

# Io' Active Eruption Plumes: Insights from HST

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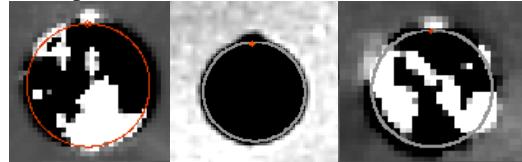
## 1. Introduction

Io is the most volcanically active body in the Solar System. Volcanism on Io has been studied extensively over the last 30 years thru imaging and spectroscopy using instrumentation aboard the earth orbiting Hubble Space Telescope (HST), at ground based observatories and using the payloads of spacecraft travelling in and thru the jovian system such as: Voyager, Galileo, Cassini and most recently the New Horizons (NH) spacecraft. From the first Voyager observations of Io it was established that all of the  $\text{SO}_2$  gas on Io ultimately originates from Io's active volcanism [1]. Although Io's volcanism styles are varied, Io's gaseous plumes have been segregated solely into two classes, namely: Pele- and Prometheus-type plumes [2, 3, 4]. Pele-type plumes are rich in diatomic sulfur ( $\text{S}_2$ ) gas [5] and produce large pyroclastic deposits of average maximum radius  $\sim 600$  km, that are dominantly red in color, due to the probable presence of  $\text{S}_3$  and  $\text{S}_4$  [3,6]. Prometheus type plumes are produced from the interaction of hot silicate lava with  $\text{SO}_2$  surface frost [6, 7, 8]; these plumes form small ( $\leq 350$  km) dominantly white pyroclastic deposits [3,6], are rich in  $\text{SO}_2$  gas and do not have a significant abundance of  $\text{S}_2$  gas [9]. Although the dimensions of the two plume types are well established, the relative significance of the two plume types in maintaining Io's tenuous atmosphere or in maintaining Io's resurfacing rate is not. Consequently, a detailed study of the composition and density of the gas and dust density species within each of Io's active plume types is needed.

## 2. Data Overview

Taking advantage of the available data, we recently [10] completed a detailed analysis of the spectral signature of Io's Pele-type Tvashtar plume as imaged by the HST Wide Field and Planetary Camera 2 (HST/WFPC2) via absorption during Jupiter transit and via reflected sunlight in 2007, as well as HST/WFPC2 observations of the 1997 eruption of

Io's Prometheus-type Pillan plume (Fig. 1). These observations were obtained in the 0.24-0.42  $\mu\text{m}$  range, where the plumes gas absorption and aerosol scattering properties are most conspicuous. By completing a detailed analysis of these observations, several key aspects of the reflectance and the absorption properties of the two plumes have been revealed. Additionally, by considering the analysis of the HST imaging data in light of previously published spectral analysis of Io's Prometheus and Pele-type plumes several trends in the plume properties have been determined, allowing us to define the relative significance of each plume on the rate of re-surfacing occurring on Io and providing the measurements needed to better assess the role the volcanoes play in the stability of Io's tenuous atmosphere.



**Fig. 1:** Io's 1997 Pillan (left) eruption plume at 0.33  $\mu\text{m}$ ; Io's Tvashtar eruption plume in reflected sunlight at 0.33  $\mu\text{m}$  (right) and during Jupiter transit (middle) at 0.26  $\mu\text{m}$ , plume at top of disk.

## 3. Summary and Conclusions

Previously published spectral observations indicate that the  $\text{SO}_2$  gas density in the Pele plume can range from  $\sim 0.5\text{-}5 \times 10^{16} \text{ cm}^{-2}$ , the resulting  $\text{SO}_2$  resurfacing rate can range from  $\sim 0.12\text{-}0.57 \text{ cm yr}^{-1}$ , and the  $\text{S}_2/\text{SO}_2$  gas density ratio is  $\sim 0.05\text{-}0.3$  [5,11]. Presuming that the  $\text{S}_2/\text{SO}_2$  for both the Pele and Tvashtar plumes is  $\sim 0.05\text{-}0.3$ , analysis of the HST Tvashtar plume transit data indicates that the plume  $\text{SO}_2$  and  $\text{S}_2$  gas densities are  $\sim 2\text{-}8 \times 10^{16} \text{ cm}^{-2}$  and  $3\text{-}6 \times 10^{15} \text{ cm}^{-2}$ , respectively, potentially producing  $\text{SO}_2$  and  $\text{S}_2$  gas resurfacing rates  $\sim 0.3\text{-}1.3 \text{ cm yr}^{-1}$  and  $0.07\text{-}0.12 \text{ cm yr}^{-1}$ , respectively. Thus, we find that the gas density in the Tvashtar plume overlaps that derived from the Pele observations. Additionally, the Tvashtar  $\text{SO}_2$  gas mass resurfacing rates are

comparable to or less than the 1.7-2.4 cm  $\text{yr}^{-1}$  rates derived from previous observations of the Prometheus plume [9, 7]. At the same time, from the color of Tvashtar in reflected sunlight we estimate that the 2007 Tvashtar eruption released  $\sim 10^9$  g of sulfur dust, and the inferred  $\text{SO}_2$  gas column density corresponds to an  $\text{SO}_2$  gas mass  $\sim 10^{13}$  g, thus the dust/gas ratio is  $\sim 10^4$ , and the  $\text{SO}_2$  gas released in the eruption is the primary resurfacing source.

Additionally, the highest  $\text{SO}_2$  gas densities ( $\sim 0.2\text{--}6.0 \times 10^{18} \text{ cm}^{-2}$ ) and resurfacing rates ( $\sim 1.7\text{--}100 \text{ cm yr}^{-1}$ ) inferred from amongst all the available HST spectral and imaging data taken between 1995 and 2007, were for the 1997 Pillan and 2001 Prometheus plume eruptions. The dust mass of the 1997 Pillan plume was  $\sim 10^{10}$  g, and so again the inferred dust/gas ratio is significantly less than 1, making the  $\text{SO}_2$  gas released in the eruption the primary resurfacing source. Each of these eruptions were driven by lava-frost interactions [7, 4] and formed in fields enriched in  $\text{SO}_2$  frost/slush [12, 3, 13]. Yet in 2003 when all known  $\text{SO}_2$  seepage had ceased, Pillan's  $\sim 2 \times 10^{16} \text{ cm}^{-2}$   $\text{SO}_2$  gas density was  $\sim 300$ x lower than in 1997, potentially producing  $\sim 0.33 \text{ cm yr}^{-1}$  of  $\text{SO}_2$  surface mass; thus, releasing gas mass comparable to that inferred from the Pele-type plume observations. This dramatic variability may indicate that the density of  $\text{SO}_2$  gas released in the detected Pillan eruptions is directly correlated with the presence/absence of  $\text{SO}_2$  slush seepage into the Pillan Paterae floor. This suggests that the mobilization of buried/trapped  $\text{SO}_2$  frost into regions of high volcanic (magma/lava) activity can significantly impact the  $\text{SO}_2$  gas density released during the resultant violent volatile driven plume eruptions. If so, this would also imply that the overall impact of the lava-frost interaction driven plumes on Io's resurfacing rate will be function of the local  $\text{SO}_2$  frost abundance at the time of eruption.

Finally, we note that these results provide important insights into the older Voyager spectral observations as it relates to the HST observations obtained during the Galileo era. In particular, the derived 1997 Pillan eruption  $\text{SO}_2$  gas density is comparable to the  $\sim 5.4 \times 10^{18} \text{ cm}^{-2}$  derived from the  $0.7 \mu\text{m}$  detections of  $\text{SO}_2$  gas absorption in the vicinity of the 1979 Loki plume (Pearl et al. 1979); thus these results validate the occurrence of high density ( $>10^{17} \text{ cm}^{-2}$ )  $\text{SO}_2$  gas eruptions on Io. Additionally, each of the plumes with high ( $>10^{17} \text{ cm}^{-2}$ )  $\text{SO}_2$  gas content are known to scatter strongly at nearly blue ( $\sim 0.33\text{--}0.4 \mu\text{m}$ ) wavelengths. In fact, the average I/F brightness level

inferred for both 1979 Loki eruption and the 1997 Pillan plumes was significantly higher in the NUV to blue ( $\sim 0.33\text{--}0.34 \mu\text{m}$ ) than at longer wavelengths; and, the 1997 Pillan  $0.33 \mu\text{m}$  I/F level was 10x brighter than that of previous detections of Io's Pele-type plumes, while Loki's  $0.34 \mu\text{m}$  brightness was 3x brighter than the Pele-type plumes. Thus, these results further establish the long suspected connection between the abundance of  $\text{SO}_2$  gas released during the eruptions and the magnitude of the plumes  $0.34 \mu\text{m}$  brightness level.

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