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Why try to use O_3 as a proxy for O_2 ?

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 \rightarrow O_2 + a reducing gas (i.e., CH_4) is a promising disequilibrium biosignature pair

There are scenarios where O_3 is more accessible than O_2

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No significant mid-IR O_2 features



Figure from the Origins Final Report

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O_3 is easier to detect at low O_2 levels



 O_3 is created by photolysis of O_2 with a UV photon

Ozone formation:

$$O_2 + h
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ightarrow O + O (\lambda < 242 \text{ nm})$$

 $O + O_2 + M
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Important: O_3 production has a nonlinear relationship with O_2 and is extremely sensitive to changing stellar UV flux

Modeling Earth-like planets with varying O_2 to study the O_2 - O_3 relationship

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 \rightarrow Photochemistry code



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Climate code



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Varied O_2 levels from 0.01 - 150% PAL for Earth-like atmospheres (PAL - Present Atmospheric Level)

Peak O_3 abundance occurs at $\sim 25\%$ Earth's Present Atmospheric Level (PAL) of O_2 Earth-Sun sytem:























The 3-body reaction that forms O_3 is faster at higher densities $(O + O_2 + M \rightarrow O_3 + M)$

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Stellar spectra are from Rugheimer et al. (2013) and MUSCLES



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Cooler stars do not have enough photons capable of O_2 photolysis (λ < 242 nm) for O_3 abundance to increase when O_2 decreases





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Ozone's 9.6 micron feature for Earth around the Sun:



 O_3 spectral feature depth is dependent on temperature difference between the planetary surface and the stratosphere

Ozone's 9.6 micron feature for Earth around the Sun:



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Ozone's 9.6 micron feature for Earth around the Sun:



Ozone's 9.6 micron feature for Earth around an M dwarf:



Ozone's 9.6 micron feature for Earth around an M dwarf:



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It's complicated, but we can understand it.



 O_3 is more accessible than O_2 in biosignature searches:

- For mid-IR wavelengths
- For low O₂ levels

Need to:

- Understand the UV spectrum of host star
- Perform photochemistry/climate atmospheric modeling

You can read more in Kozakis et al. (2022) in A&A