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## A Next Generation Ocean Carbon Isotope Model for Climate Studies

Rolf Sonnerup and **Mariona Claret**

university of washington, Joint Institute for Study of the Atmosphere and Ocean, United States of America (rolf@uw.edu)

The  $^{13}\text{C}/^{12}\text{C}$  of dissolved inorganic carbon ( $\delta^{13}\text{C}$  DIC) carries valuable information on ocean biological C-cycling, air-sea  $\text{CO}_2$  exchange, and circulation. Paleo-reconstructions of oceanic  $^{13}\text{C}$  from sediment cores provide key insights into past as changes in these three drivers. As a step toward full inclusion of  $^{13}\text{C}$  in the next generation of Earth system models, we implemented  $^{13}\text{C}$ -cycling in a  $1^\circ$  lateral resolution ocean-ice-biogeochemistry Geophysical Fluid Dynamics Laboratory (GFDL) model driven by Common Ocean Reference Experiment perpetual year forcing. The model improved the mean of modern  $\delta^{13}\text{C}$  DIC over coarser resolution GFDL-model implementations, capturing the Southern Ocean decline in surface  $\delta^{13}\text{C}$  DIC that propagates to the deep sea via deep water formation. The model is used here to quantify controls on modern and anthropogenic  $\delta^{13}\text{C}$  DIC as well as to test their sensitivity to wind speed/gas exchange parameterizations.

We found that reducing the coefficient for air-sea gas exchange following OMIP-CMIP6 protocols reduces deep sea modern  $\delta^{13}\text{C}$  DIC by 0.2 permil and improves the depth-integrated anthropogenic  $\delta^{13}\text{C}$  DIC relative to previous gas exchange parameterizations. This is because the  $\delta^{13}\text{C}$  DIC of the endmembers ventilating the deep sea and intermediate waters are highly sensitive to the wind speed dependence of the air-sea  $\text{CO}_2$  gas exchange. Additionally, meridional gradients of surface modern  $\delta^{13}\text{C}$  DIC are better resolved with OMIP-CMIP6. While this model was initially constructed to study the anthropogenic  $^{13}\text{C}$  response, it has promising applications toward longer time scales. For example, BLING 13 C includes controls on the biological C-pump thought to be important in the glacial ocean: light and iron limitation, and controls on  $^{13}\text{C}$  of organic matter formation, and thus on ocean  $\delta^{13}\text{C}$  DIC and its vertical gradient, that depend on  $\text{pCO}_2$ .