

like behaviour under the radiative influence of lifted dust [6].

3. The influence of atmospheric circulations on dust lifting

We then focus on the influence of atmospheric circulations on dust lifting. Dust lifting in the GCM at low and medium (climate modeling) resolutions is strongly dependent on model resolution, not only for its ability to resolve topographic slopes, but also baroclinic ‘flushing’ storms and the regions of high vorticity associated with them. The GCM on its own also does not reproduce the same degree of interannual variability that is observed in the martian atmosphere [4]. In the mesoscale model, slope winds play a key role, modulated by thermal tides. Broadly consistent results are found between the two different dynamical strategies, if the resolution of the GCM is increased: mesoscale modelling and high-resolution (~ 40 km grid) global circulation modelling.

On the microscale, both daytime convective gusts and vortices contribute to lifting, though the latter are more efficient. Large-eddy simulations show in addition that near-surface friction velocity would be expected to be significantly larger in high plateaus than in low plains [9]. The influence of background winds on convective turbulent circulations is discussed.

4. Summary and further work

This work only represent a tentative first step towards understanding the coupling between dust and winds. Many further simulations are planned, including: (a) running longer climate experiments with GCMs under conditions appropriate to Mars both in the present day and with varying orbital parameters; (b) running global circulation models with enhanced dust in certain regions, such as Tharsis, where lifting and dust transport is thought to occur (cf. Arsia example in [7]) in order to examine the influence on large-scale climate; and, (c), running large-eddy simulations with radiatively-active, dusty convective vortices (dust devils) so as to study the influence of transported dust on dust devil characteristics.

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