



POTSDAM INSTITUTE FOR  
CLIMATE IMPACT RESEARCH

**nagra** ● aus verantwortung

# Evolution of the climate in the next Million years:

## A Simple Model for Glacial Cycles and Impact of fossil fuel CO<sub>2</sub>

- Generate a **Process-based Simple Model** to describe the Temporal Natural Evolution of the Climate System as forced by Changes in Incoming Insolation only

Forcing:	Summer Maximum Insolation at 65°N	( <b>f</b> )
Climate System Variables to Model:	Northern Hemisphere (NH) Continental Ice-Volume	( <b>v</b> )
	Atmospheric CO <sub>2</sub> concentration	( <b>CO<sub>2</sub></b> )
	Global Mean Temperature Anomaly	( <b>dT</b> )

- Use the Model to **predict the Evolution of the Climate System in the Next 1 Million years**, including the effect of Anthropogenic CO<sub>2</sub> Emissions

- Process-based Simple Model
- 3 Equations with 3 unknown variables: **v**, **CO<sub>2</sub>**, **dT**
  - **v** is non-dimensional (normalized to vary between [0,1])
  - **CO<sub>2</sub>** is measured in ppm
  - **dT** is measured in C
- For time-scales longer than 1000 years

Equation (1): Temporal Evolution of NH Continental Ice-Volume ( $v$ )

$$\frac{dv}{dt} = \frac{b_1 v + b_2 v^{3/2} + b_3 (f - \bar{f}) + b_4 \log(CO_2)}{1 + b_5 v_{memory}} + b_6$$

$$v_{memory}(t) = \frac{\delta}{T} \int_{t-T}^t v(\tau) d\tau$$

$$\begin{aligned} T &= 30 \text{ kyr} \\ \delta &= 1 \text{ if } \frac{dv}{dt} < 0 \\ \delta &= 0 \text{ otherwise} \end{aligned}$$

Ice-growth / Ice-melting depends on:

- Magnitude of the current Ice-Volume (through linear and a 3/2 power terms)
- Orbital Forcing
- $\log(CO_2)$
- Memory Term: Registering the magnitude of the Ice-Volume in the previous 30 kyr

Note:

- $v$  must be non negative at all times
- Before -400 kyr  $v$  must be larger than 0.05 (in Normalized units)

Equation (2): Temporal Evolution of Atmospheric CO<sub>2</sub> Concentration (CO<sub>2</sub>)

$$CO_2 = c_1 dT + c_2 v + c_3 \min\left(\frac{dv}{dt}, 0\right) + c_4$$

CO<sub>2</sub> evolution depends on:

- dT term: Solubility, Ocean Circulation and Deep Ocean Ventilation effects
- V term: Salinity, vegetation Covered Area and Iron Fertilization effects
- Atlantic Meridional Overturning Circulation changes (term that is only active when melting occurs i.e.  $dv/dt < 0$ )

Equation (3): Temporal Evolution of Global Mean Temperature Anomaly (dT)

$$dT = d_1 v + d_2 \log\left(\frac{CO_2}{278}\right)$$

dT evolution depends on:

- Ice-volume
- $\log(CO_2)$

Model:

$$(1) \quad \frac{dv}{dt} = \frac{b_1 v + b_2 v^{3/2} + b_3 (f - \bar{f}) + b_4 \log(CO_2)}{1 + b_5 v_{memory}} + b_6$$

$$(2) \quad CO_2 = c_1 dT + c_2 v + c_3 \min\left(\frac{dv}{dt}, 0\right) + c_4$$

$$(3) \quad dT = d_1 v + d_2 \log\left(\frac{CO_2}{278}\right)$$

$$v_{memory}(t) = \frac{\delta}{T} \int_{t-T}^t v(\tau) d\tau$$

$$T = 30 \text{ kyr}$$

$$\delta = 1 \text{ if } \frac{dv}{dt} < 0$$

$$\delta = 0 \text{ otherwise}$$

Number of Model Parameters: 12

3 Conditions on the Parameters

Pre-industrial conditions:

if  $v=0$  &  $CO_2 = 278$  ppm  $\rightarrow dT = 0$  C

Last Glacial Maximum conditions:

if  $v=1$  &  $CO_2 = 194$  ppm  $\rightarrow dT = -5$  C

Climate Sensitivity:

if  $CO_2 = 2 \times$  Pre-industrial  $CO_2 \rightarrow dT = [2, 4]$  C

**$\rightarrow$  9 Parameters left to adjust**

Model:

$$(1) \quad \frac{dv}{dt} = \frac{b_1 v + b_2 v^{3/2} + b_3 (f - \bar{f}) + b_4 \log(CO_2)}{1 + b_5 v_{memory}} + b_6$$

$$(2) \quad CO_2 = c_1 dT + c_2 v + c_3 \min\left(\frac{dv}{dt}, 0\right) + c_4$$

$$(3) \quad dT = d_1 v + d_2 \log\left(\frac{CO_2}{278}\right)$$

$$v_{memory}(t) = \frac{\delta}{T} \int_{t-T}^t v(\tau) d\tau$$

$T = 30 \text{ kyr}$   
 $\delta = 1 \text{ if } \frac{dv}{dt} < 0$   
 $\delta = 0 \text{ otherwise}$

Fit of the Model:

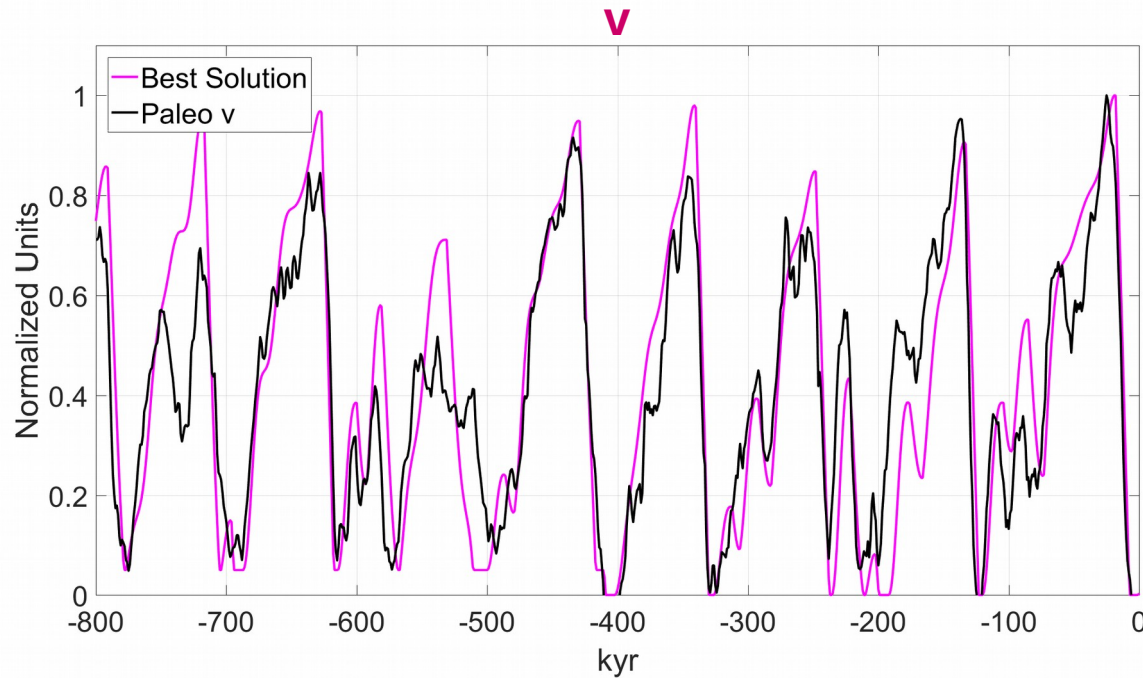
Use of **Paleo Data from the last 800 kyr as Learning Set**

Approach: **Non-Linear Optimisation Problem** with Constraints

1. Optimisation target: Maximise correlation between **v** and **paleo data v**
2. Constraint:  $v(\text{Pre-industrial} \dots +20 \text{ kyr}) < \varepsilon$ , being  $\varepsilon$  small (No Glacial Inception in the next 20kyr permitted)

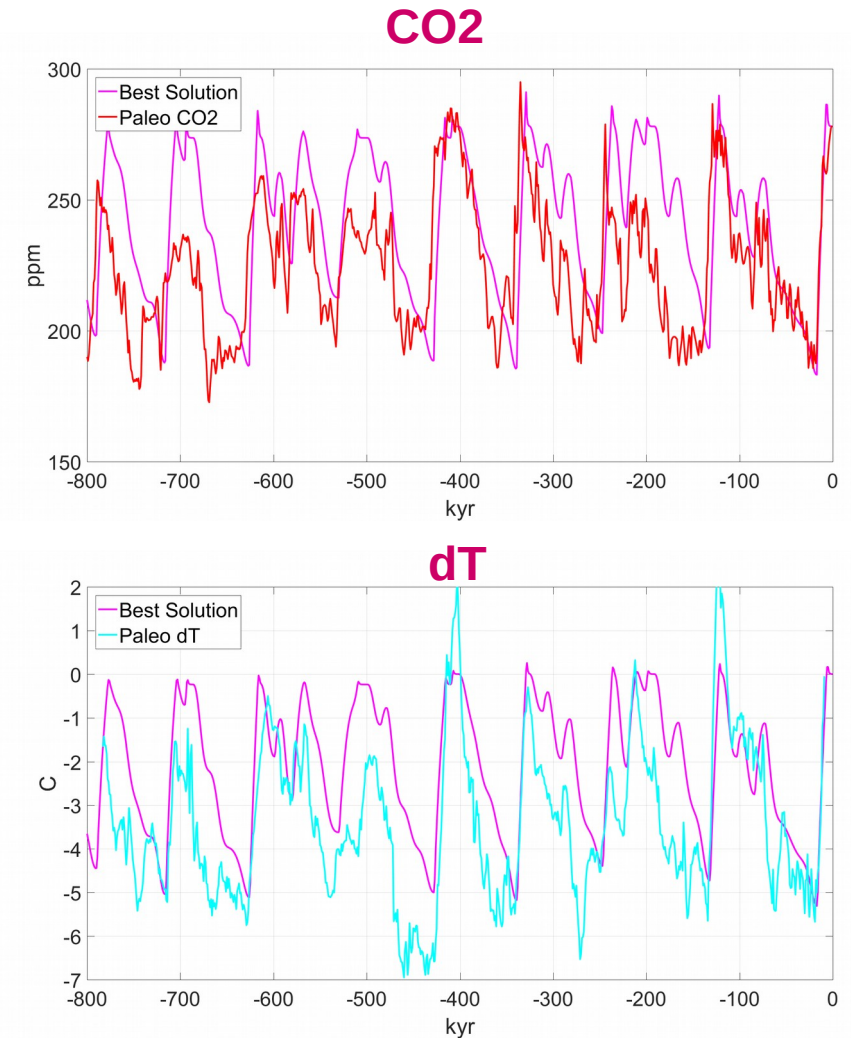


## Best Solution

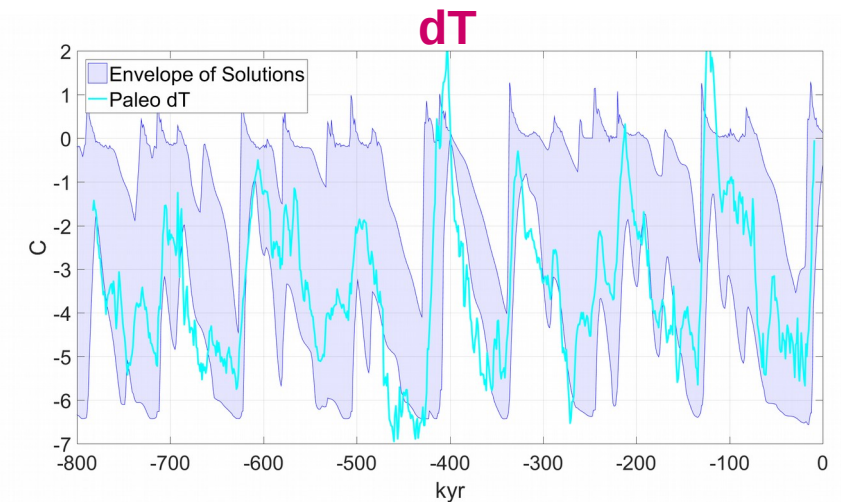
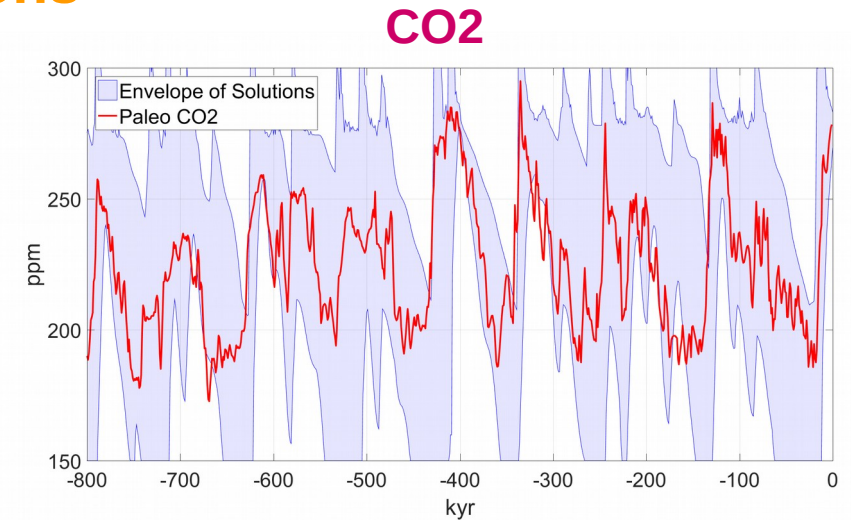
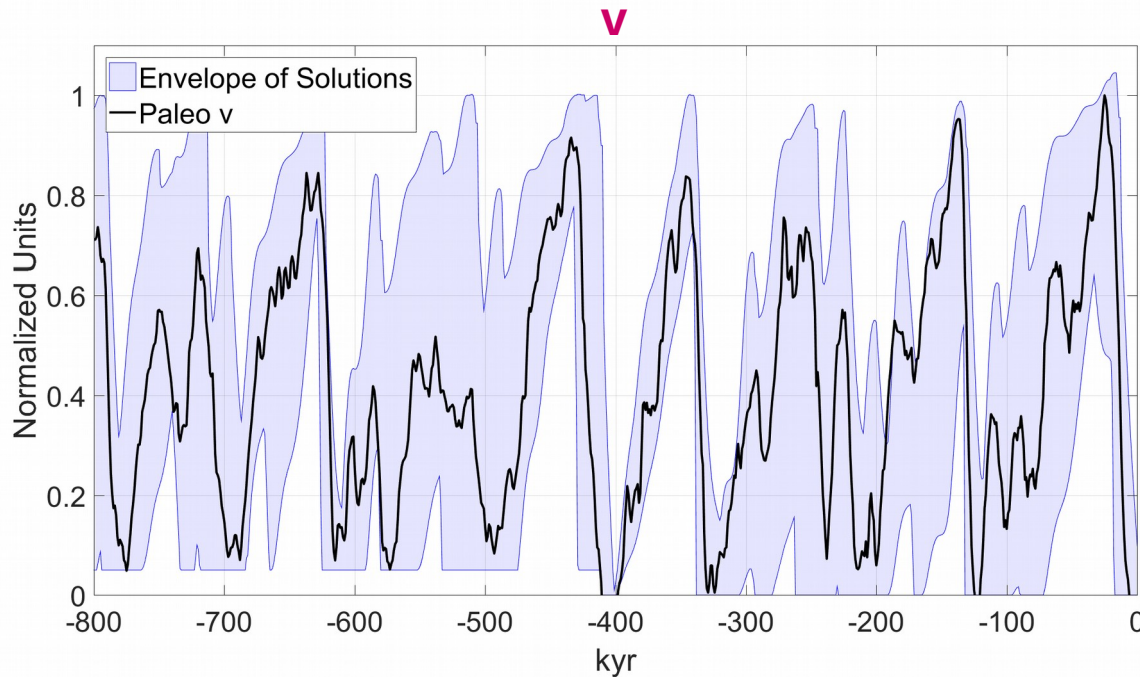


The model is able to reproduce the Natural Temporal Evolution of the Climate System

Correlation between $v$ and Paleo $v$	= 0.87
Correlation between CO <sub>2</sub> and Paleo CO <sub>2</sub>	= 0.60
Correlation between $dT$ and Paleo $dT$	= 0.61



## Ensemble of Accepted Solutions



**Ensemble Accepted Solutions:** Those with correlation between  $v$  and Paleo Data  $v \geq 0.7$  ( **N = 353** )

Mean Correlation between  $v$  and Paleo  $v$  = 0.76

Mean Correlation between CO<sub>2</sub> and Paleo CO<sub>2</sub> = 0.50

Mean Correlation between dT and Paleo dT = 0.56

- The Model has high skill when Validated against the same data used for Learning (Apparent Skill)
- Estimation of Prediction Skill:

Separate Data into independent **Learning** and **Validation** Sets

- Learning: paleo data in period [-800,-400] kyr // Validation: paleo data in period [-400,0] kyr
- Learning: paleo data in period [-400,0] kyr // Validation: paleo data in period [-800,-400] kyr

Loss in Skill from Learning to Validation:

Mean **Correlation between  $v$  and paleo  $v$**  in Validation  
Period = **0.49**

Mean **Correlation between CO<sub>2</sub> and paleo CO<sub>2</sub>** in  
Validation Period = **0.37**

# Model Prediction: Next 1 Milion Years



We add to the Model **Anthropogenic CO2 Emissions**:

$$(1) \quad \frac{dv}{dt} = \frac{b_1 v + b_2 v^{3/2} + b_3 (f - \bar{f}) + b_4 \log(CO_2)}{1 + b_5 v_{memory}} + b_6$$

$$v_{memory}(t) = \frac{\delta}{T} \int_{t-T}^t v(\tau) d\tau$$

$T = 30 \text{ kyr}$   
 $\delta = 1 \text{ if } \frac{dv}{dt} < 0$   
 $\delta = 0 \text{ otherwise}$

$$(2) \quad CO_2 = c_1 dT + c_2 v + c_3 \min\left(\frac{dv}{dt}, 0\right) + c_4 + CO_{2 \text{ Anthropogenic}}$$

$$(3) \quad dT = d_1 v + d_2 \log\left(\frac{CO_2}{278}\right)$$

Anthropogenic Emissions:

- Delta Pulse at Present  $CO_{2 \text{ Anthropogenic}}(t) = E \sum_{i=1}^5 \alpha_i * e^{-t/\tau_i}$   $\alpha_i = \alpha_i(E)$   $\tau_i = \tau_i(E)$   
 Following Lord et al. (2016)  
 $\alpha$  and  $\tau$  polynomials of degree 3

- Magnitude of Cumulative Emission E

E = 500 Pg:

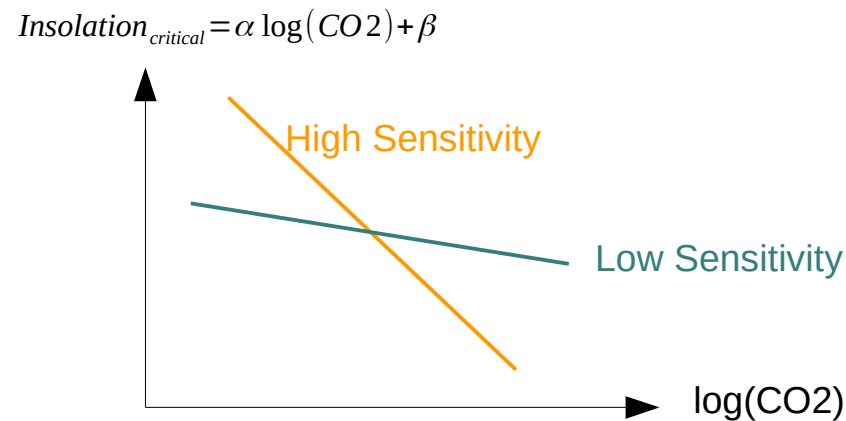
Already emitted

E = 3000 Pg:

Maximum Expected total Emissions in the next 200-300 years



- We show the results of the **Ensemble of Accepted Solutions (N = 353)**
- For each Accepted Solution we calculate the dependency of the Critical Insolation on CO2 levels (as in Ganopolski et al., 2016)

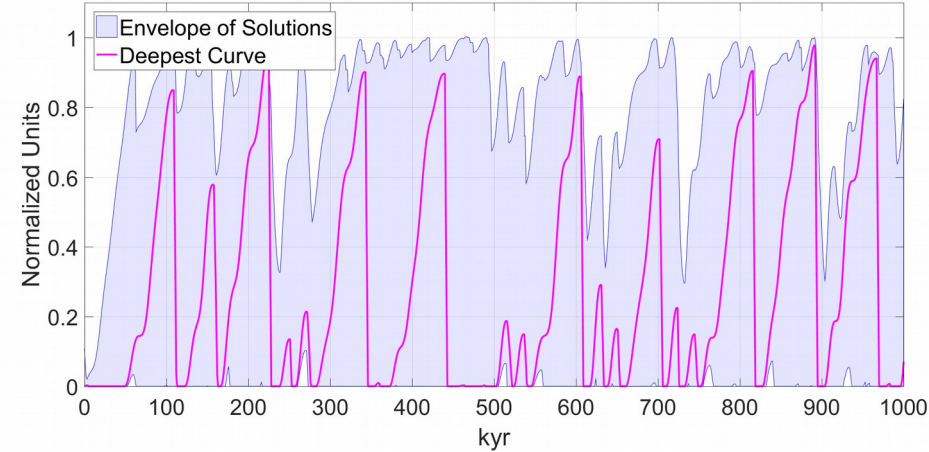


- We separate the solutions into those with **High Sensitivity (N<sub>High</sub> = 324)** or **Low Sensitivity (N<sub>Low</sub> = 29)**

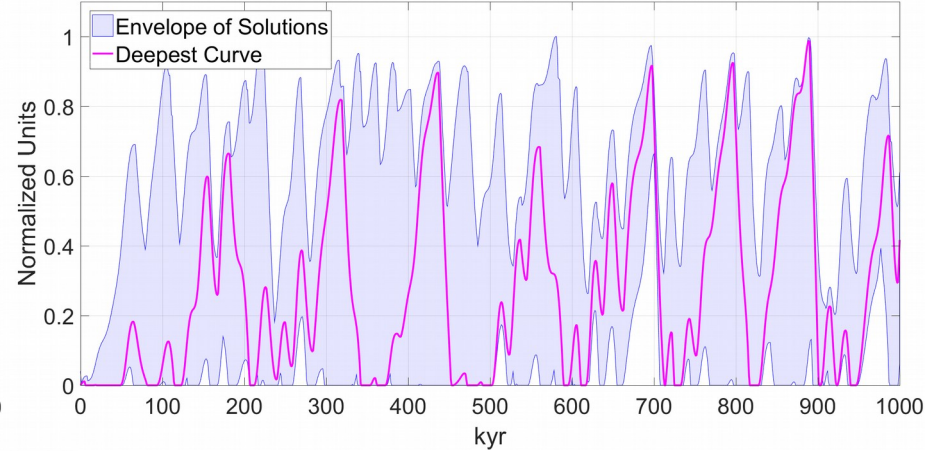


# Model Prediction: Next 1 Milion Years

Ice-Volume, High Sensitivity Solutions,  $E = 0$  Pg



Ice-Volume, Low Sensitivity Solutions,  $E = 0$  Pg



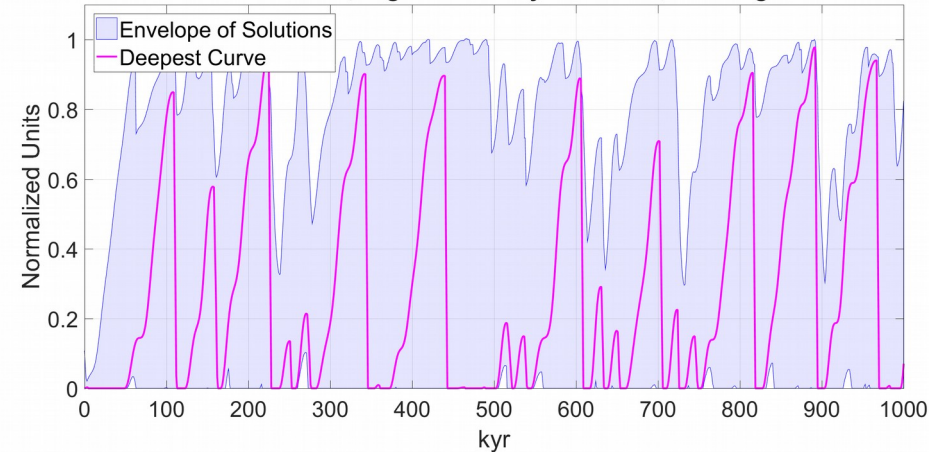
## Natural Scenario ( $E = 0$ )

- High Sensitivity Solutions:
  - Next 1 Myr: Large Uncertainty
  - Next Glacial Inception: Likely in ~50 kyr
- Low Sensitivity Solutions:
  - Next 1 Myr: Large Uncertainty
  - Next Glacial Inception: Likely in ~120 kyr

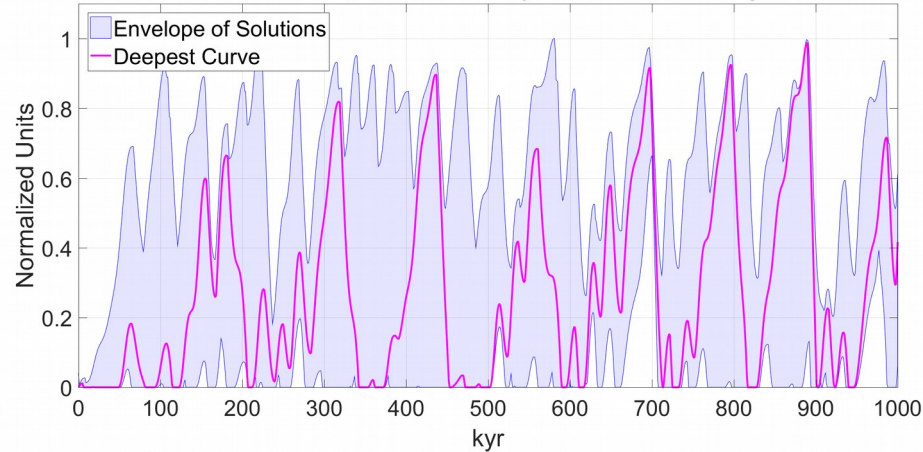
# Model Prediction: Next 1 Milion Years



Ice-Volume, High Sensitivity Solutions, E = 0 Pg



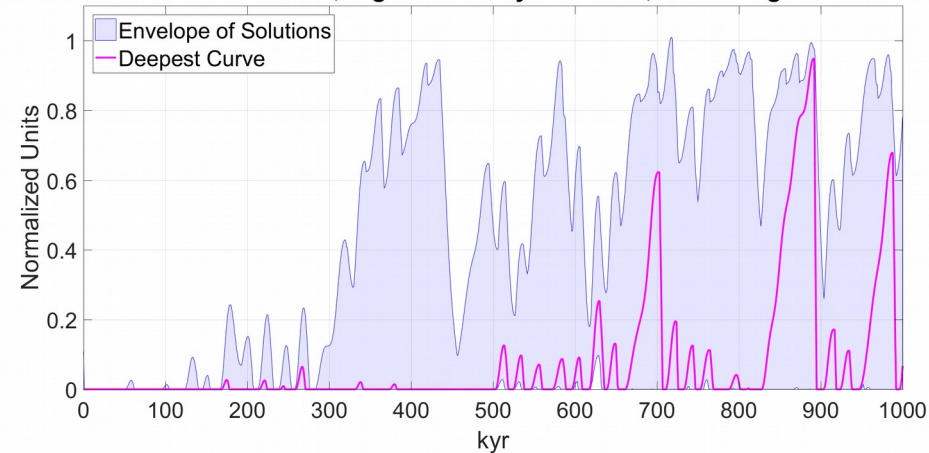
Ice-Volume, Low Sensitivity Solutions, E = 0 Pg



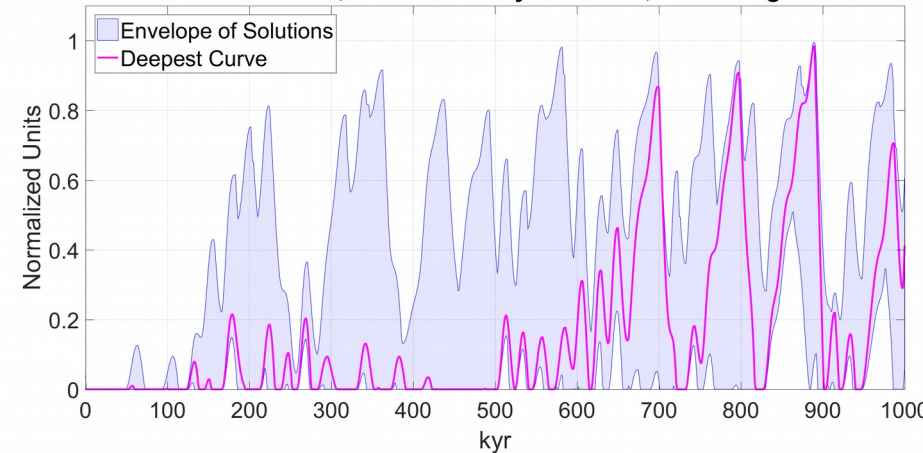
## Natural Scenario (E = 0)

- High Sensitivity Solutions:
  - Next 1 Myr: Large Uncertainty
  - Next Glacial Inception: Likely in ~50 kyr
- Low Sensitivity Solutions:
  - Next 1 Myr: Large Uncertainty
  - Next Glacial Inception: Likely in ~120 kyr

Ice-Volume, High Sensitivity Solutions, E = 500 Pg



Ice-Volume, Low Sensitivity Solutions, E = 500 Pg



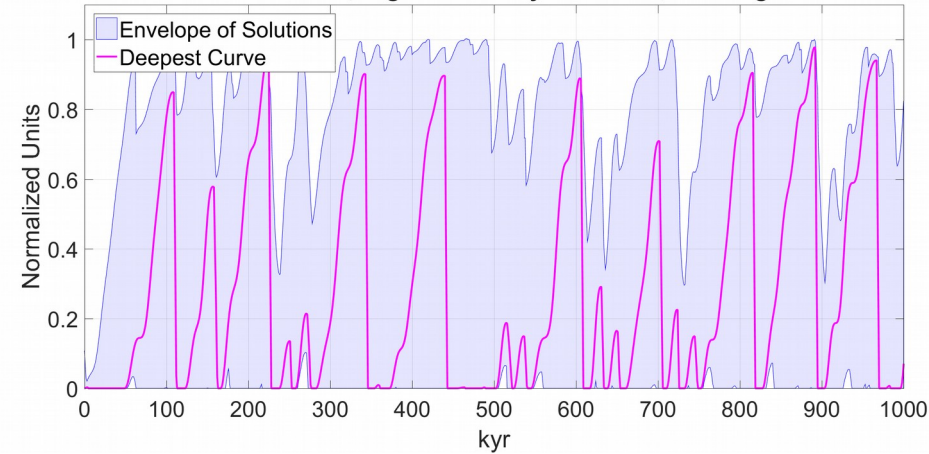
## Emissions Scenario (E = 500 Pg)

- High Sensitivity Solutions:
  - Next 300 kyr: Conditions significantly different from Natural, Low Uncertainty of almost ice-free conditions
- Low Sensitivity Solutions:
  - Next 120 kyr: Conditions significantly different from Natural, Low Uncertainty of almost ice-free conditions

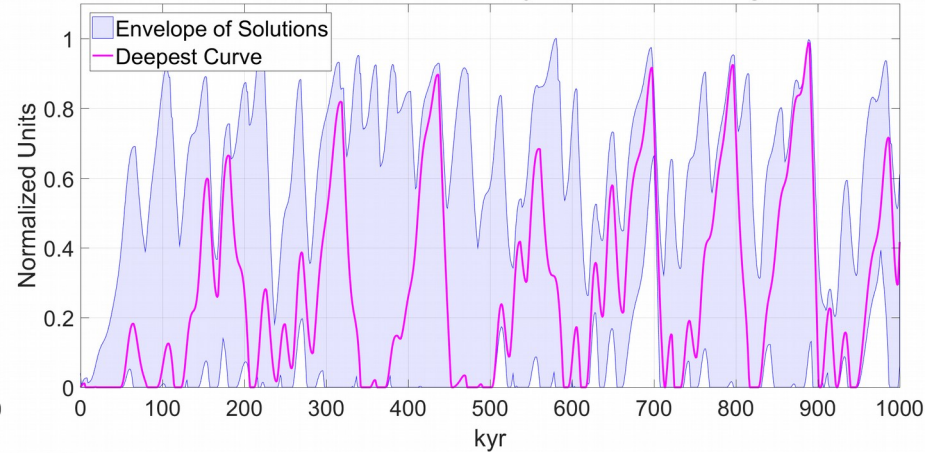
# Model Prediction: Next 1 Milion Years



Ice-Volume, High Sensitivity Solutions, E = 0 Pg



Ice-Volume, Low Sensitivity Solutions, E = 0 Pg



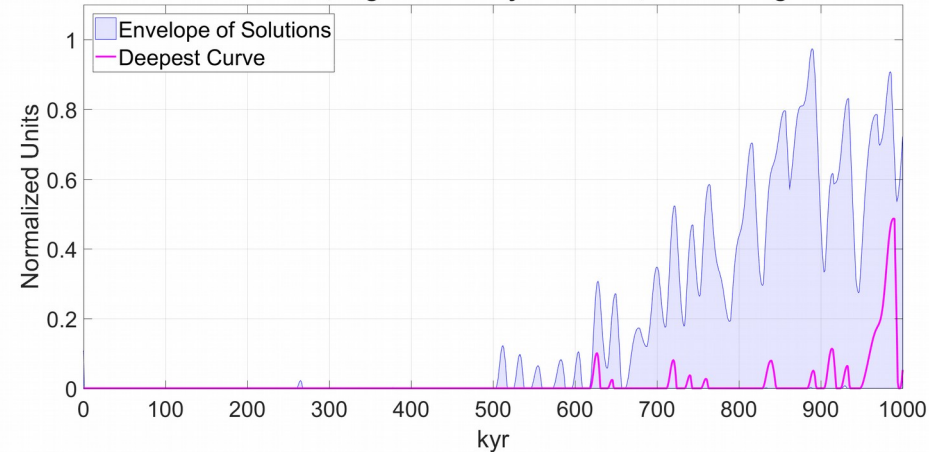
## Natural Scenario (E = 0)

- High Sensitivity Solutions:
  - Next 1 Myr: Large Uncertainty
  - Next Glacial Inception: Likely in ~50 kyr
- Low Sensitivity Solutions:
  - Next 1 Myr: Large Uncertainty
  - Next Glacial Inception: Likely in ~120 kyr

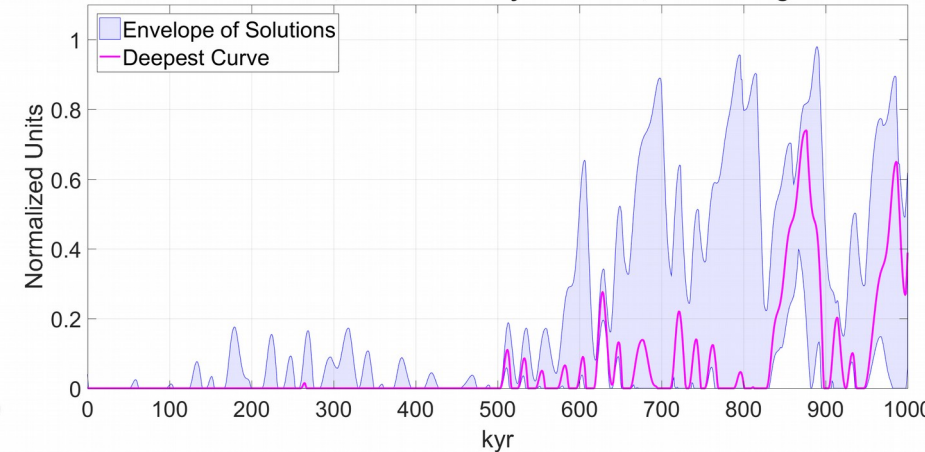
## Emissions Scenario (E = 3000 Pg)

- High Sensitivity Solutions:
  - Next 700 kyr: Conditions significantly different from Natural evolution. Low Uncertainty of almost ice-free conditions
- Low Sensitivity Solutions:
  - Next 550 kyr: Conditions significantly different from Natural evolution. Low Uncertainty of almost ice-free conditions

Ice-Volume, High Sensitivity Solutions, E = 3000 Pg



Ice-Volume, Low Sensitivity Solutions, E = 3000 Pg





- The **Model is successful** in reproducing the Natural Evolution of the Climate System in the **last 800 kyr**
- The **Model is skillful** in **Predictive Mode** (when evaluated with data in an independent Validation Set)
- For the Future:
  - **Natural Scenario:**
    - **Large Uncertainty** in the Next 1 Myr: Indication that the past does not constraint the future of the Climate Evolution subject only to Orbital Forcing
    - **Next Glacial Inception:** Most likely to occur in **~50 kyr** from now
  - **Anthropogenic Emissions Scenarios:**
    - Even **already achieved CO<sub>2</sub> Anthropogenic Emissions** (500 Pg) are capable of affecting the Climate Evolution for long periods: **Significantly different from Natural behaviour in, at least, the next 120 kyr**
    - **High CO<sub>2</sub> Anthropogenic Emissions** (3000 Pg or larger): **Significantly different from Natural behaviour in, at least, the next 550 kyr**

# Thanks!

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- Ganopolski, A., Winkelmann, R., & Schellnhuber, H. J. (2016). Critical insolation–CO<sub>2</sub> relation for diagnosing past and future glacial inception. *Nature*, 529(7585), 200-203.
- López-Pintado, S., & Romo, J. (2006). Depth-based classification for functional data. *DIMACS Series in Discrete Mathematics and Theoretical Computer Science*, 72, 103.
- Lord, N. S., Ridgwell, A., Thorne, M. C., & Lunt, D. J. (2016). An impulse response function for the “long tail” of excess atmospheric CO<sub>2</sub> in an Earth system model. *Global Biogeochemical Cycles*, 30(1), 2-17.

## Paleo Data of the last 800 kyr:

- Friedrich, T., Timmermann, A., Tigchelaar, M., Timm, O. E., & Ganopolski, A. (2016). Nonlinear climate sensitivity and its implications for future greenhouse warming. *Science Advances*, 2(11), e1501923.
- Petit, J. R., Jouzel, J., Raynaud, D., Barkov, N. I., Barnola, J. M., Basile, I., ... & Delmotte, M. (1999). Climate and atmospheric history of the past 420,000 years from the Vostok ice core, Antarctica. *Nature*, 399(6735), 429.
- Spratt, R. M., & Lisiecki, L. E. (2016). A Late Pleistocene sea level stack. *Climate of the Past*, 12(4), 1079-1092.