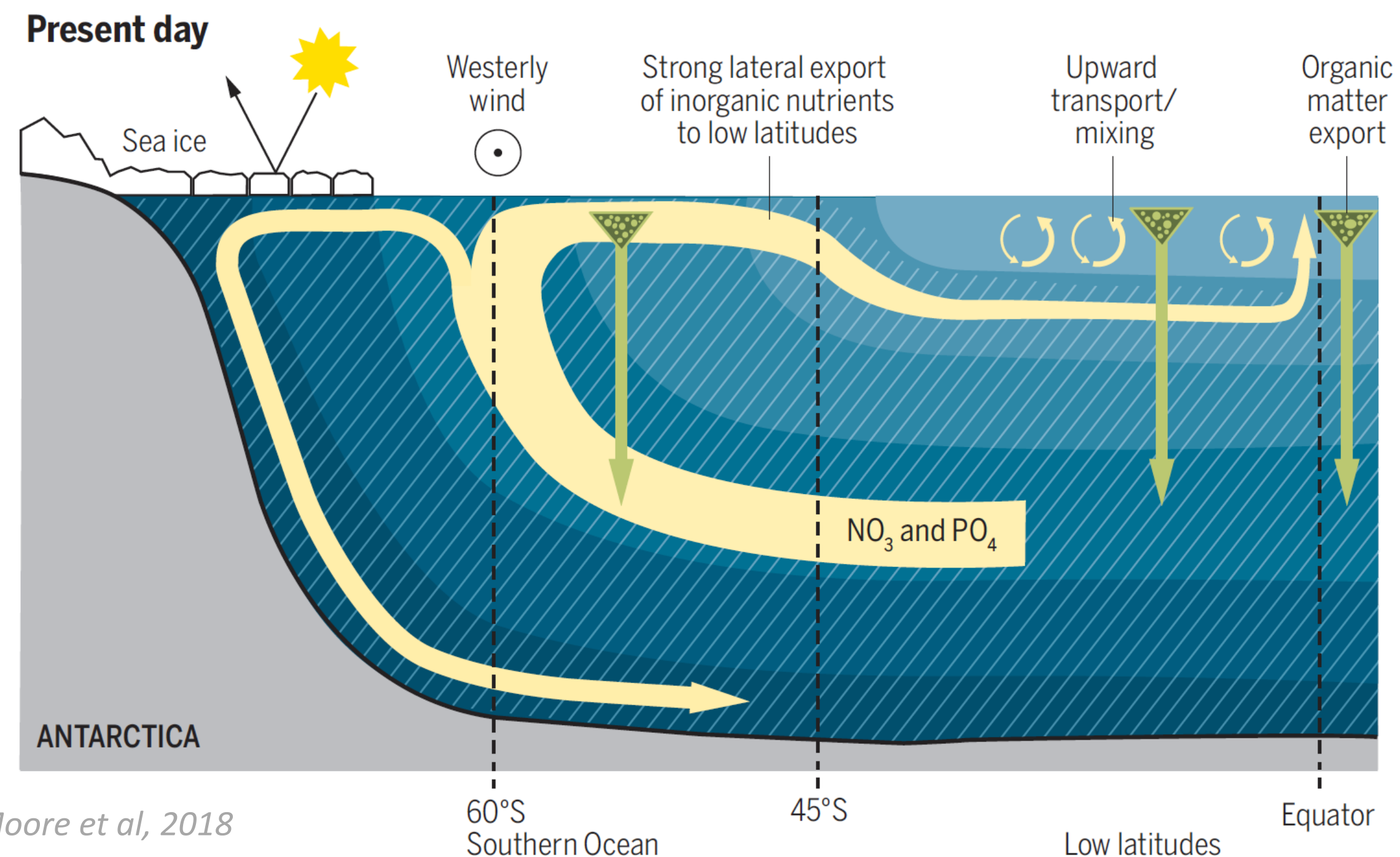
(1) Quantify nitrate (NO_3) reshuffling across Southern Ocean water masses

Example quantification: 9 Tmol N/yr exported in SAMW (2/3) + AAIW (1/3)

(2) Assess bottleneck processes

Example assessment: Support of SAMW and AAIW export:

Next to water mass transformation, i.e., overturning (driven mainly by air-sea fluxes), roughly 1/4 tracer mixing



... NEXT STEPS

- Update of data to higher resolution iteration of B-SOSE ($1/6^\circ$)
- Assessment of spatio-temporal variability, including mixed layer versus interior ocean processes

Example: Variability of net biological fluxes (Bio) with latitude and mixed layer vs interior ocean

