

# A daily estimate of phase speed to explore the link between Arctic Amplification and Rossby waves

**Jacopo Riboldi\***, François Lott, Fabio D'Andrea  
and Gwendal Rivière

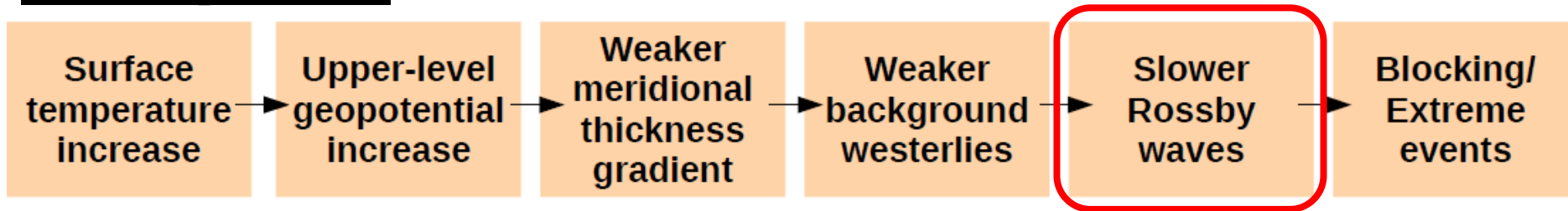
Laboratoire de Météorologie Dynamique, PSL University, ENS/IPSL/CNRS  
Paris, France



# Motivation

Enhanced Arctic warming with respect to midlatitudes (Arctic Amplification) can slow down eastward Rossby wave propagation and induce extreme weather events (Francis and Vavrus 2012, Cohen et al. 2020).

## Chain of processes:



## How to assess whether Rossby waves have become slower?

$$c_p = \frac{\omega}{k} = \frac{\omega a \cos(\phi)}{n}$$

**Problem:** many waves with different  $\omega$  and  $k$  co-exist in the atmosphere: how to obtain a global value?

## Objectives:

- 1) Develop a phase speed diagnostic accounting for midlatitude Rossby waves variability (e.g., blocking).
- 2) Assess whether Arctic Amplification impacted phase speed trends.

# Daily phase speed diagnostic

## Superposition principle:

Large-scale flow evolution results from a superposition of waves across a broad range of wavenumbers and frequencies (and, therefore, of phase speeds).

Spectral analysis tells us how much each  $(n, c_p)$  harmonic contributes to the overall phase speed.

**Phase speed  
metric**

$c$

$$c = \frac{\sum_{n=1}^{15} \sum_{c_p=-30}^{30} S(n, c_p) \cdot c_p}{\sum_{n=1}^{15} \sum_{c_p=-30}^{30} S(n, c_p)}$$

Sum over wavenumbers and phase speeds

Spectral coefficients as weights

$n$  a-dimensional zonal wavenumber

$c_p$  phase speed

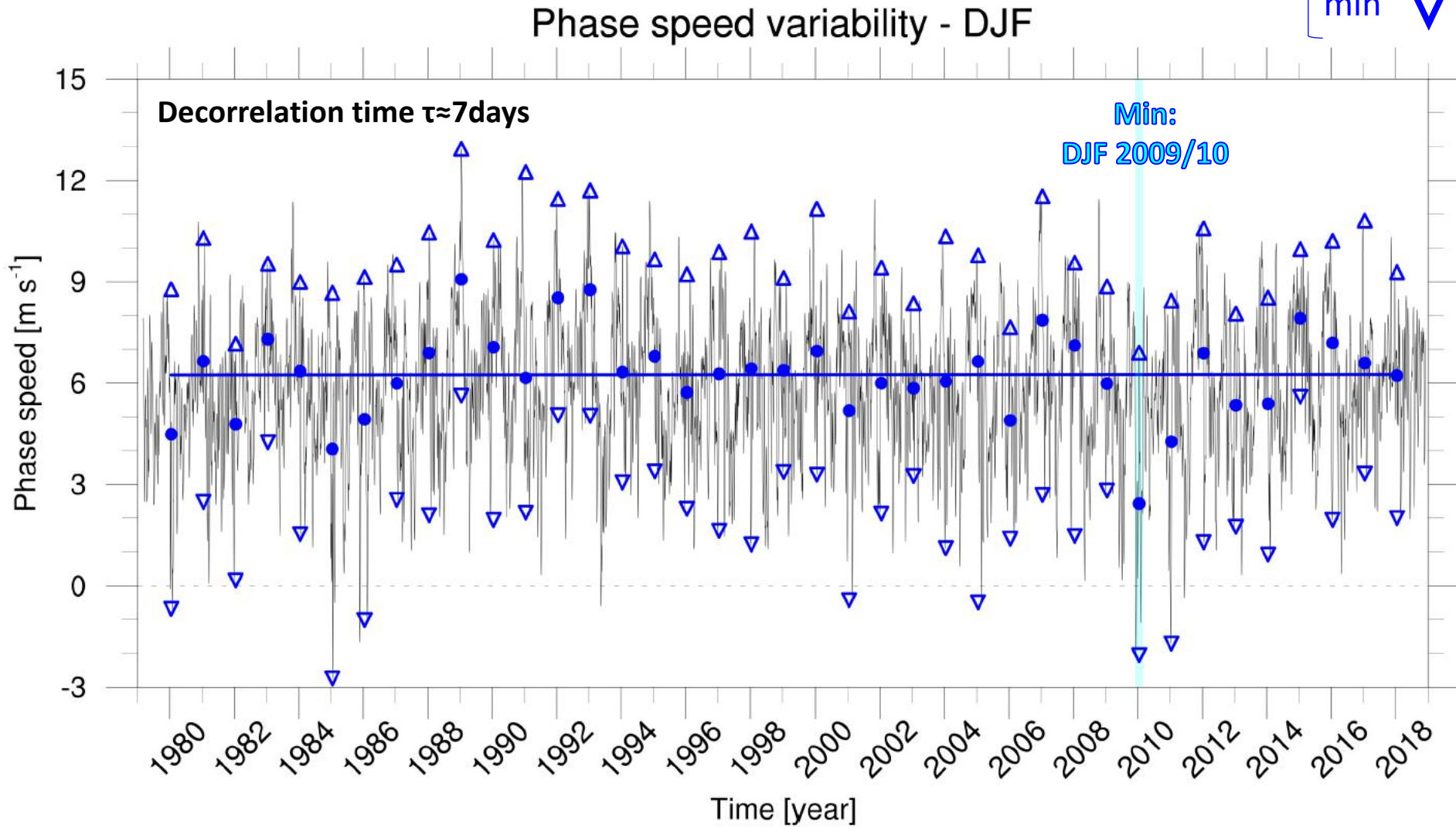
$S(n, c_p)$  spectral coefficients of meridional wind for 61 days period (37 days with tapering), as in Randel and Held (1991).

# Daily phase speed diagnostic

Daily phase speed evolution from Feb. 1979 to Aug. 2018

*Blue line is linear regression (trend is absent,  $p < 0.01$ ).*

Seasonal {  
max  $\triangle$   
mean  $\bullet$   
min  $\nabla$

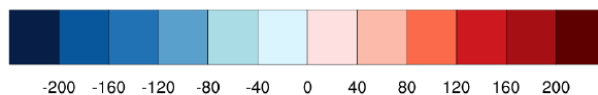
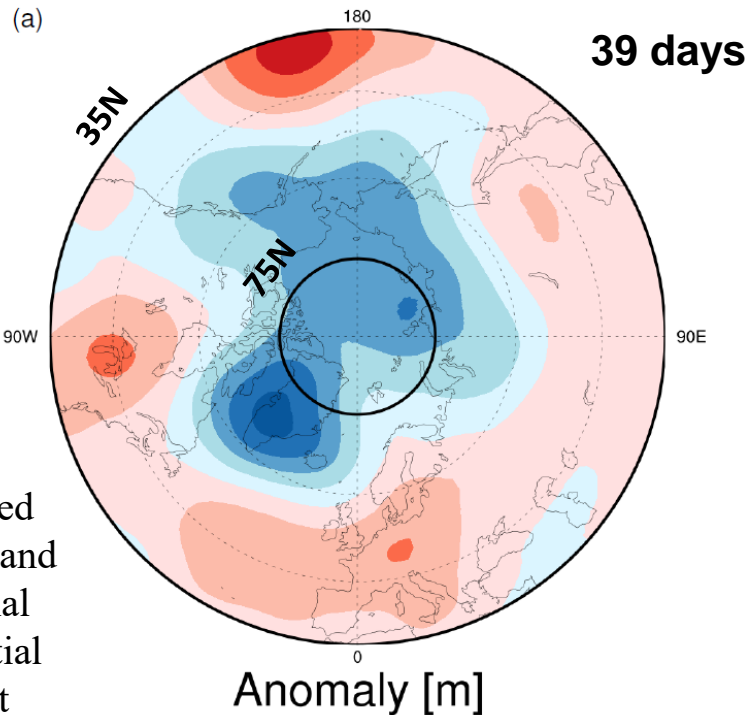


# Large-scale circulation during high/low phase speed days

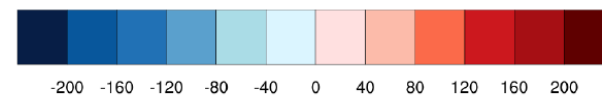
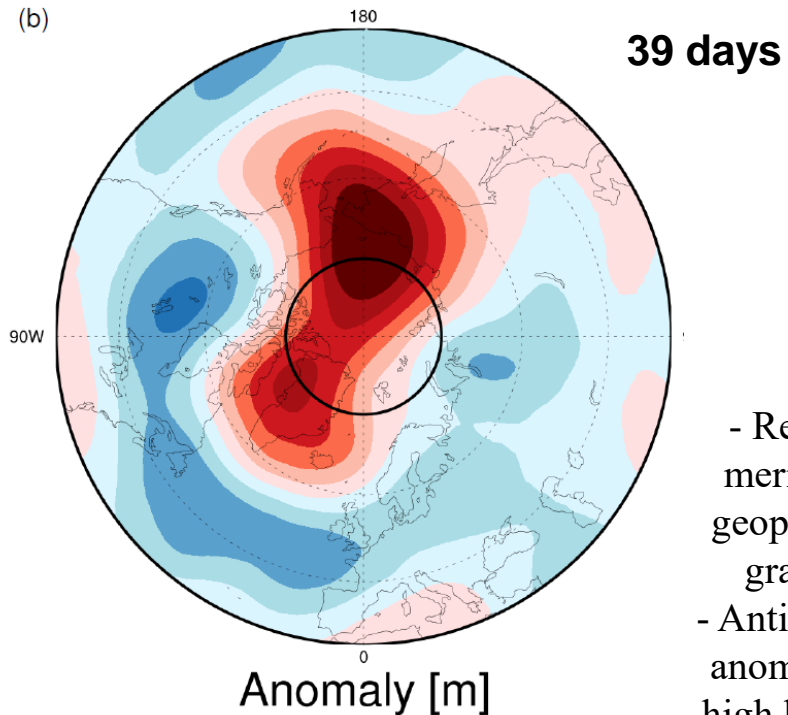
DJF 1979/80→2017/18 (39 winters): composite for the day of maximum and minimum phase speed in each winter.

Low phase speed days correspond to anomalously high blocking activity ([hyperlink](#)) and to extreme temperature events over midlatitudes, especially in winter ([hyperlink](#)).

**High phase speed - DJF**  
250hPa geopt. anomaly



**Low phase speed - DJF**  
250hPa geopt. anomaly



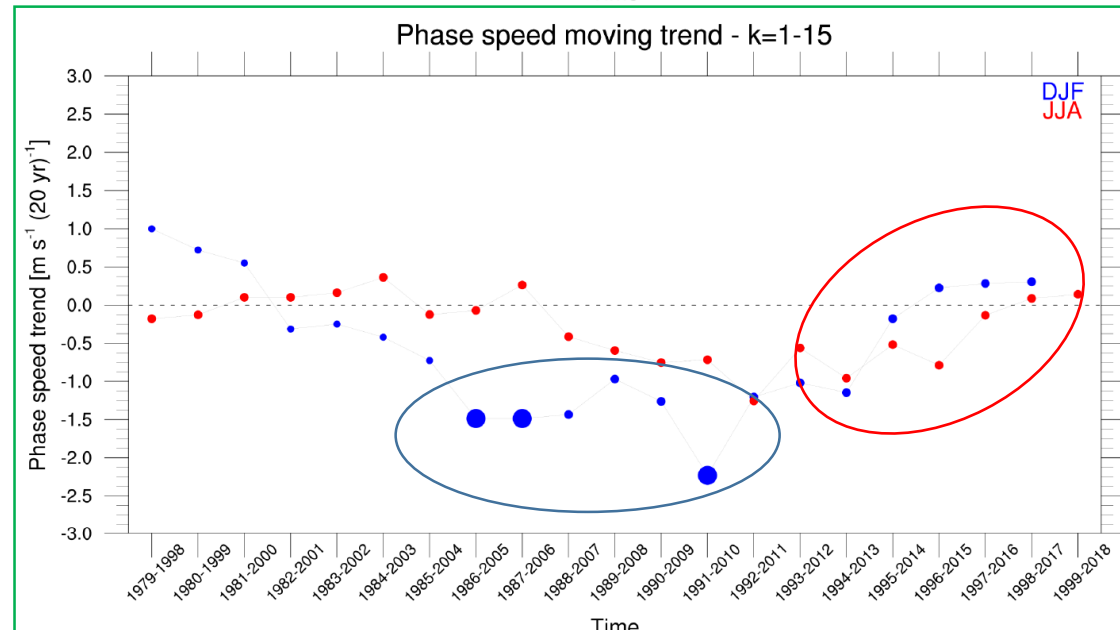
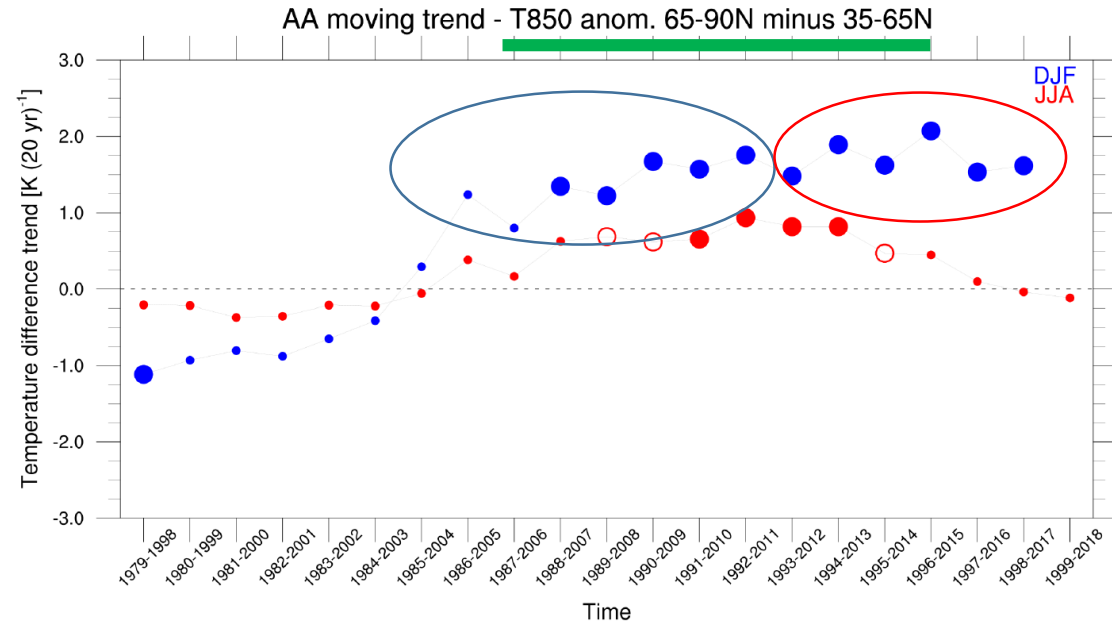
# Arctic Amplification and phase speed trends (1)

p-value > 0.95 ●  
p-value > 0.90 ○

Trends over consecutive  
20-year periods: **T850 anom**  
difference vs phase speed

- Day-to-day Pearson correlation  
 $r(c, \text{T850 anom.}) = -0.39$
- [Arctic-to-midlatitude T850 difference \(hyperlink\)](#)  
consistently increasing since 1988-2007 period (Arctic Amplification).
- No corresponding phase speed trend. Significant negative trend in 1991-2010 period (likely because of extreme 2009/2010 winter).

Riboldi et al. 2020, in review (GRL)

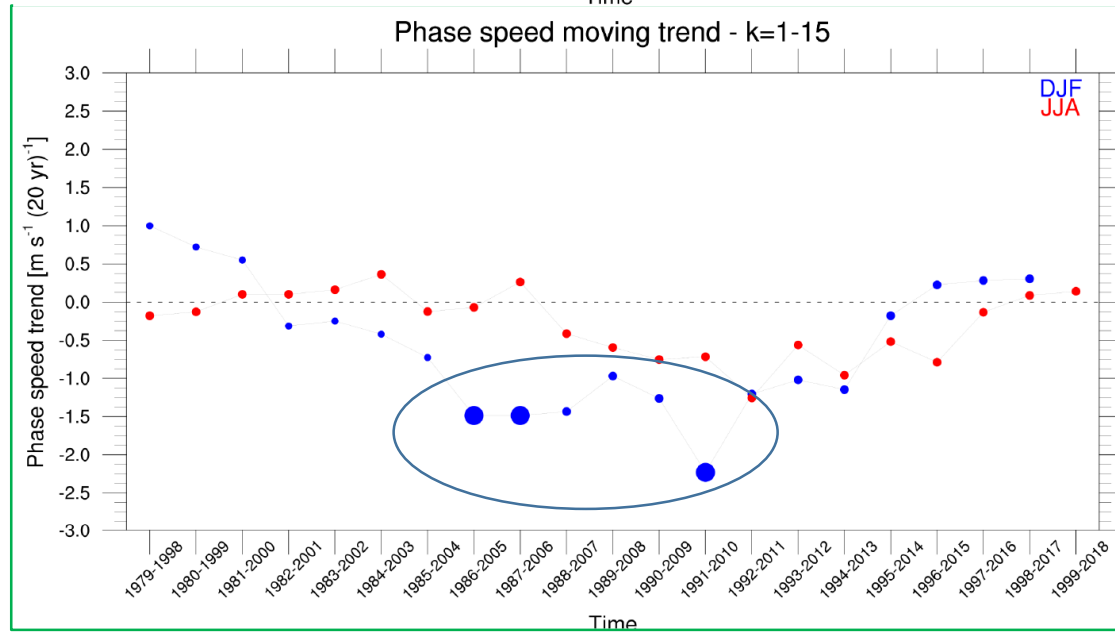
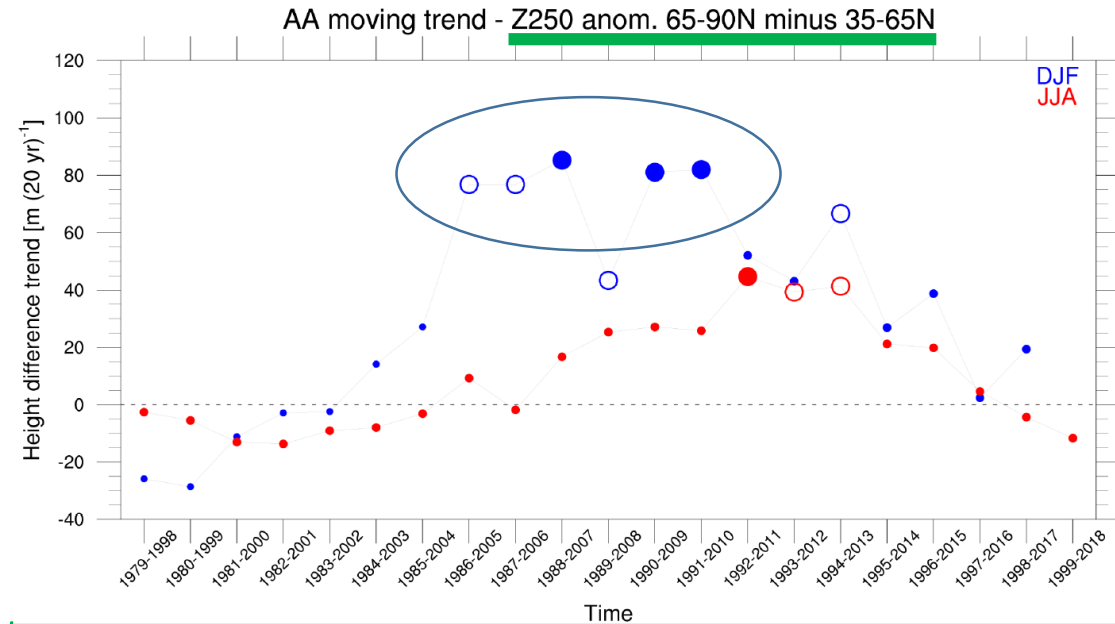


# Arctic Amplification and phase speed trends (2)

p-value > 0.95 ●  
p-value > 0.90 ○

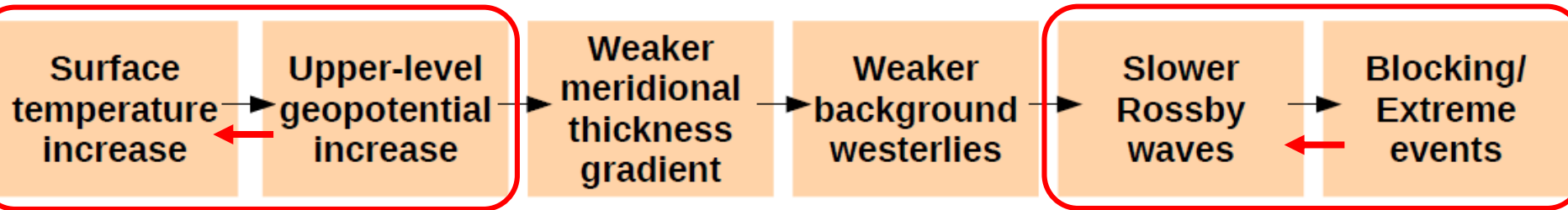
Trends over consecutive  
20-year periods: **Z250 anom**  
difference vs phase speed

- Day-to-day Pearson correlation  
 $r(c, \text{Z250 anom.}) = -0.70$
- [Arctic-to-midlatitude Z250 anomaly \(hyperlink\)](#) trends occur in similar periods as phase speed trends.
- Phase speed trend follows evolution of Z250 gradient, rather than of T850.





# Conclusions



- **Reduction of low-level meridional T gradient not related to reduction in Rossby wave phase speed.**

→ Apparent thermal wind violation, but consistent with eddy-driving of the polar jet.

→ Extratropical variability has affected phase speed more strongly than Arctic Amplification.

- **Conversely, Rossby wave phase speed correlates well with upper-level geopotential gradient and blocking activity.**

→ Upper-level processes modulate Rossby wave phase speed, rather than low-levels.

→ Atmospheric blocking, related to meridional geopotential gradient reversal, can modulate phase speed variability.

**Thanks for your attention!**



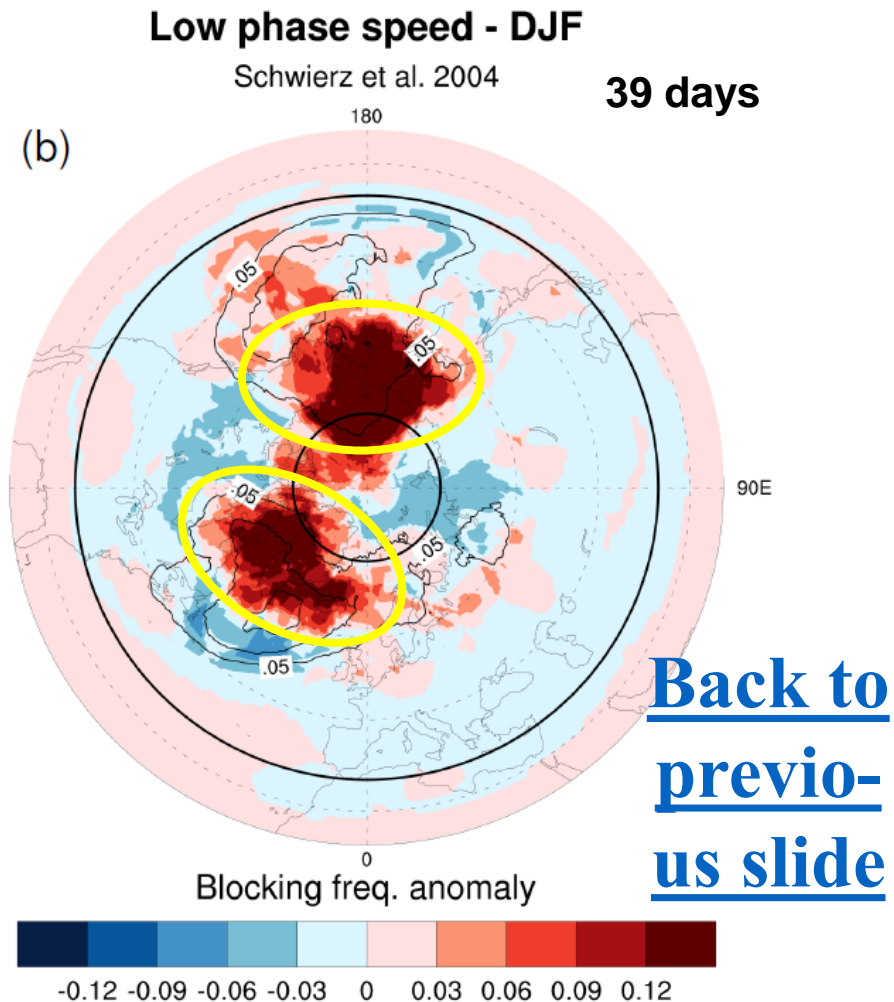
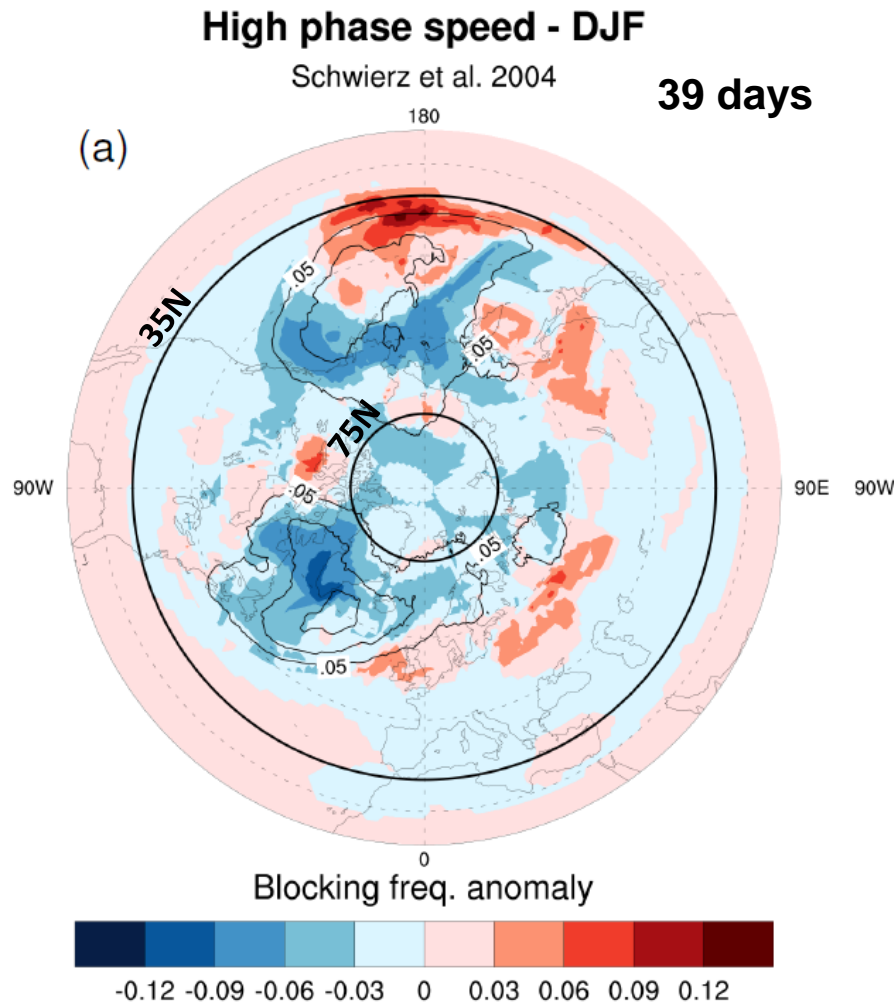
# Bibliography

- Francis, J. A., and Vavrus, S. J. ( 2012), Evidence linking Arctic amplification to extreme weather in mid-latitudes, *Geophys. Res. Lett.*, 39, L06801, doi:[10.1029/2012GL051000](https://doi.org/10.1029/2012GL051000).
- Cohen J., Zhang X., Francis J., et al. ARCTIC CHANGE AND POSSIBLE INFLUENCE ON MID-LATITUDE CLIMATE AND WEATHER: A US CLIVAR White Paper. *US CLIVAR Rep.* (2018) doi:[10.5065/D6TH8KGW](https://doi.org/10.5065/D6TH8KGW)

# LINKED SLIDES

# Atmospheric blocking during high/low phase speed days

Low phase speed occurrence is linked to positive blocking frequency anomalies at the end of the storm tracks (consistent across blocking indices).



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# Extreme temperatures during high/low phase speed days

## Midlatitude EXtreme (MEX) index:

Areally averaged (over N gridpoints between 35°N and 75°N) squared standardized 2-m temperature anomalies (see also Coumou et al. 2014).

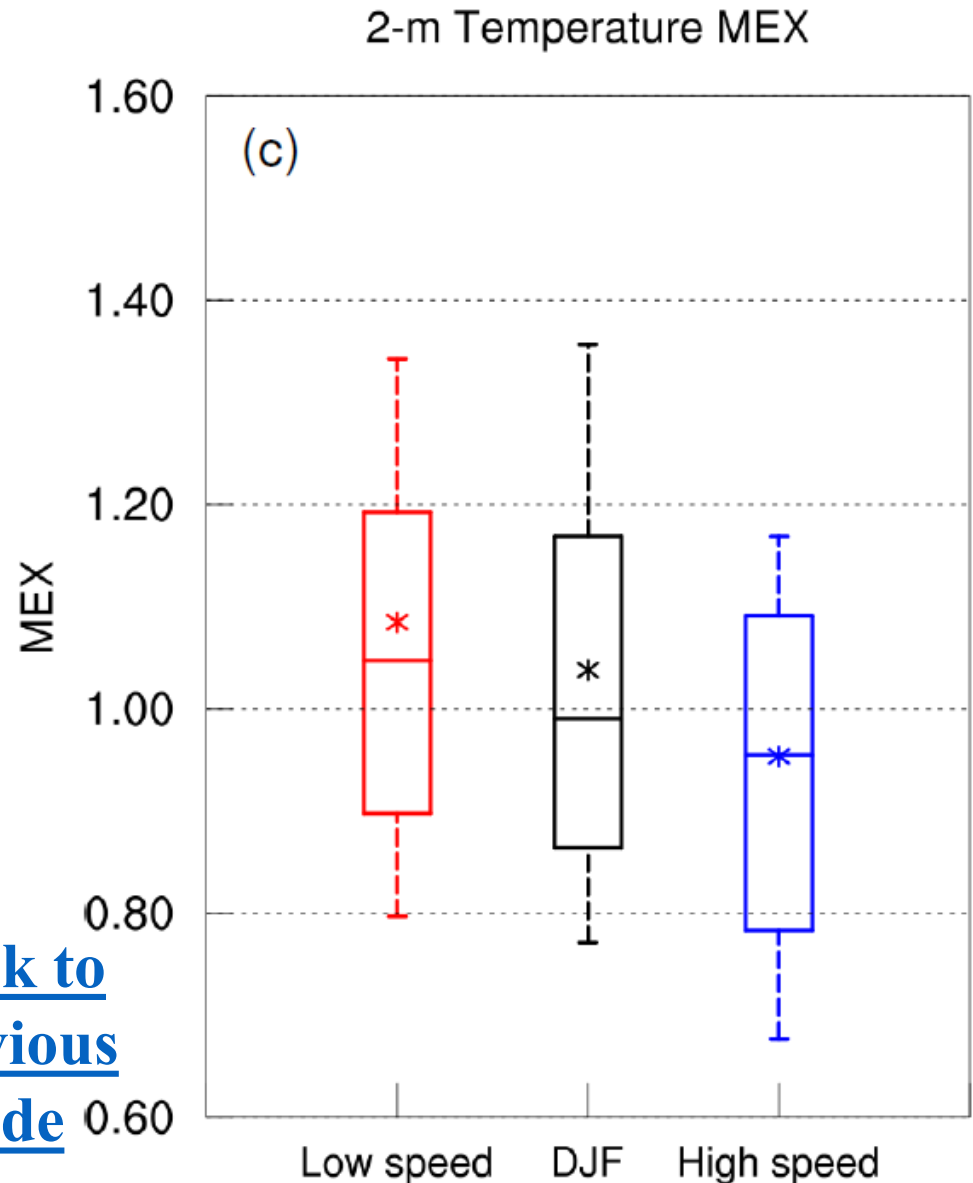
$$MEX(x, t) = \frac{1}{N} \sum_i^N \left( \frac{x_i(t) - \overline{x_i(t)}}{\overline{\sigma}(x_i(t))} \right)^2$$

**Low phase speed → high MEX values → stronger/more extended temperature anomalies than normal.**

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## MEX definition:

Coumou D., Petoukhov V., Rahmstorf S., Petri S., Schellnhuber H. J.: *Quasi-resonant circulation and extreme weather* PNAS, (2014) DOI: 10.1073/pnas.1412797111

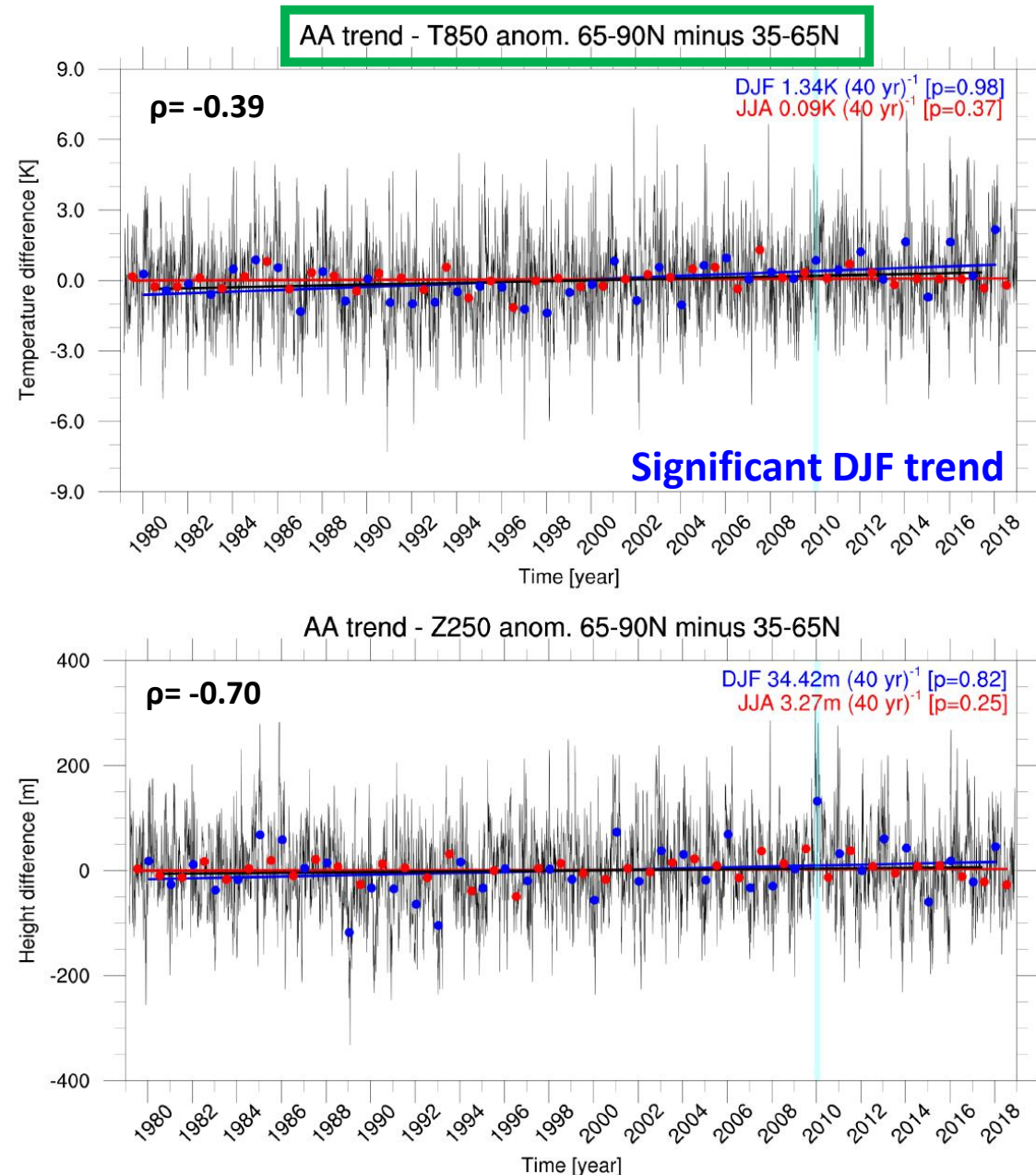


# Arctic Amplification and phase speed trends

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## Two different Arctic Amplification metrics:

- 1) **Thermal metric:**  
850hPa temperature anomaly difference (65°N-90°N minus 35-65°N)
- 2) **Dynamical metric:**  
250hPa geopotential anomaly difference (65°N-90°N minus 35-65°N)



# Arctic Amplification and phase speed trends

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