

## The Erebus Montes, Mars – Investigation of Ages and Amazonian Erosion Rates

S. van Gasselt (1), C. Orgel (1,2), J. Schulz (1)

(1) Institute of Geological Sciences, Department of Earth Sciences, Freie Universität Berlin, Malteserstr. 74–100, D-12249 Berlin (Stephan.vanGasselt@fu-berlin.de).

(2) Department of Physical and Applied Geology, Eötvös Lorand Tudományegyetem, H-1117 Budapest, Pázmány Péter sétány 1/c, Hungary.

### Overview

Lobate debris aprons are considered to be indicators for the presence of ice and water reservoirs on Mars and are therefore sensitive to climate variability. The northern hemisphere of Mars is characterized by three major populations of debris aprons (see, e.g. [12]): (1) the Tempe Terra/Mareotis Fossae region [2, 5], (2) the Deuteronilus/Protonilus Mensae [1, 4, 8], and (3) the Phlegra Montes (PM) [3]. The broader PM area can be subdivided into a number of smaller populations dispersed across parts of Arcadia Planitia (see figure 1) of which the Erebus Montes located at 180–195°E, 25–41°N form a well-confined set of features. We here focus on age and erosional characteristics of the northern Erebus Montes (see inset in figure 1).

Our study makes use of panchromatic image data obtained by the High Resolution Stereo Camera (HRSC) [9, 6] onboard Mars Express and the Context Camera (CTX) [7] onboard Mars Reconnaissance Orbiter. Image data analyses are supported by digital terrain-model data derived from HRSC based stereo imaging [10] and from Mars Orbiter Laser Altimeter (MOLA) [11].

We performed detailed geologic mapping at a scale of 1:10,000 and analysed age relationships and erosion rates based on a similar approach as outlined in [5] for the northern part of the Erebus Montes. The aim of this study is to compare feature characteristics to other populations in order to assess timing and the overarching control of landforms evolution in the Martian northern hemisphere.

The EM compare geologically relatively well with the Phlegra Montes in terms of individual feature morphologies. The concentration based on cluster analysis (figure 1) shows an up to 10 times higher concentration of remnants per 25 km<sup>2</sup> area peaking at  $3.4 \times 10^{-3}$  features for Erebus Montes. Debris aprons

show well-defined age signals ranging from 15 Myr up to 145 Myr. Some units even show continuous degradation implying active denudation of the Noachian to Hesperian-aged remnant massifs.

Based on the current status of investigations latitudinally dependent age trends cannot be observed which is likely to be related to the small extent of the northern region. Erosion rates determined at selected remnants are comparable to the Tempe Terra region with 0.1–0.3 mm·a<sup>-1</sup> (100–300 B) [5], depending on the model that has been used for our calculations. An explanation for such high Amazonian rates could be that much of the apron material has not been accumulated through denudation processes but by atmospheric deposition and removal of material from high-relief areas.

### Acknowledgements

This work is supported by the DLR Space Administration on behalf of the Federal Ministry for Economic Affairs and Energy, grant 50QM1301 (HRSC on Mars Express). We thank the HRSC experiment team at DLR Berlin and the HRSC operations team at ESOC for their successful planning, acquisition and processing of the HRSC data, and we thank the CTX team for their successful acquisition of data and for making it publicly available.

### References

- [1] Carr, M. H. & Schaber, G. G. (1977): Martian permafrost features.– *J. Geophys. Res.* **82**(28) 4039–4054.
- [2] Chuang, F. & Crown, D. A. (2005): Surface characteristics and degradational history of debris aprons in the Tempe Terra/Mareotis fossae region of Mars.– *Icarus* **179**(1), 24–42.
- [3] Dickson, J. L., Head, J. W., Marchant, D. R. (2010): Kilometer-thick ice accumulation and glaciation in the

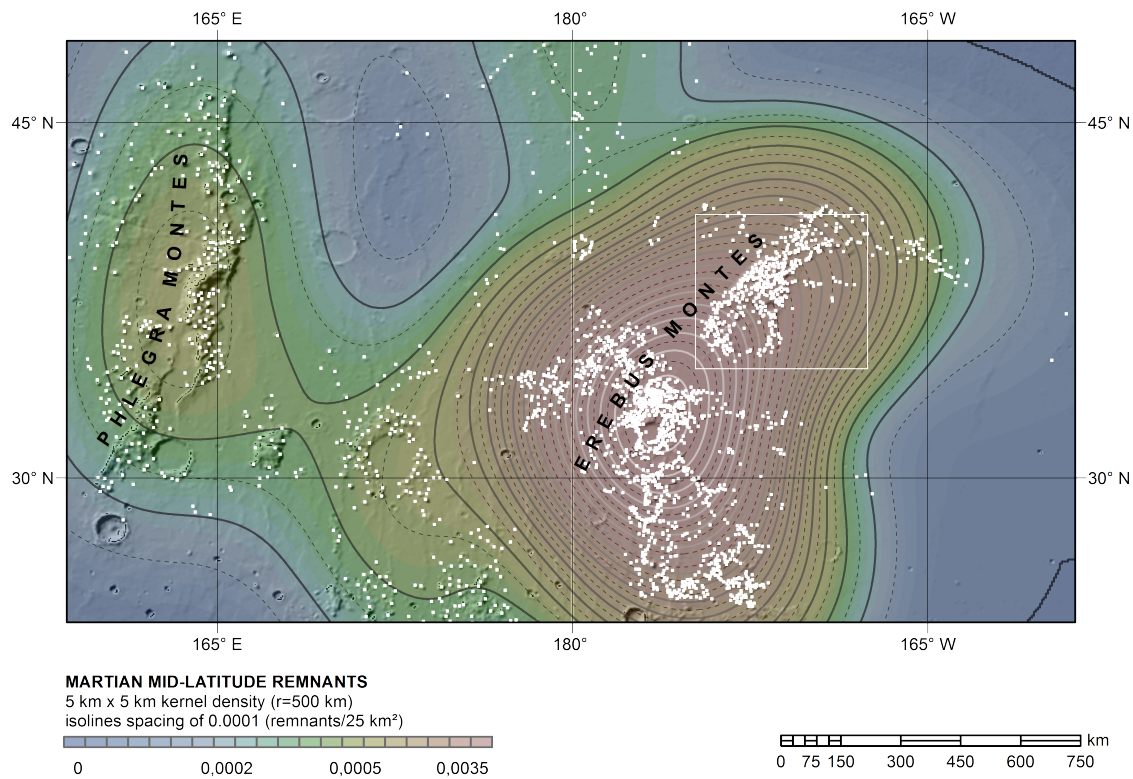


Figure 1: Distribution of remnant features in Arcadia Planitia (white dots) and kernel density plot superimposed on MOLA hillshade topography.

- northern mid-latitudes of Mars: Evidence for crater-filling events in the Late Amazonian at the Phlegra Montes.— *Earth Planet. Sci. Lett.* **294**(3–4), 332–342.
- [4] Fastook, J. L., Head, J. W., Marchant, D. R. (2010): Formation of lobate debris aprons on Mars: Assessment of regional ice sheet collapse and debris-cover armoring.— *Icarus* **228**, 54–63.
- [5] van Gasselt, S., Hauber, E., Rossi, A.-P. et al. (2011): Periglacial geomorphology and landscape evolution of the Tempe Terra region, Mars.— *Geological Society, London Special Publications* **356**, 43–67.
- [6] Jaumann, R., Neukum, G., Behnke et al. (2007): The high-resolution stereo camera (HRSC) experiment on Mars Express: Instrument aspects and experiment conduct from interplanetary cruise through the nominal mission.— *Planet. Space Sci.*, **55**(7–8), 928–952.
- [7] Malin, M. C., Bell, J. F., Cantor, B. A. et al. (2007): Context Camera Investigation on board the Mars Reconnaissance Orbiter.— *J. Geophys. Res.* **112**(E5), CiteID E05S04.
- [8] Mangold, N., Allemand, P., Duval, P. et al. (2002): Experimental and theoretical deformation of ice?rock mixtures: Implications on rheology and ice content of martian permafrost.— *Planet. Space Sci.* **50**(4), 385–401.
- [9] Neukum, G., Jaumann, R. and the HRSC Co-Investigator and Experiment Team (2004): HRSC: the High Resolution Stereo Camera of Mars Express.— in: Wilson, A., (ed.) *Mars Express – The Scientific Payload*.— ESA SP-1240, 17–35, ESA/ESTEC (Noordwijk).
- [10] Scholten, F., Gwinner, K., Roatsch, T. (2005): Mars Express HRSC Data Processing – Methods and Operational Aspects.— *Photogram. Eng. Rem. Sensing* **71**(10), 1143–1152.
- [11] Smith, D. E., Zuber, M. T., Frey, H. V. et al. (2001): Mars Orbiter Laser Altimeter: Experiment summary after the first year of global mapping of Mars.— *J. Geophys. Res.* **106**(E10), 23,689–23,722.
- [12] Squyres, S. W. & Carr, M. H. (1986): Geomorphic evidence for the distribution of ground ice on Mars.— *Science* **231**(4735), 249–252, doi: 10.1126/science.231.4735.249.