

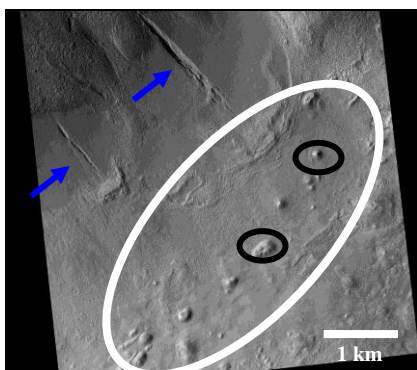
# Possible open-system (hydraulic) pingos in the Argyre impact region of Mars

R.J. Soare (1), S.J. Conway (2), J.M. Dohm (3) and R. El-Maarry (4)

(1) Geography Department, Dawson College, Montreal, Canada H3Z 1A4 (rsoare@dawsoncollege.qc.ca), (2) Department of Physical Sciences, Open University, Milton Keynes, United Kingdom MK7 6AA, (3) Earth-Life Science Institute, Tokyo Institute of Technology, Meguro, Tokyo, Japan 152-8551, (4) Physikalisches Institut, Bern Universität, Berne, Switzerland 3012.

## 1. Introduction

A group of small-sized mounds (~100–750m, see large white oval, Fig. 1) occur to the north of the Argyre impact basin, Mars. The shape, size, summit characteristics and clustering of the mounds, as well as a suite of geomorphological and geological features in the surrounding terrain, are suggestive of open-system (hydraulic) pingos (*HPs*) on Earth [1–7]. *HPs* are perennial water-ice cored-mounds that occur in permafrost (cold-climate and non-glacial) environments; they form and grow as the result of artesian pressure [1–7]. Here, we compare and contrast the Martian mounds and their geological context with a possible and terrestrial permafrost-analogue. Some preliminary work has discussed the possibility of *HPs* on Mars but, as of yet, strong candidate-sites on the Red Planet have been few in number [8–9].

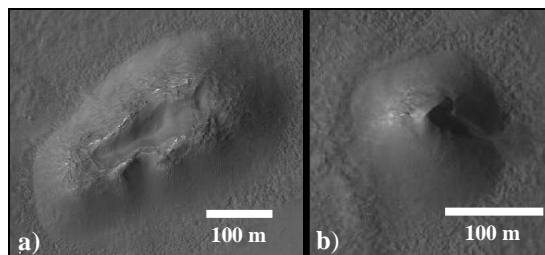


**Figure 1:** Small-sized mounds downslope of graben-like features. HiRISE image ESP\_020720\_1410, 317.714° E, 38.431° S; 25 cm/pixel. North is at the top. Image credit, NASA/JPL/ University of Arizona.

## 2. A putative periglacial-landscape in Argyre Planitia (AP)

North of the Argyre impact basin, Mars (317.714° E, 38.431° S), small-sized mounds are clustered downslope of numerous linear (graben-like) cavities (see blue arrows in Fig. 1). Mound morphology ranges from circular or sub-circular to elongate. Numerous mounds show summit

depressions, shaped irregularly; in some instances depression margins are highlighted by patchy bright material (Fig. 2a–b). One of the mounds (see small black upper oval in Fig. 1) displays a debris-fan at its base and an associated erosional-valley cut into its eastern flank (Fig. 2b). The pristine and unmodified morphology of the scar and fan suggest a relatively youthful age.



**Figure 2a-b):** Sub-images of HiRISE image ESP\_020720 that shows (see small black ovals in Fig. 1) a) elongate mound with an irregular summit-depression and bright patchy materials around its margin. b) circular mound with a linear valley cutting its eastern flank and an associated debris-fan at its base; note the bright patchy materials near its the summit to the left of the erosional channel. North is at the top.

Linear (graben-like) cavities also are observed upslope of the mounds; they could be indicative of faulting and of basement structural-control of the local if not regional landscape; some structures could be deeply-seated due to the Argyre impact-event and post-impact adjustment. The linear cavities form in and are bridled by high-albedo, smooth-textured terrain. In the literature, this type of terrain often is described as a latitude dependent and possibly ice-rich or cemented mantle (*LDM*) [10–12]. In turn, the linear cavities and the *LDM* are deformed by small-sized polygons (~5–20m in diameter); the latter could be the work of thermal-contraction cracking [13–14]. The unaltered or intact geometry of the polygons suggests that they postdate the formation of the linear cavities. Sinuous braided-channels that dissect the cavity floors and the polygonal-patterned ground point to subsequent hydrological activity. Arcuate ridges occur upslope of the *HP*-like mounds and

downslope of the linear cavities. These ridges could be terminal moraines, staking a line of furthest advance for glaciers that otherwise have ablated by sublimation or thaw [15]. Depending upon ambient conditions at the time of their formation, surface or near-surface melt-water might have been associated with the glacial system of which the ridges were a part.

### 3. Hydraulic pingos on Earth

Hydraulic pingos are perennial ice-cored mounds that occur in permafrost environments on Earth [1–5]. Mound morphology is varied, i.e. circular to elongate, and mound long-axes extend from metres to hundreds of metres [1–7]. Often, the **HPs** occur in groups or clusters, principally on valley floors and sides or in outwash plains [1–7]. Sometimes, the **HPs** occur on polygonised terrain formed by thermal-contraction cracking and underlain by ice or sand wedges at the polygon margins [1–7]. The **HPs** are thought to be the work of artesian pressure delivering sub- or intra-permafrost water to the location of mound development. Here, under freezing temperatures, an ice core forms and uplifts the sediments overlying it, creating a permafrost mound [1–7]. When an **HP** degrades, as the result of its ice core dissipating by thaw or sublimation, a summit depression or cavity may form [1–7].

Three principal geological-pathways are generally invoked to explain the formation of the **HPs**:

- (1) the presence of geological faults and structural discontinuities, which putatively deliver juvenile water to or near-to the surface where the **HPs** occur [1-7,16].



**Figure 3:** An **HP** in Greenland. Note the crater-like depression at the pingo summit. Also, a small spring has initiated the formation of a run-off channel [16].

- (2) a (potential) hydraulic gradient that moves sub- or intra-permafrost melt-water downslope, in areas of topographical relief, and towards a point(s) of emergence under artesian pressure where the permafrost is thin or weak [2].

- (3) a sub-category or variant of (2),: permafrost thaw in the accumulation area of glaciers, allowing melt-water infiltration [4], or thaw at the base of wet-based glaciers [17]. In either case, the melt-water migrates downslope, through the sub-permafrost that lies beyond the glacier's margin and, once again, emerges under the influence of artesian pressure to form **HPs** where the permafrost is relatively thin or weak.

### 4. Hydraulic pingos in AP?

Absent of a spade or a shovel, validating or invalidating a Martian landscape interpretation is difficult. However, we have constructed an **HP** formation-hypothesis based upon a possible cold-climate analogue on Earth that shows striking similarities to the studied Martian landscape.

As noted above, the shape, size, summit characteristics, clustering and slope-side location of the Martian mounds are consistent with permafrost regions on Earth where **HPs** are commonplace. In addition, the marginal presence of structural cavities, cavity-centred gully-flow, the **LDM** and arcuate ridges, are suggestive of a landscape where freeze-thaw cycling has been active. The fact that the mounds occur within this landscape assemblage brings the discussion concerning the possibility of **HPs** on Mars a step further than has been the case hitherto.

To the extent that this permafrost mound-hypothesis highlights recent activity in and around the Argye basin, it points to a treasure trove of far-reaching (~3.93 GA to the present and local to possibly global scales) geological, hydrological, and climatological data [18-19].

### 5. References

- [1] Müller, F., MoG 153, 3, 1-117, 1959. [2] Mackay, J.R., GPQ 52, 3, 1-53, 1998. [3] O'Brien, R., MoG 195, 1, 1-20, 1971. [4] Liestöl, O., NPÁ 7-29, 1975. [5] Yoshikawa, K., & Harada, K., PPP 6, 361-372, 1995. [6] Worsley, P. & Gurney, S.D., JQS 11, 3, 249-262, 1996. [7] Gurney, S.D., PPG 22, 3, 307-324, 1998. [8] Burr, D.M., et al., PSS 57, 541-555, 2009. [9] Hauber, E., et al., GSL, 356, 111-131, 2011. [10] Mustard, J.F., et al., Nature, 412, 411-414, 2001. [11] Milliken, R.E., et al., JGR 108, E6, 5057, 2003. [12] Head, J.W., et al., Nature 426, 797-802, 2003 [13] Mellon, M.T., JGR 102, E11, 25,617-25,628, 1997. [14] Levy, J.S., et al., Icarus 206, 229-252, 2010. [15] Arfstrom, J., & Hartmann, W.K., Icarus 174, 321-335, 2005. [16] Scholz, H., & Baumann, M., GGSB 104-108, 1997. [17] Humlum, O., et al., PR 22, 2, 191-215, 2003. [18] Dohm, J.M., et al., 44th LPSC, Abstract #2255, 2011. [19] Dohm, J.M. et al., this conference.