

A thermal plume model to investigate convection in Jupiter's troposphere

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1. Introduction

1.1. Observational background

Jupiter's atmosphere contains three condensable species (ammonia, ammonium hydrosulfide and water) what allows for moist convective activity. Storms and lightnings have been observed by Galileo [8], Cassini [3] and Juno [1] missions. Storms are typically few thousands kilometers large but we can also see smaller (~ 100 km) convective clouds. Juno mission revealed recently that they unexpectedly occur mainly in polar regions, less often at lower latitudes in belts and hardly ever in zones.

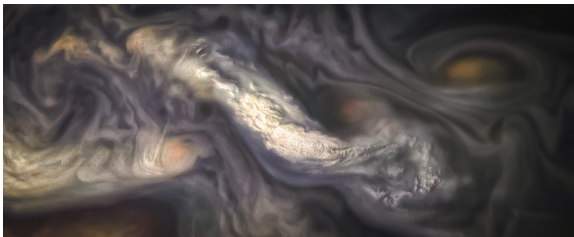


Figure 1: Jupiter's storm (credits: NASA Junocam)

1.2. Modeling context

Convection has already been modeled in Large Eddies Simulations (LES) (for example [5], [10]). In these simulations, vertical speed can reach several tens of m s^{-1} and stronger plumes can range from water condensation level up to ammonia clouds. That is why this activity could impact on the large scale circulation [7] and supply jet streams through inverse cascade of energy.

2. Thermal plume model

We seek to model moist convection in Jupiter's weather layer in order to determine its impact on the large-scale circulation and particularly on the jet streams structure. For that, we adapted a thermal

plume model coming from Earth LMD GCM [9] to gas giant. It's a 1D mass flux parametrization which computes a representative plume in an atmospheric column as soon as it detects unstable layers. For each layer penetrated by the plume: vertical speed w , entrainment e , detrainment d and vertical mass flux f (cf. figure 2 for dry case result) are computed thanks to the following equations :

$$\frac{\partial f}{\partial z} = e - d \quad (1)$$

$$\frac{\partial fw}{\partial z} = -dw + \alpha \rho \Gamma \quad (2)$$

Where Γ is an acceleration term due to buoyancy, ρ the density and α the updraft fraction. Then plume properties such as temperature or tracers mass mixing ratios are inferred. This parametrization has the advantage of consistently computing mixing and plume top with a possible overshoot.

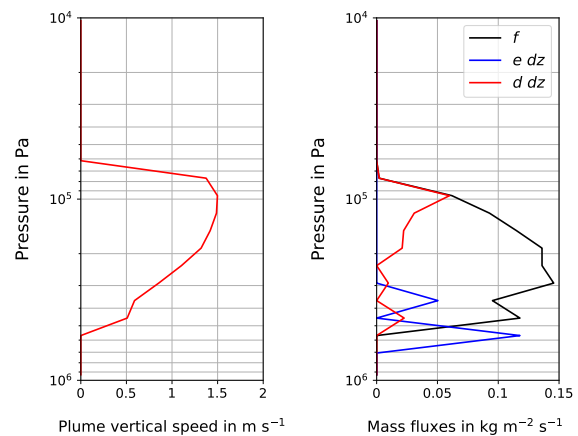


Figure 2: Dry case vertical speed and mass flux ($e dz$ and $d dz$ are respectively entrained and detrained mass flux)

3. First results

We performed several 1D simulations using a complete radiative transfer scheme ([4]) and the thermal plume model, with and without water. In both cases, we get radiatively entailed plumes but when there is water, competition between latent heat release and molecular weight leads to stronger but less high plumes (cf. figure 3). Moreover, the greater the water abundance is (from 0 to 9 solar abundance), the deeper convection ranges. The thermal plume model was coupled with the 3D GCM Dynamico Jupiter ([2], [4]) and some simulations was also performed.

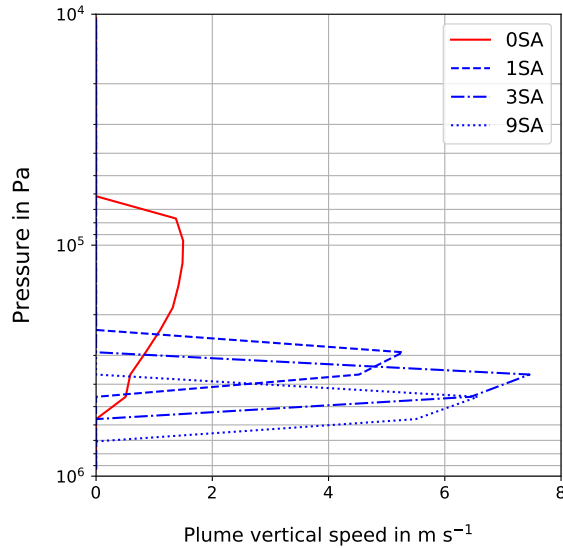


Figure 3: Vertical speed for various water solar abundance (SA)

Conclusion

During the conference, we will present further investigations about convection impact on Jupiter's troposphere and we will compare our results with storm distribution and altitude-latitude map of ammonia [6] measured by Juno.

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