


TIDAL WALKING ON EUROPA'S STRIKE-SLIP FAULTS INSIGHT FROM NUMERICAL MODELING

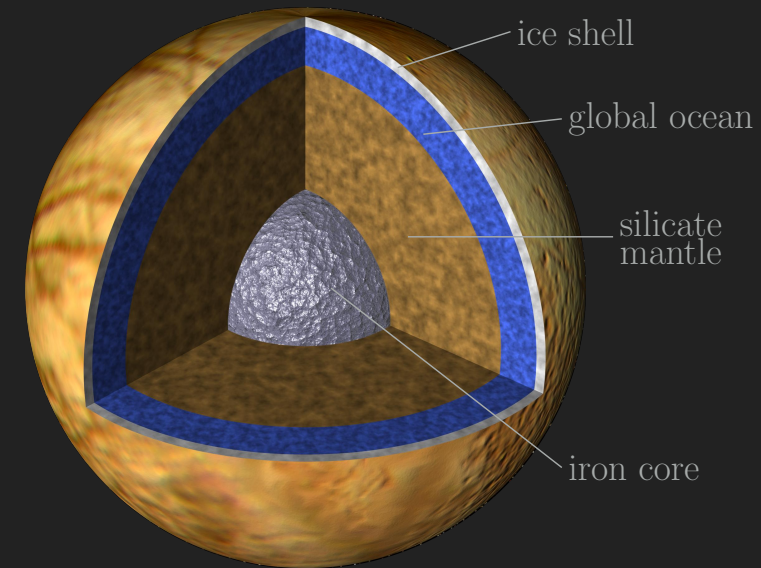
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EUROPA:

- ▶ is divided into iron core, silicate mantle, global ocean and ice shell
- ▶ has a young surface (tens of kyr) with multitude of superposed crosscutting lineaments
 - ▶ on some of the faults, strike-slip offset of a few kilometers was observed

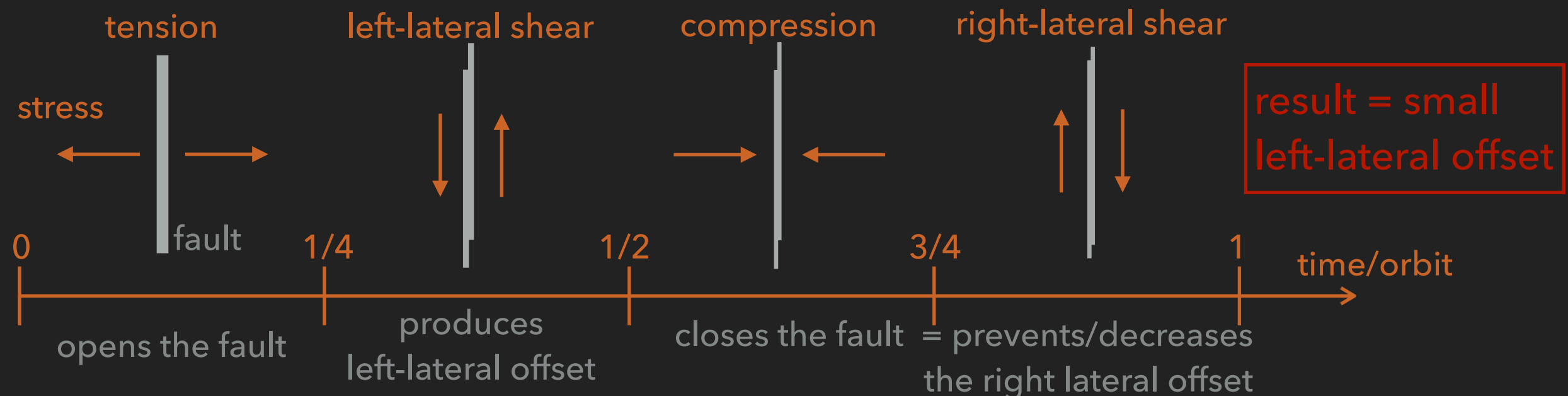


Background image: courtesy of JPL NASA.



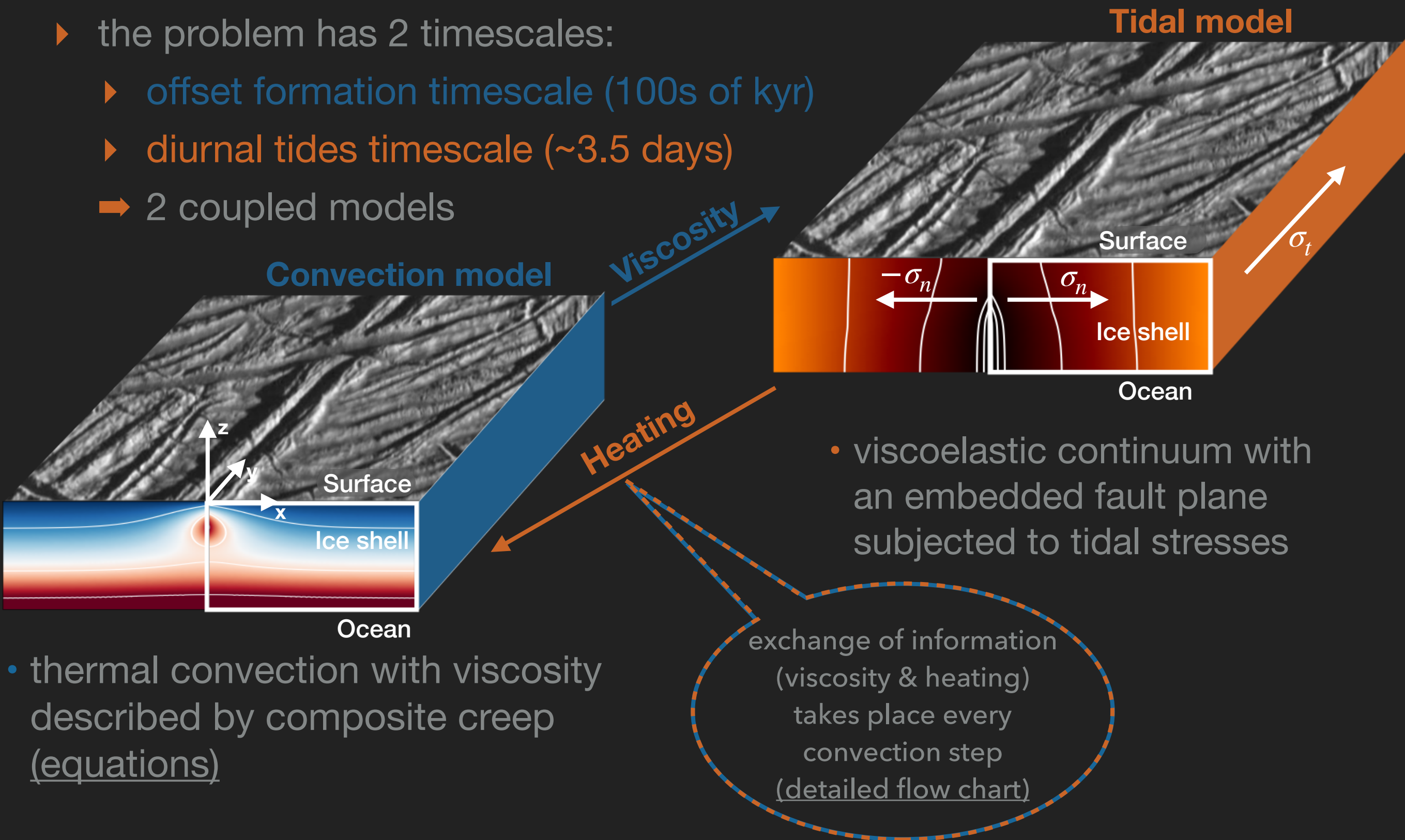
→ STRIKE-SLIP FAULTS

- ▶ formation of strike-slip offset?
 - ▶ reactivation of pre-existing faults
 - ▶ by diurnal tides => **tidal walking model** (Hoppa et al. 1999)
in the simplest setting



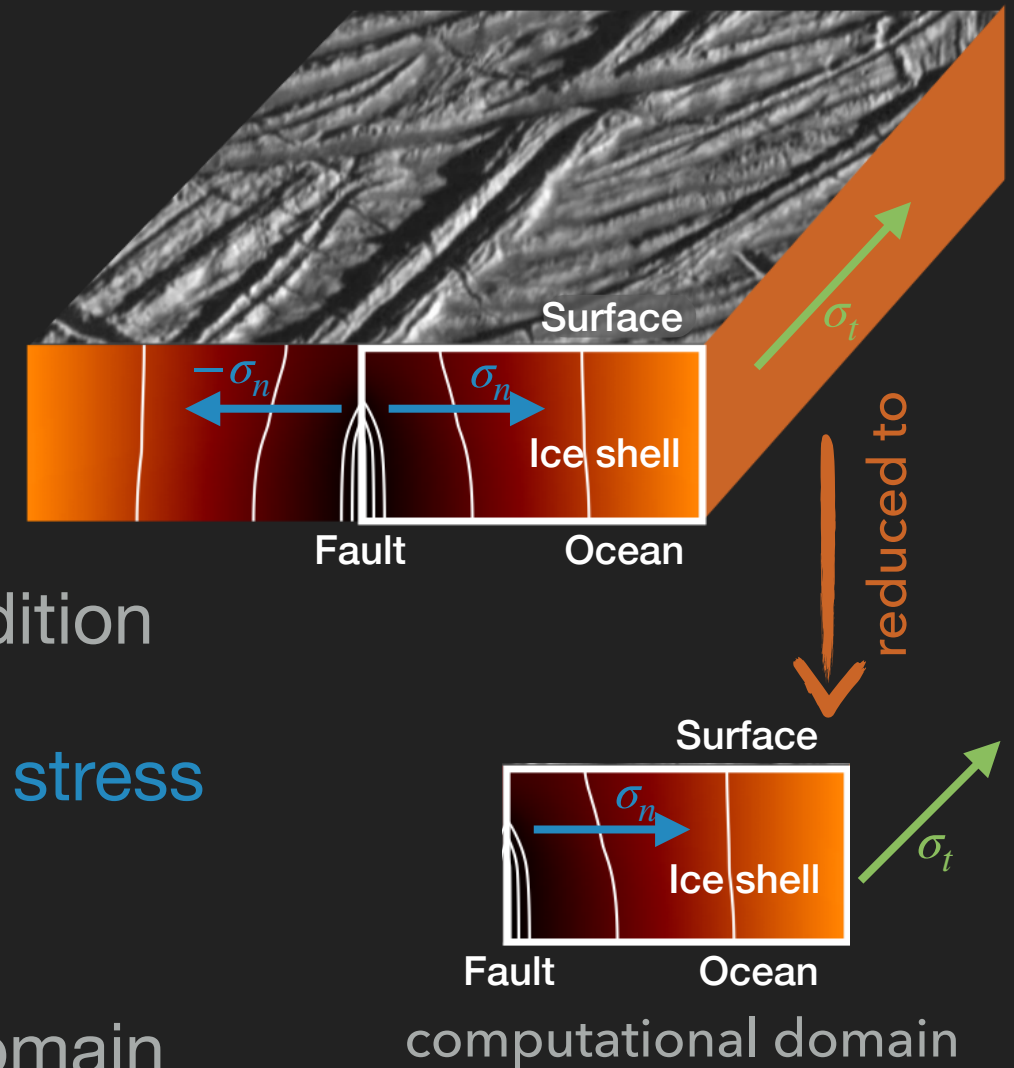
MOTIVATION —> MODEL

- ▶ aim: to test the tidal walking hypothesis with numerical model, quantify the offset and thermal signatures on surface
- ▶ the problem has 2 timescales:
 - ▶ offset formation timescale (100s of kyr)
 - ▶ diurnal tides timescale (~3.5 days)
- ➔ 2 coupled models



TIDAL MODEL

- ▶ solves mass and momentum balances of incompressible Maxwell medium
- ▶ with preexisting fault subjected to diurnal stresses



- ▶ tangential: $\sigma_t = \sigma_0 \sin(\omega t)$ - boundary condition
amplitude of loading stress
orbital frequency
- ▶ normal: $\sigma_n = \sigma_0 [\sin(\omega t - \varphi)]$ - in the yield stress
phase shift

- ▶ uniformity of the fault => 2D cut
- ▶ anti-symmetry of the problem => 1/2 of the domain

FAULT

- ▶ Mohr-Coulomb criterion - 2 states:
 - stick $v_{slip} = 0 \Leftrightarrow |s_x| < \sigma_Y$
 - slip $v_{slip} \neq 0 \Leftrightarrow |s_x| = \sigma_Y$slip velocity shear stress on the fault yield stress
- ▶ approximated with Navier-slip boundary condition and effective coefficient of friction (pseudoplasticity)
- ▶ yield stress $\sigma_Y = \max(0, \mu_f(\sigma_n + \rho g h))$
ice density magnitude of gravity acceleration depth
coefficient of friction
- ▶ varies in time due to normal diurnal tides => opening/closing of the fault

RESULTS


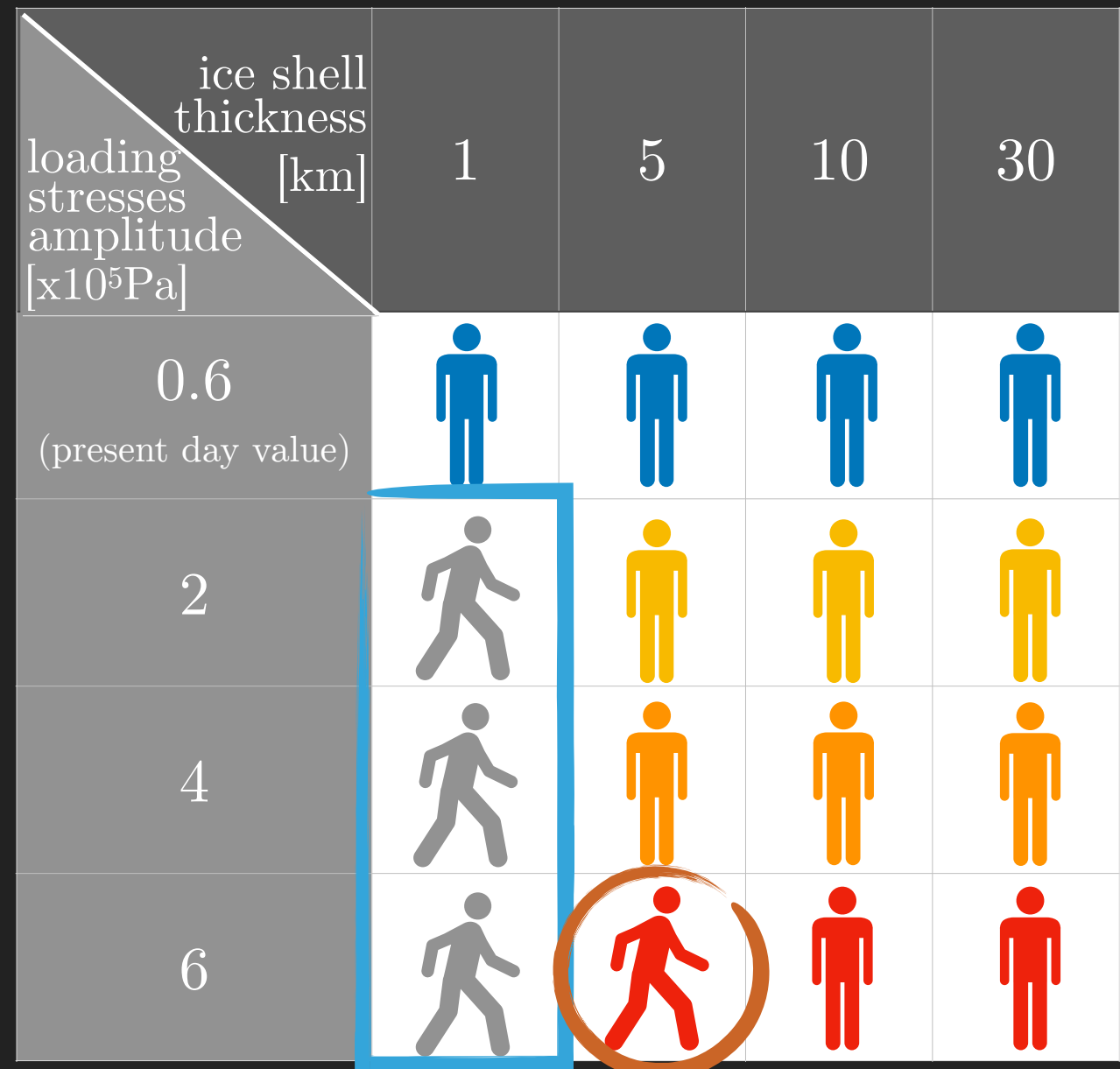
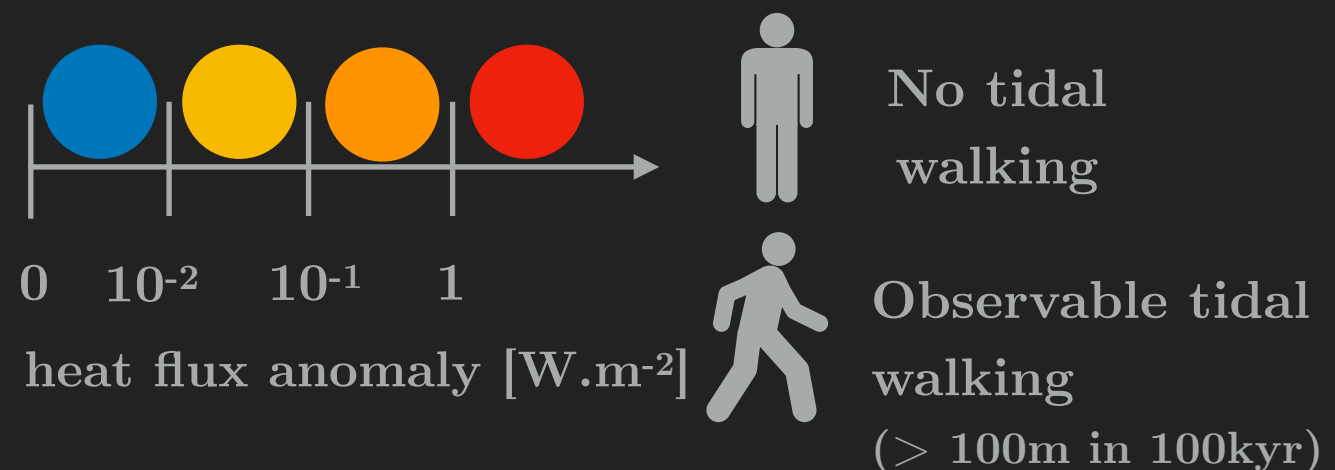
- ▶ model run for 100kyr
- ▶ heat flux anomaly = difference in heat flux above the fault at the end and at the start of the computation
- challenging to produce observable offset (few walking figures)
- unrealistically thin shell and high stresses needed except for:
 - ▶ thermally activated run = 
 - ▶ the fault and its surroundings are heated by frictional and shear heating, viscosity drops enabling the accumulation of the offset
- measurable heat flux anomalies can be produced without observable offset

Figure from Sládková et al., 2020



 stabilization (only qualitative results)



CONCLUSIONS

- ▶ an **observable offset** can be produced if **the active part** of the fault reaches either the **bottom** of the shell or penetrates to **a very low viscosity** zone (e.g. produced by extensive heating => thermally activated run)
- ➔ **thermo-mechanical** coupling is **important**
- ▶ whole shell penetration:
 - ▶ improbable in present days
 - ▶ possible if the **cracks** are **filled with water** compensating the overburden pressure and enabling a **significant offset** even for the **present-day** estimates of the tidal **forcing** amplitudes and **ice-shell thickness ≤ 10 km**
- ▶ **moderate surface heat flux** anomaly between 10 and 100 mW m⁻² is observed (even when the accumulated offset is negligible) = a possible tool for identifying active strike-slip faults on Europa

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