The Effects of Cracks and Fluids on Post-Seismic healing

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Motivation



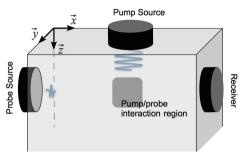
We observe changes in seismic velocity after large earthquakes. What parameters control these effects?

- "Comprehensive observation and modeling of earthquake and temperature-related seismic velocity changes in northern Chile with passive image interferometry." Richter et al (2013)
- "The mechanism by which seismic velocities decrease in response to stress perturbations is commonly described as related to the opening of cracks (9, 10)" Brenguier et al. (2014).
- "The largest coseismic drops are observed close to the fault zones." Hobiger et al, 2016

We design a lab experiment to look at these effects in more detail.

Basic Experiment





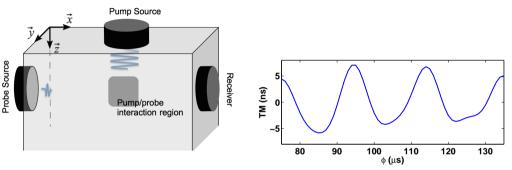
Two waves:

- PUMP (Proxy for earthquake): $\epsilon \sim 10^{-6}$ perturbs rock
 - $\lambda pprox$ 40 mm
- probe (Proxy for noise): $\epsilon \sim 10^{-8}$ senses perturbation $\lambda pprox$ 6 mm

 ϵ – strain

Basic Experiment





- Strong PUMP wave slows weak probe wave
- Directly sense the PUMP with the probe
- Similar to Dynamic Acousto-Elasticity Testing (DAET, Renaud et al., 2012), but 2D rather than 1D
- More detail in Gallot et al., 2015, TenCate et al, 2016

Making Measurements



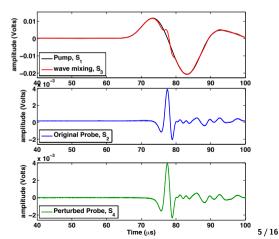
We measure a time delay in the probe as a function of the phase of the pump, which we control by controlling the timing between the pump and the probe. These two slides show two different values of ϕ .

For transmission delay ϕ , we record:

- probe S1
- 2 PUMP S₂
- OUMP+probe S3

Compute:

-) perturbed probe: $S_4 = S_3 S_2$
- 2 time delay:
 - $\blacktriangleright S_4 * S_1$
 - interpolate peak
 - time delay(ϕ) = peak time



Making Measurements



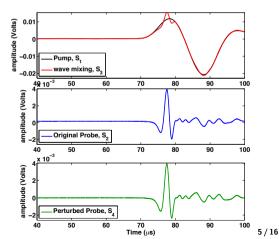
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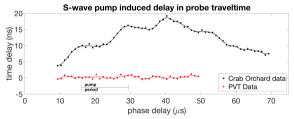
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Data

First, we verify that we see an effect, and that the effect is in the rock and not the apparatus, by comparing the data in a linear material (PVT, a type of plastic) and a rock (Crab Orchard Sandstone). The x-axis is the ϕ from above, which is the time delay between the release of the PUMP wave and the release of the probe wave. Changing ϕ changes the part of the PUMP waveform that the probe senses.



Conclusion:

• The nonlinear effect is in the rock, not the apparatus

Observations:

- We see two frequencies in our data:
 - One is at the PUMP frequency (74 kHz)
 - One is at a much lower frequency (controlled by the envelope of the pump)



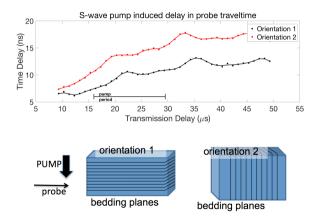


We have looked at how several rock and experimental parameters change our experimental results

- Crack orientation (experimental and modelled)
- Applied load (experimental and modelled)
- Humidity

Effect of Crack Orientation





Observations:

• The part of the signal at the frequency of the PUMP envelope is affected by crack orientation

Questions:

- Why is the high-f signal not affected?
- Can we explain these results with changes only in cracks?
- Are there field datasets where we might see this kind of orientation difference in velocity perturbations?

Modeling Velocity Change



We model the experiment by calculating the change in probe wavespeed, V_p as a function of the strain in the PUMP, ϵ_{ij} . We model the PUMP propagation with the standard five-constant nonlinearity model and the cracks with linear slip theory.

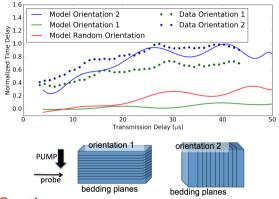
$$V_{\rho} = \sqrt{\frac{1}{\rho}} \left(\lambda + 2\mu + 2\left[(B+C)(\epsilon_{11} + \epsilon_{22}) + (A+3B+C)\epsilon_{33}\right]\right)$$

- V_p : P-wave speed
 - ρ : mass density
- λ, μ : Lamé parameters
- A, B, C: 3rd-order moduli
 - ϵ : strain tensor

More details: Rusmanugroho et al, 2020

Modeling Velocity Change





Questions:

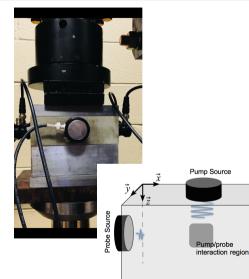
- Why is the high-f part of the signal independent of orientation?
- What do we stand to gain by using a more realistic nonlinearity model?

Observations:

- The basic structure is well-captured, but the details are not. This is perhaps because our rock is more complicated and our source/receive model simpler than reality.
- Our model shows a much larger change due to orientation than do our data. This indicates that it is likely there are more than one set of fractures in the sample.
- In both the data and the model, the signal at the PUMP frequency is independent of orientation.

Applying Uniaxial Load



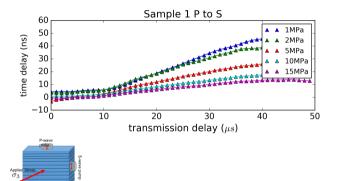


- We apply a range of values of uniaxial stress, with the load held steady at each stress.
- Up to pprox 18 MPa
- 4 experiments
 - 2 samples, with different fracture orientations
 - P and S probes

More details, Hayes et al, 2018

Applying Pressure





Observations:

- Our signal shrinks as we increase the load.
- This is independent of PUMP/probe/fracture orientation (other data not shown).

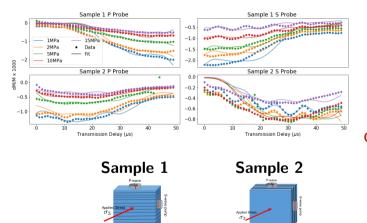
Questions:

 The signal decreases even with a load where we would expect the fractures to open (applied stress ⊥ to fracture normals). Why are they still closing? Or are we simply not sensing them open?

Evaluating data fit



Fit data to Sens-Schönfelder's (2019) model.



Observations:

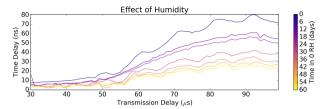
- Capture trend on all data
- Capture details on some data
- Conspicuously missing $|\epsilon|$ dependence (no signal above PUMP frequency, $|\epsilon|$ should be at twice the PUMP freq)

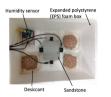
Questions:

• We see only some of the mechanisms seen in DAET. Why?

The effect of humidity







Questions:

 What is the pore-scale mechanism? Are we breaking and re-attaching water bridges across pores?

Observations:

- We see a significant drop in the nonlinear signal as a function of humidity.
- The amount of water in the sample is very small, but enough to potentially cover all of the pore surface with a single water molecule.
- The effect is once-again primarily in the signal at the PUMP envelope frequency not at the frequency of the PUMP itself.

Conclusions



- Cracks completely dominate the nonlinear effect
- Low-f signals decrease under applied load ⇒ closing fractures? Easier crossing of connections?
- Higher-f signals show no consistent trend
- Simple models can capture the trends, even without detailed experimental matching.
- Humidity has a strong effect, even though only very small amounts of water move around, but again only on the low-frequency part of our signal.

References



References in red are our papers, that pertain directly to these experiments, other references give motivation and background information.

- Benson, Philip M., Philip G. Meredith, and Ellen S. Platzman. "Relating pore fabric geometry to acoustic and permeability anisotropy in Crab Orchard Sandstone: A laboratory study using magnetic ferrofluid." Geophysical research letters 30.19 (2003).
- Brenguier, F., et al. "Mapping pressurized volcanic fluids from induced crustal seismic velocity drops." Science 345.6192 (2014): 80-82.
- Gallot, Thomas, et al. "Characterizing the nonlinear interaction of S-and P-waves in a rock sample." Journal of applied physics 117.3 (2015): 034902.
- Hayes, L., Malcolm, A., Moravej, K., Butt, S., "Nonlinear interactions of P and S waves under uniaxial stress." Proceedings of Meetings on Acoustics 21ISNA. Vol. 34. No. 1. Acoustical Society of America, 2018.
- Hobiger, Manuel, et al. "Coseismic and post-seismic velocity changes detected by passive image interferometry: comparison of one great and five strong earthquakes in Japan." Geophysical Journal International 205.2 (2016): 1053-1073.
- Renaud, G., P-Y. Le Bas, and P. A. Johnson. "Revealing highly complex elastic nonlinear (anelastic) behavior of Earth materials applying a new probe: Dynamic acoustoelastic testing." Journal of Geophysical Research: Solid Earth 117.B6 (2012).
- Richter, Tom, et al. "Comprehensive observation and modeling of earthquake and temperature-related seismic velocity changes in northern Chile with passive image interferometry." Journal of Geophysical Research: Solid Earth 119.6 (2014): 4747-4765.
- Rusmanugroho, Herurisa, Alison Malcolm, and Meghdad Darijani. "A numerical model for the nonlinear interaction of elastic waves with cracks." Wave Motion 92 (2020): 102444.
- Sens-Schönfelder, Christoph, Roel Snieder, and Xun Li. "A model for nonlinear elasticity in rocks based on friction of internal interfaces and contact aging." Geophysical Journal International 216.1 (2019): 319-331.
- TenCate, James A., et al. "The effect of crack orientation on the nonlinear interaction of a P wave with an S wave." Geophysical Research Letters 43.12 (2016): 6146-6152.