

Pressure variations in the Monte Rosa nappe: new results from staurolite bearing metapelites

Joshua D. Vaughan-Hammon*, Cindy Luisier, Lukas P. Baumgartner, Stefan M. Schmalholz

*Institute of Earth Sciences, University of Lausanne, Lausanne 1015, Switzerland

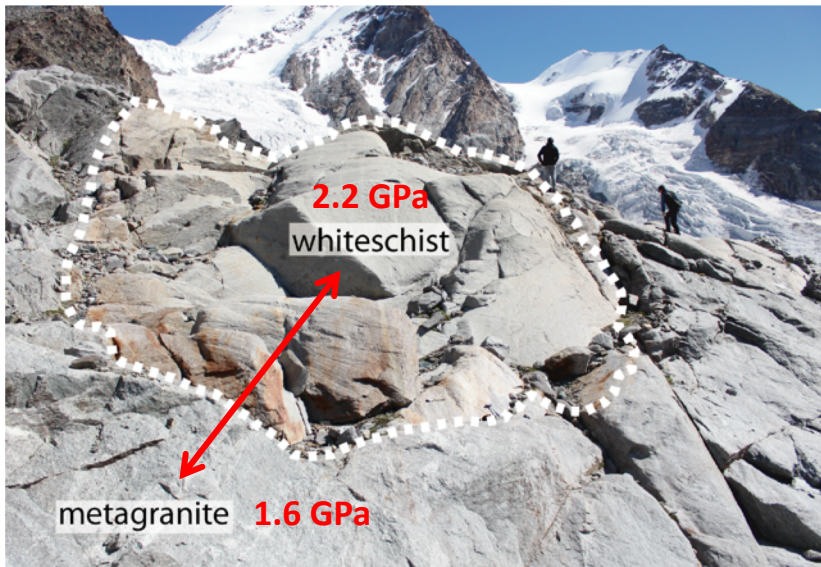
joshua.vaughan-hammon@unil.ch

ARTICLE

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Metamorphic pressure variation in a coherent Alpine nappe challenges lithostatic pressure paradigm

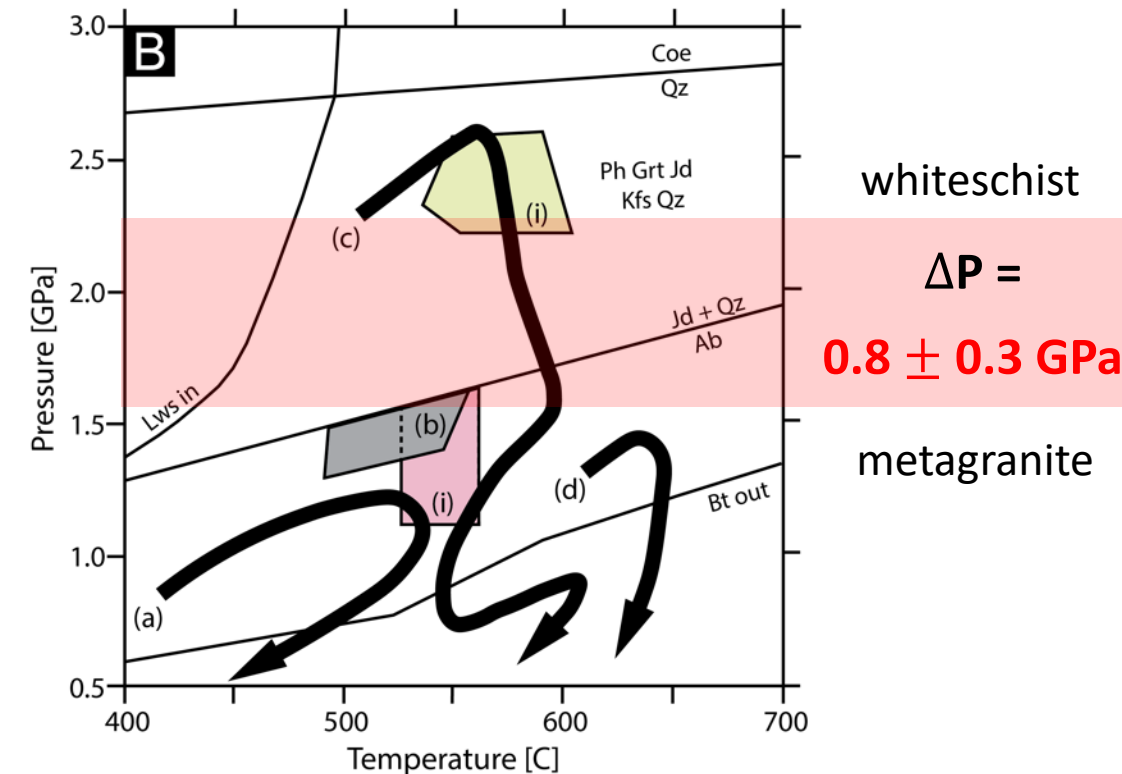
Cindy Luisier ^{1*}, Lukas Baumgartner¹, Stefan M. Schmalholz¹, Guillaume Siron^{1,2} & Torsten Vennemann³**Alpine peak pressure:****Whiteschist: ca 2.2 GPa****Metagranite: ca 1.6 GPa****Why pressure difference?****Possible explanations:**

- 1) Tectonic mélange
- 2) Granite did not record peak-P
 - (a) sluggish kinetics
 - (b) retrogression
- 3) Thermodynamic database
- 4) Mechanical P variations

Aims of this study:

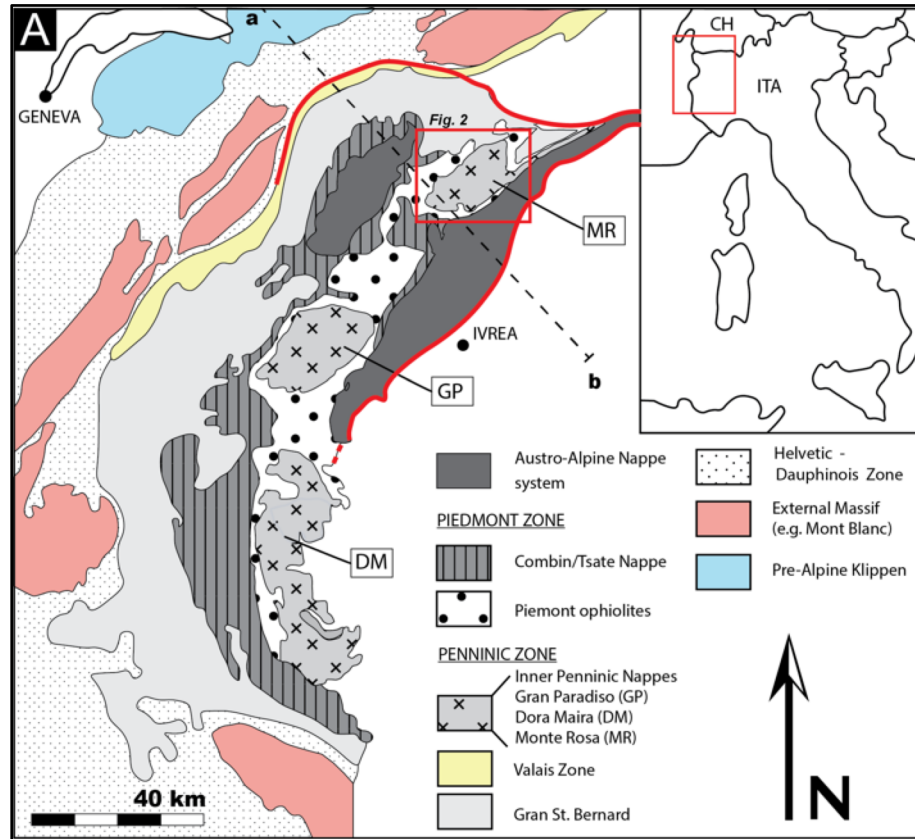
(*manuscript under review)

- Continue to investigate P-variations in the Monte Rosa nappe
- Analyse basement metapelite samples:
Newly discovered peak Alpine assemblages
Calculate P and T
- Pressure variations:
Mechanically induced
Chemically induced
- Geodynamic implications

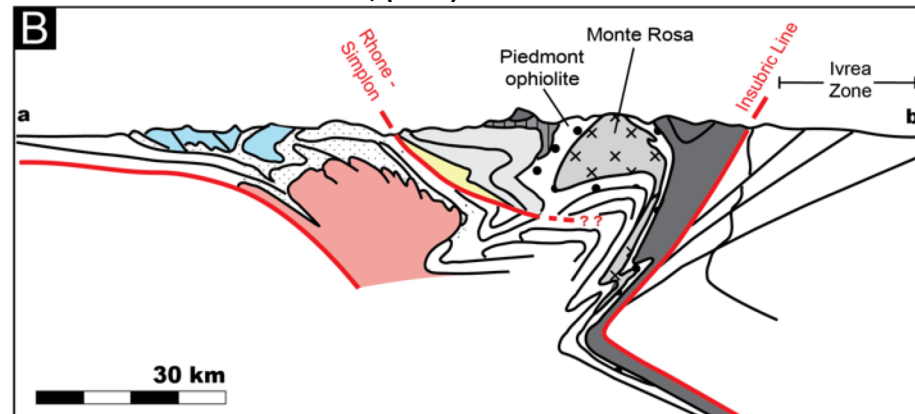


- (a) Borghi et al., 2004
- (b) Dal Piaz and Lombardo 1986
- (c) Gasco et al., 2011
- (d) Keller et al., 2004
- (e) Le Bayon et al., 2006
- (f) Lapen et al., 2007
- (g) Le Bayon et al., 2006
- (h) Chopin and Monie 1984
- (i) Luisier et al., 2019

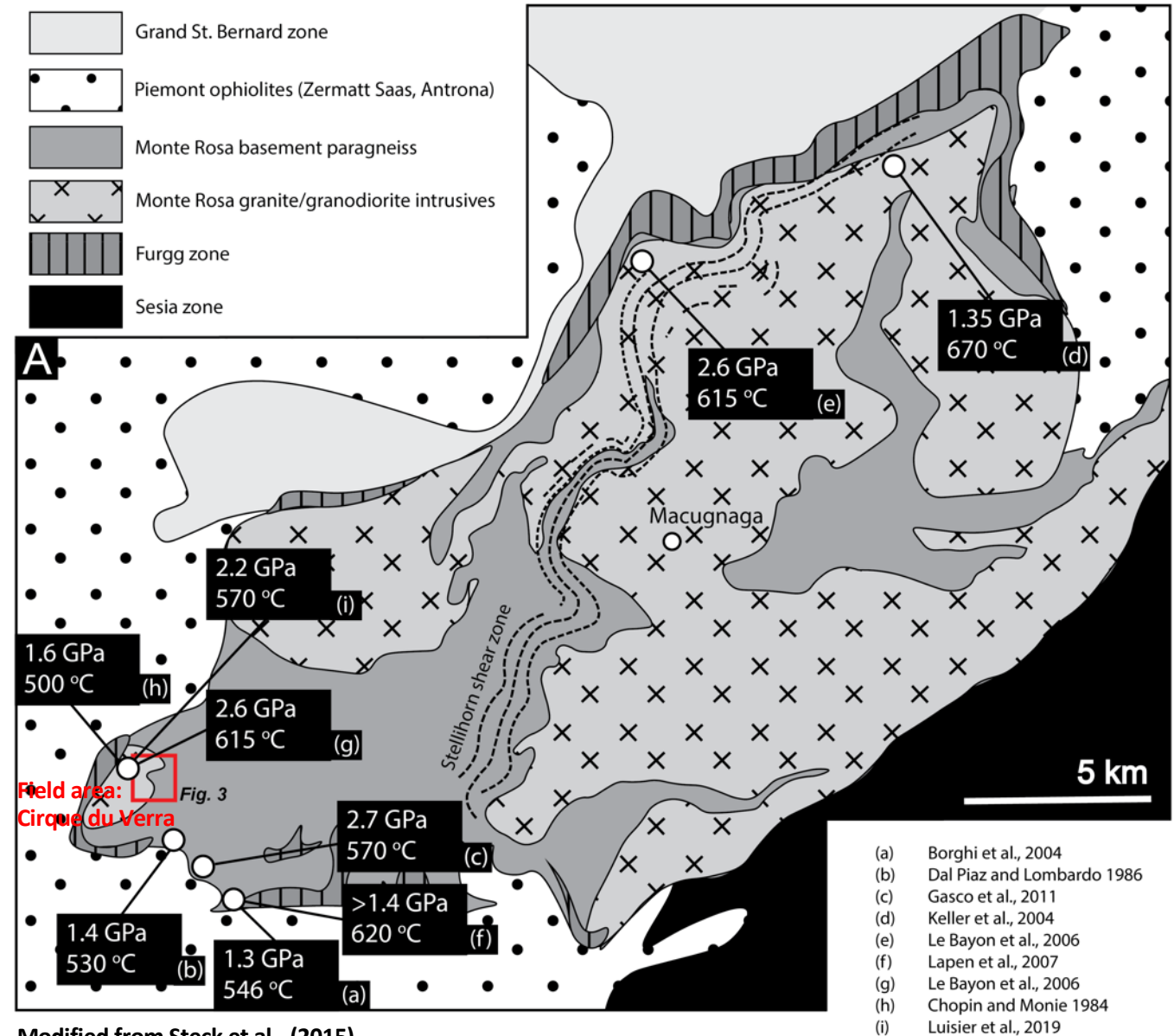
Tectonic location of the Monte Rosa:



Modified from Beltrando et al., (2010)



Modified from Steck et al., (2015)



Modified from Steck et al., (2015)

Field area: "Cirque du Verra"

- Far western extent of Monte Rosa nappe
- Recent glacial retreat has uncovered fresh exposure for mapping



Metapelite sample: outcrop observations

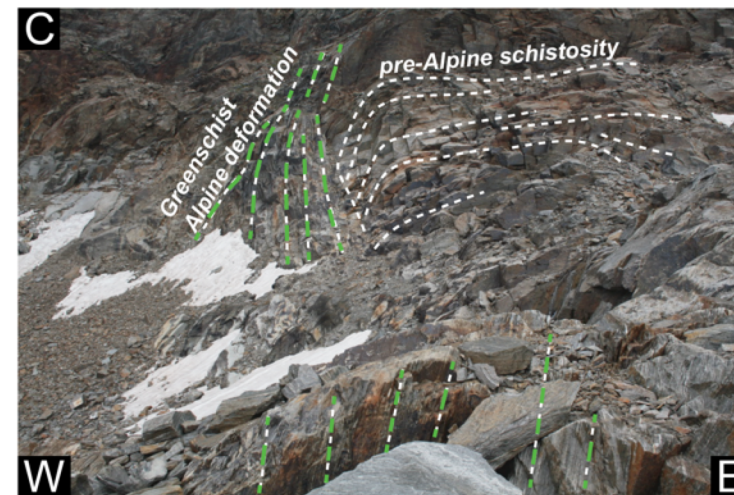
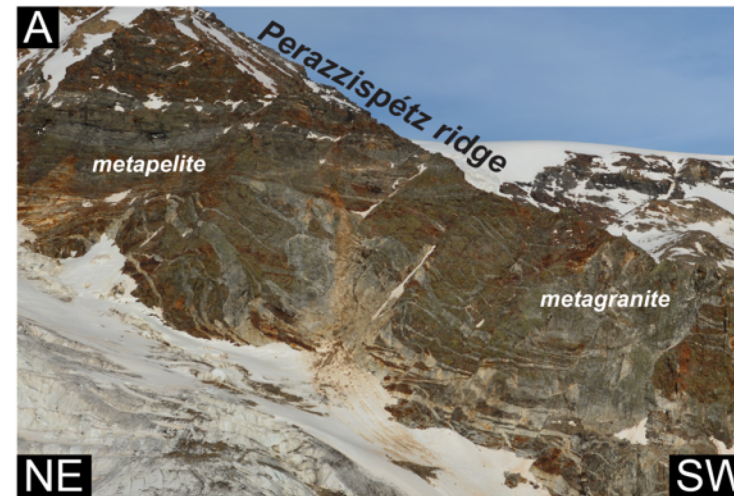
A - Large scale igneous textures.

B - Igneous contact => coherent unit and therefore

not a tectonic mélange.

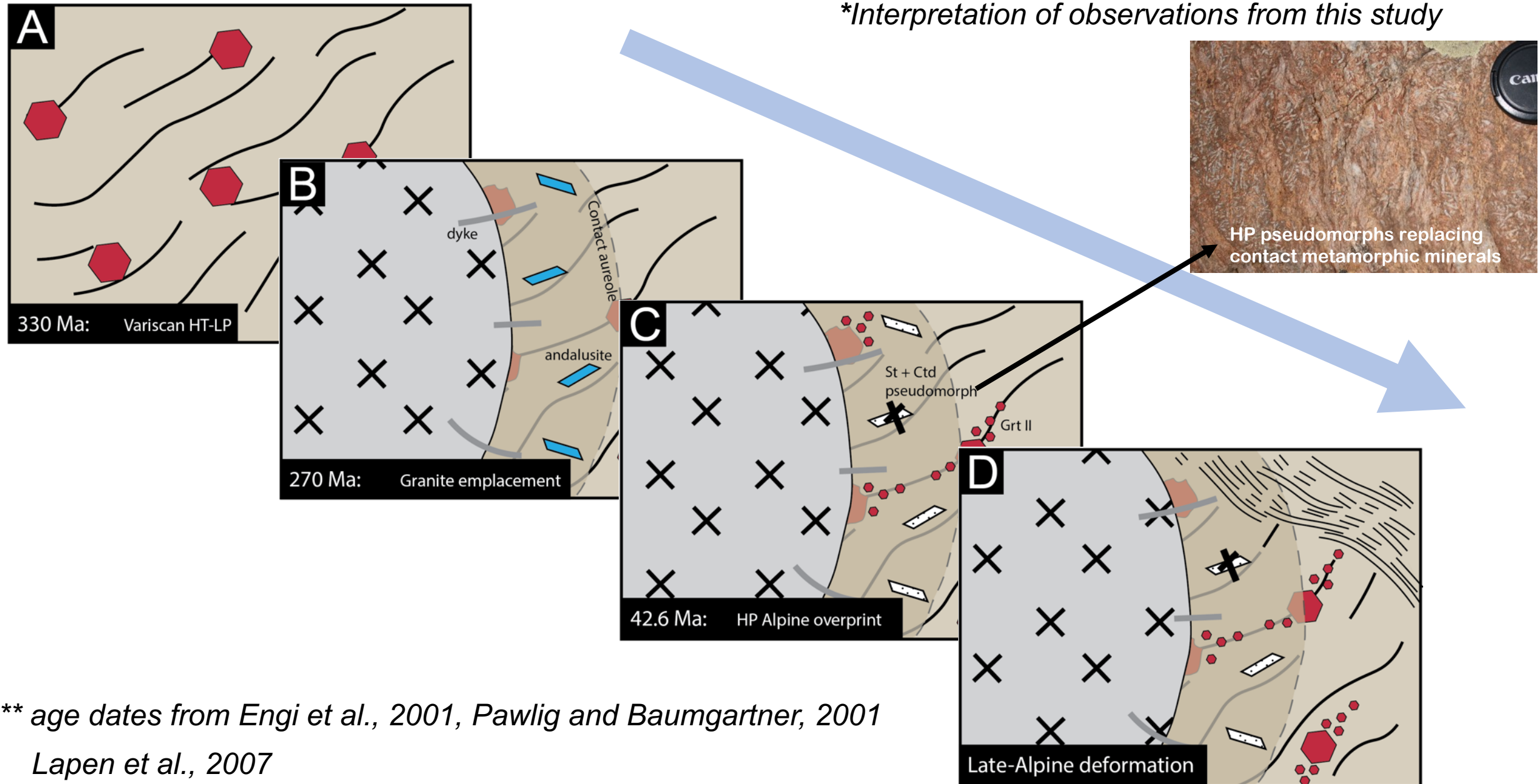
C - High pressure Alpine imprint is preserved within a pre-Alpine schistosity unaffected by late-Alpine greenschist overprinting.

D - High pressure assemblages within pseudomorphs replacing former contact metamorphic andalusite....



Schematic geological history of metapelite:

**Interpretation of observations from this study*



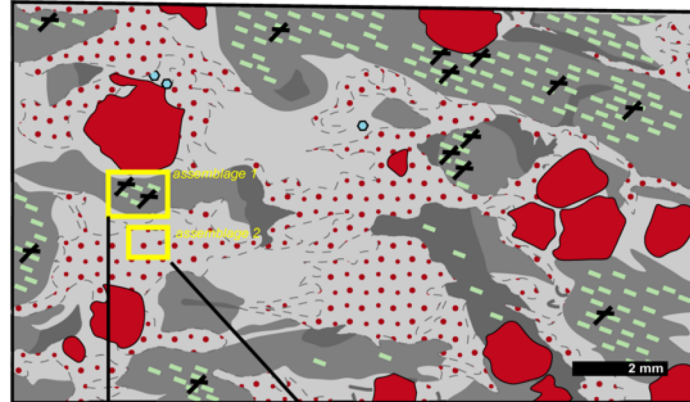
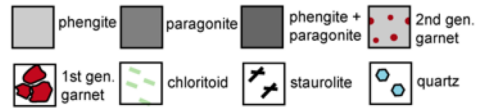
****** age dates from Engi et al., 2001, Pawlig and Baumgartner, 2001
Lapen et al., 2007

Metapelite petrology: 3x equilibrium assemblages

Sample 16MR-17 & 19MR-04: unique staurolite + chloritoid bearing assemblages:

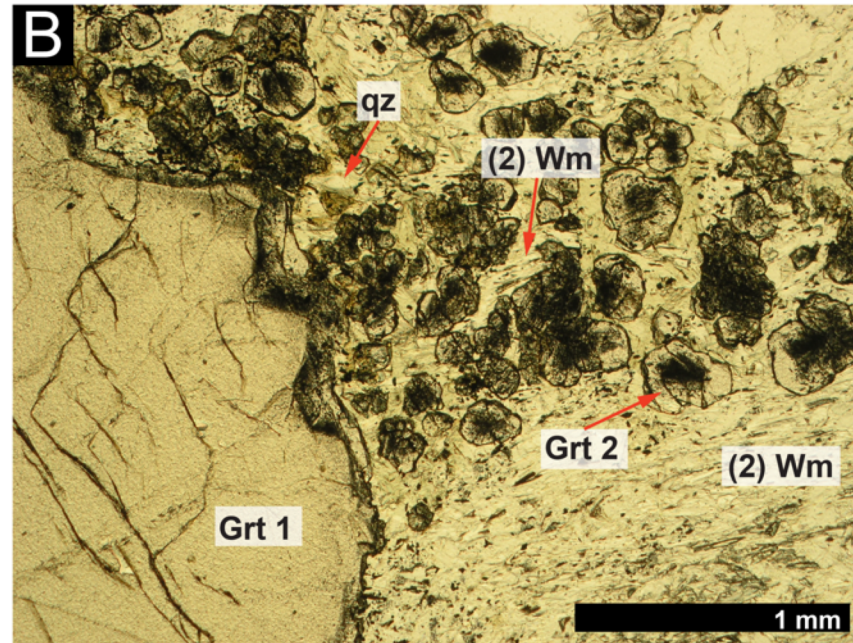
Grt + Ms + Pg + Cld + St + Chl + Bt + Qtz + Als (+ accessory mineral Ap, Rt and Mz)

- Representing peak Alpine metamorphism
- Water saturated conditions (no sluggish kinetics)

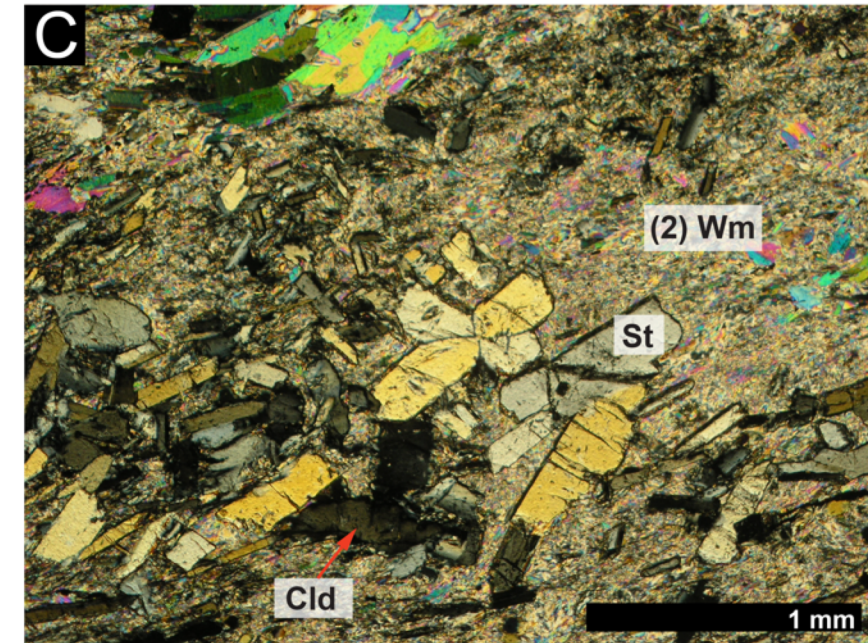


Assemblage 1:
St + Cld + Ms + Pg

Assemblage 2:
Grt + Chl + Ms + Pg



Assemblage 2:
Grt + Chl + Ms + Pg



Assemblage 1:
St + Cld + Ms + Pg

Assemblage 3:
Cld + Ms + Pg

**Similar to Assemblage 1
(without St)**

Chemistry of peak metamorphic minerals:

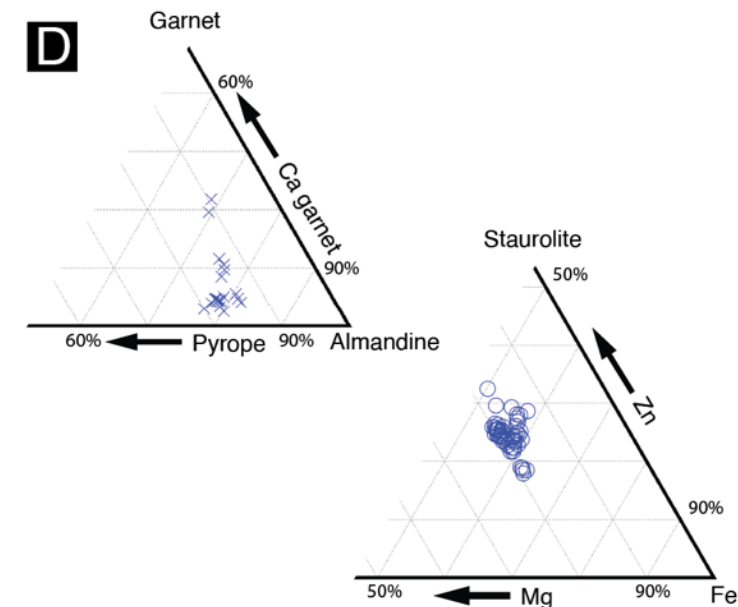
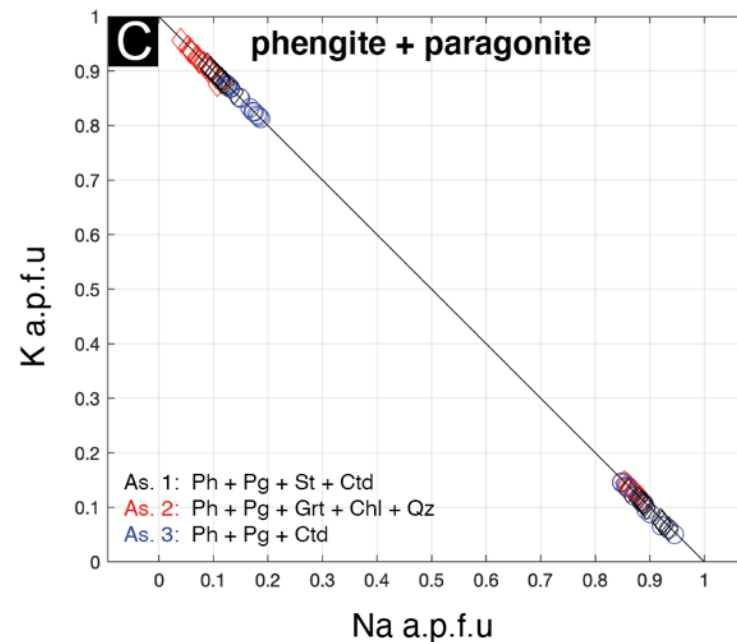
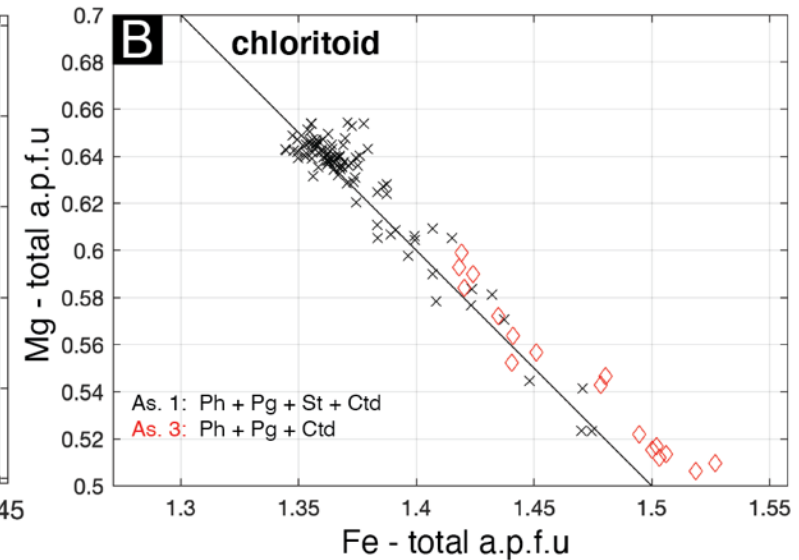
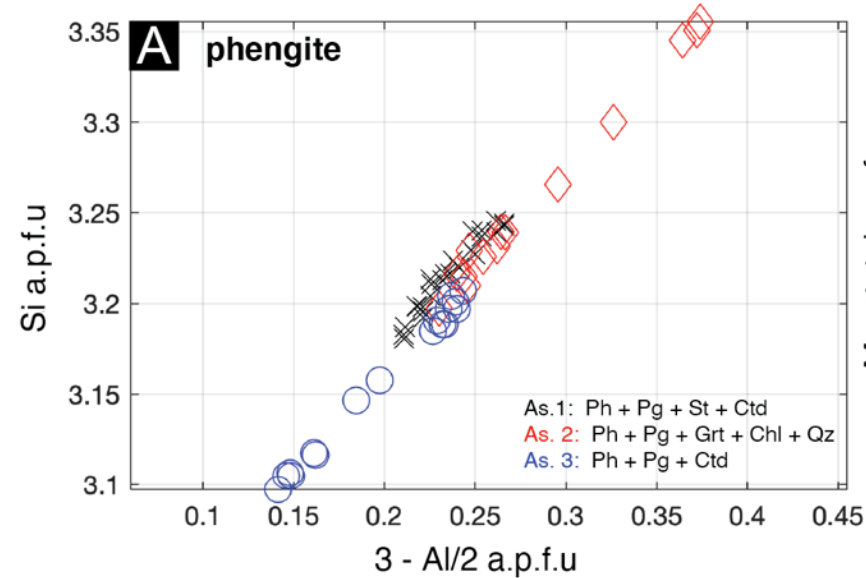
A - Si in phengite.

B - Mg and Fe-total in chloritoid.

C - Na in paragonite and K in phengite mixing gap.

D - Ternary plot for garnet compositions in assemblage 2, and ternary plot for staurolite compositions in assemblage 1.

***note non-negligible Zn**



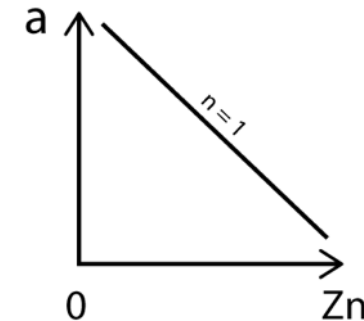
Zn in staurolite activity reduction:

- In order to account for the lack of solution models for Zn in staurolite we have employed a method to adjust the activity of available solid solution end-member data.
- Only Mg and Fe end-member data is available, therefore an entropy adjustment is needed:

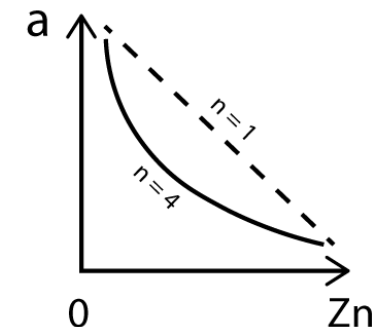
$$S^{corr} = S^o - R \ln a$$

- $a = 1$ for a pure phase.
- Site multiplicity of staurolite being 4 ($\text{Fe}^{2+} = \text{Mg} = \text{Zn} = \text{Mn}$).

Molecular mixing model: $a = XMg = \left(1 - \left(\frac{Zn}{4}\right)\right)$

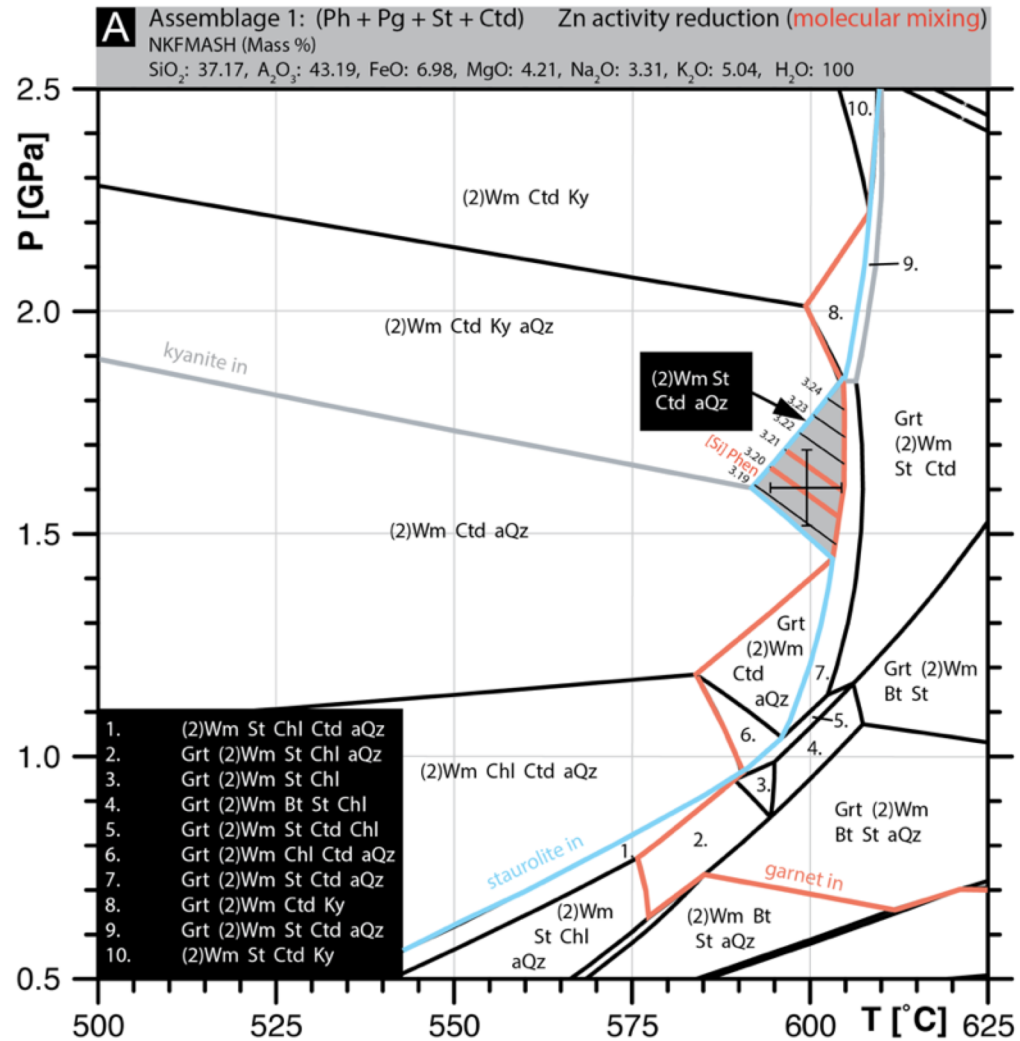


Site mixing model: $a = XMg = \left(1 - \left(\frac{Zn}{4}\right)\right)^4$

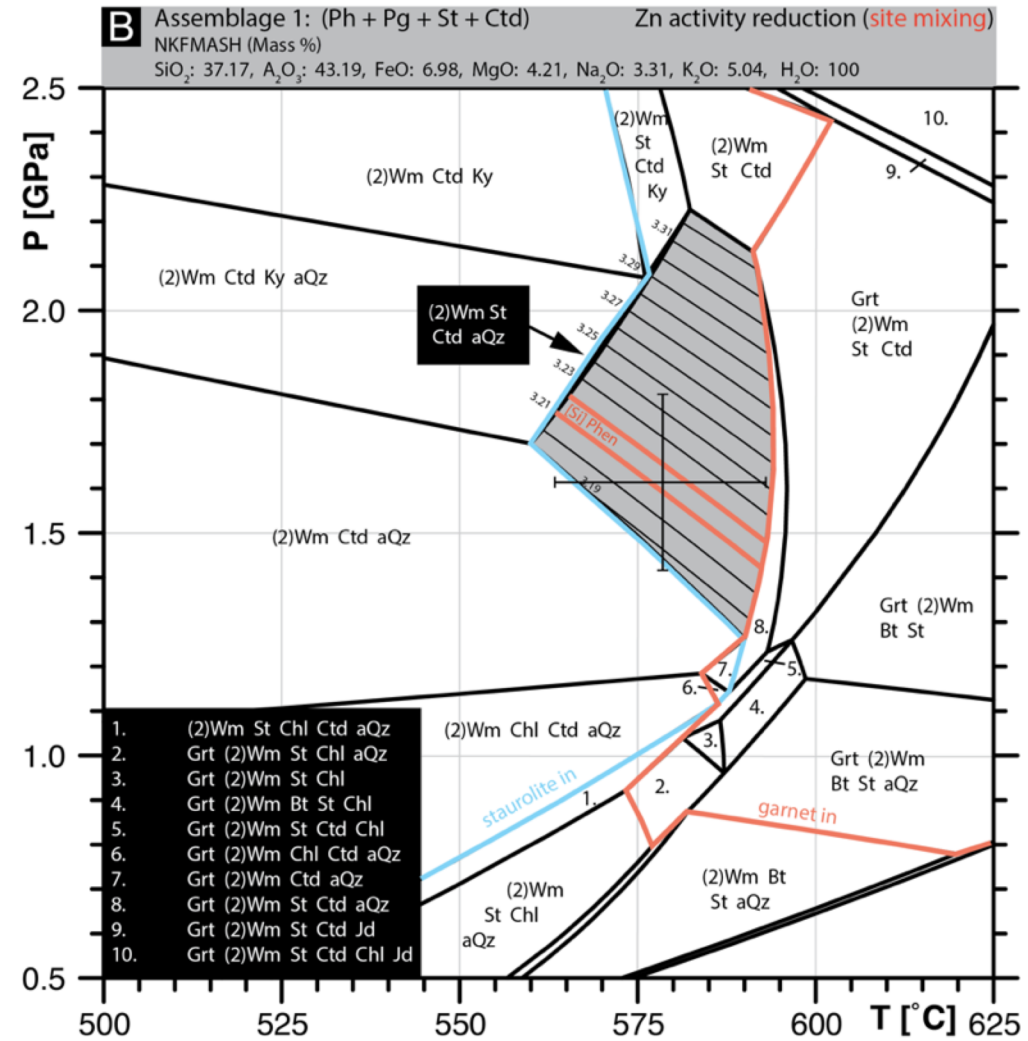


Pseudosection results: assemblage 1

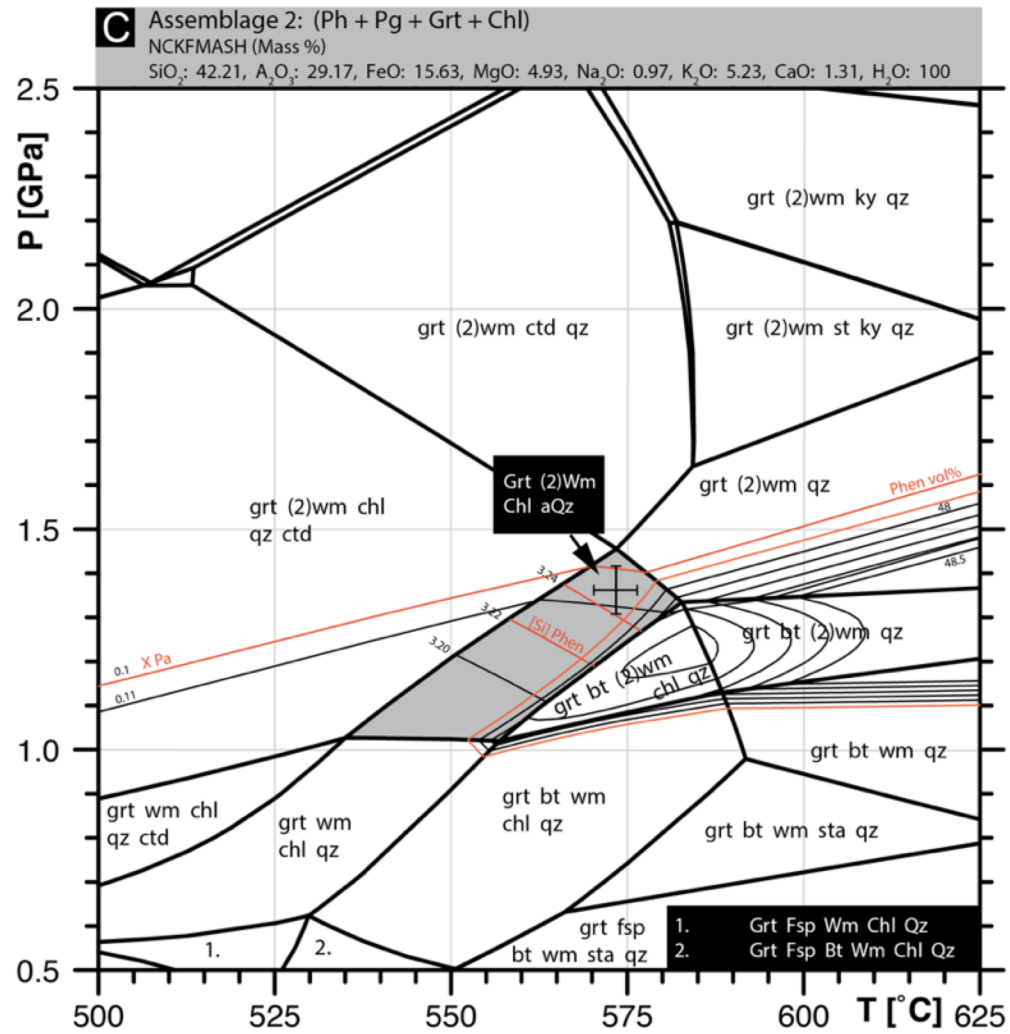
Molecular mixing:



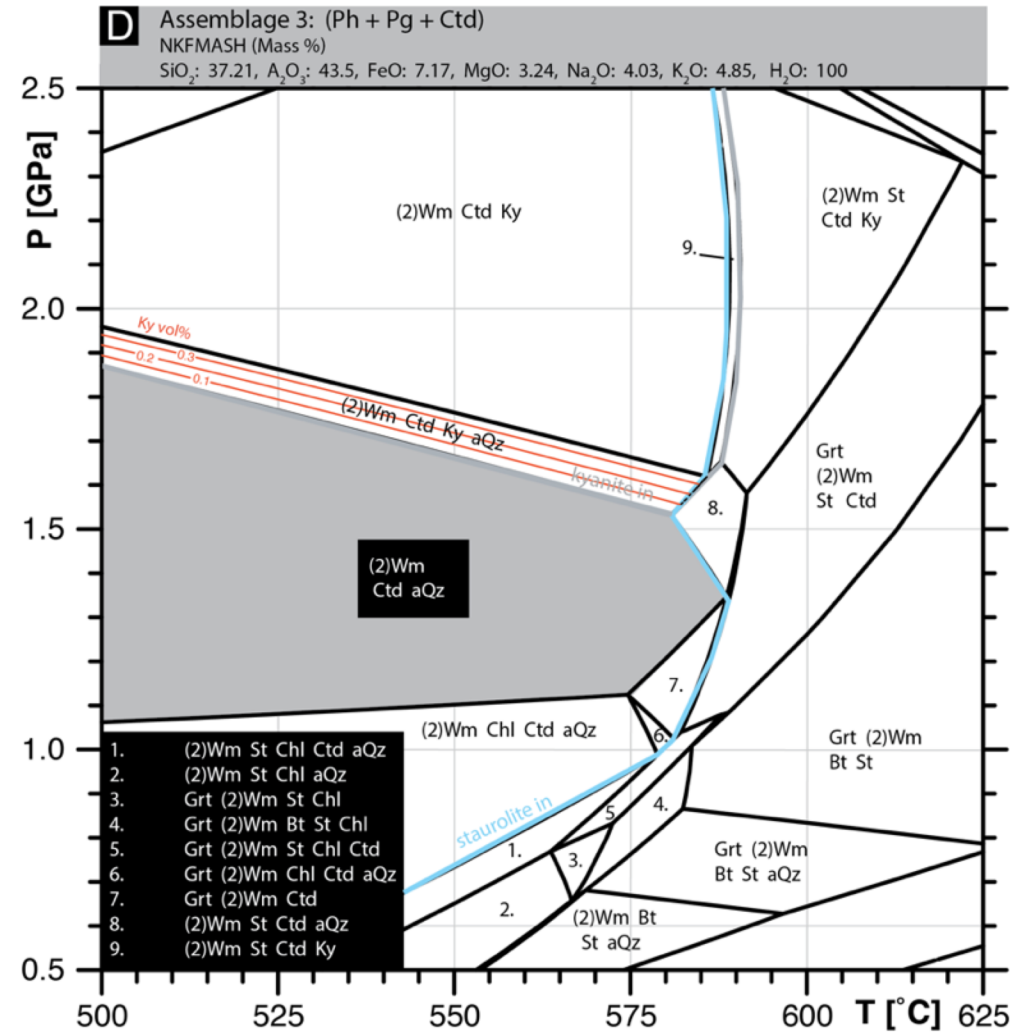
Site mixing:



Pseudosection results: assemblages 2 and 3

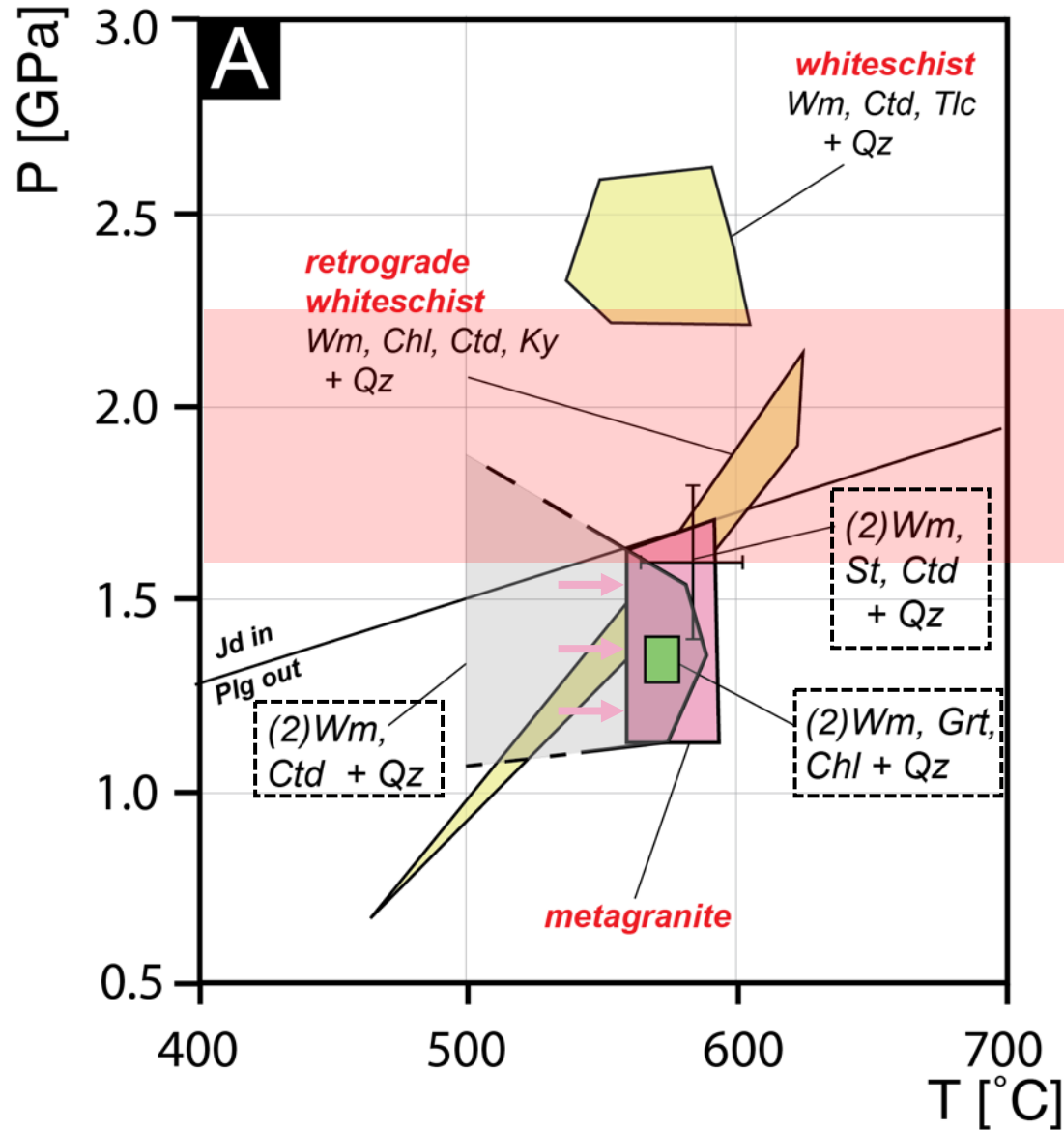


**1.3 – 1.4 GPa,
ca. 575 °C**



**ca. 1.6 GPa
at 575 °C**

Comparison with whiteschist:



$$\Delta P = 0.6 \pm 0.2 \text{ GPa}$$

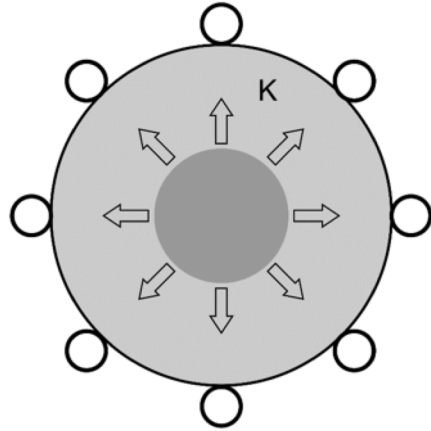
- Peak Metamorphic conditions from metapelitic samples re-affirm pressure variations
- Whiteschist is consistently at a higher pressures compared to all metagranite and metapelite lithologies examined
- Varying P but consistently similar T conditions => isothermal decompression?
- Rapid isothermal exhumation, or mechanical P variations?

Mechanical P variation:the options

Reaction-induced stress

Dehydration under isochoric conditions

0.4 – 0.5 GPa

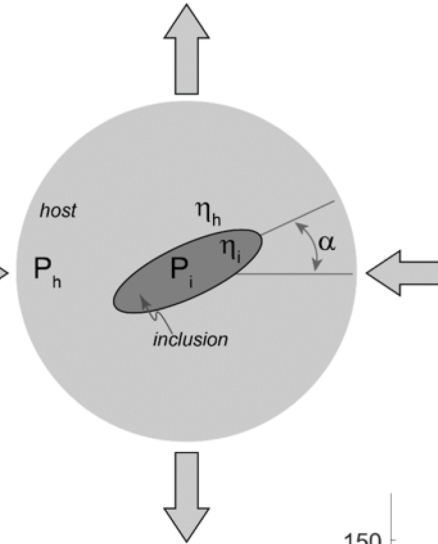


In reality it can be a mixture of both processes...

What is needed are **rheological heterogeneities**

Compression-induced stress

Weak inclusion in strong host

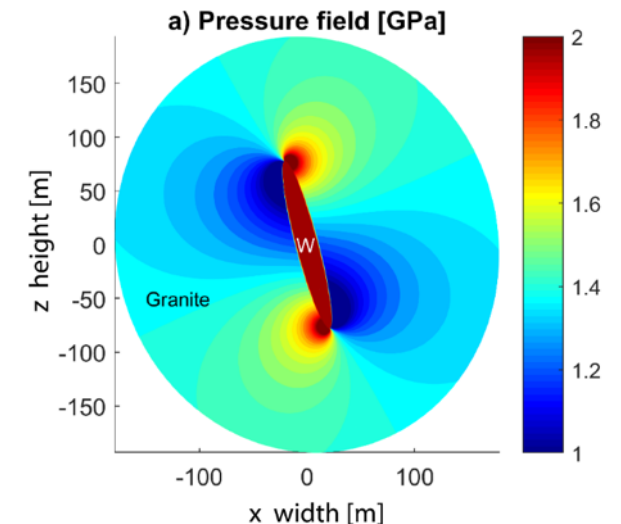
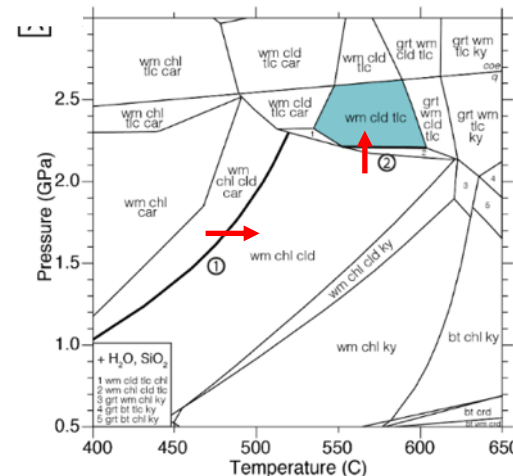


$\Delta P \sim T_{xx}$

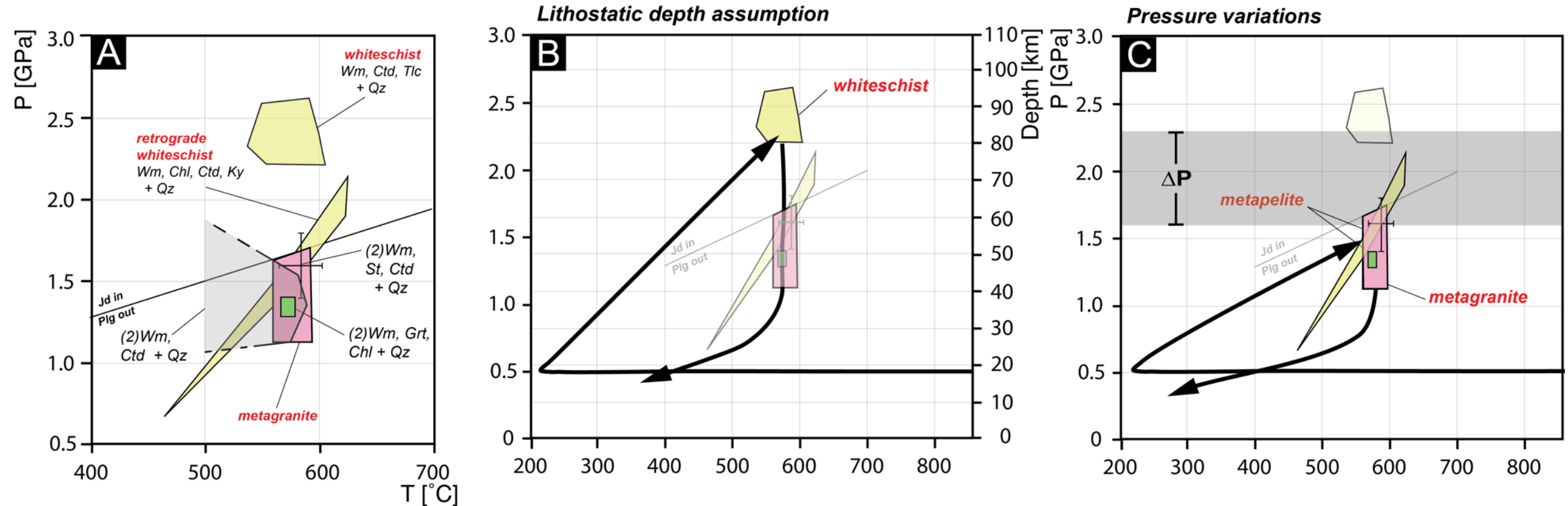
$$\Delta V_{tot} = \Delta V_{ela} + \Delta V_{rea} = 0$$

$$\Delta p = K \frac{\Delta V_{ela}}{\Delta V_0} = -K \frac{\Delta V_{rea}}{\Delta V_0}$$

$$\Delta p = -\frac{1}{\beta} \frac{\Delta V_{rea}}{\Delta V_0} = -\rho \frac{dp}{d\rho} \frac{\Delta V_{rea}}{\Delta V_0}$$



Geodynamic implications:



- **Deep subduction**
- **Mechanically weak and homogeneous rock unit**

- **Moderate subduction**
- **Mechanically heterogeneous rock unit**

- Lithostatic pressure
- Pressure directly related to Depth
- Peak P = deepest burial of the Monte Rosa nappe

- P variations
- Whiteschist represents local pressure variations
- Metapelite and Metagranite lithologies represent regional peak pressure of the Monte Rosa nappe

Thank you