

## Third corner frequency of moderate Kamchatka earthquakes: station estimates agree over a network

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An acceleration spectrum of a local earthquake, after reduction to the source, is typically approximated (following Aki, 1967; Brune, 1970) by high frequency (HF) plateau. At the low-frequency side this plateau is bounded by a cutoff at the corner frequency  $f_c$ ; often this corner is split into two located at  $f_{c1}$  and  $f_{c2}$ . Hanks (1982) emphasized the “ $f_{max}$ ” phenomenon: even after correction for along-path attenuation the spectral plateau shows HF cutoff at a certain frequency  $f_{max}$ . Gusev (1983) and Papageorgiou and Aki (1983) ascribed its origin to the source. Anderson and Hough (1984), however, have shown that a near-site constant- $Q$  attenuation layer of limited thickness commonly exists, whose effect mimics source-related cutoff. Therefore, source acceleration spectrum can be thought as flat up to infinity. Still, evidence was accumulating year by year which indicates that the “ $f_{max}$ ” feature may represent, at least sometimes, a combined effect of two constituents: “source-controlled  $f_{max}$ ” and “site-controlled  $f_{max}$ ”. Recently Gusev and Guseva (2014) showed the existence of “source-controlled  $f_{max}$ ”, relabeled as  $f_{c3}$ , for majority of moderate Kamchatka subduction earthquakes using single-station data for 1993-2005. In that study, observed spectra were approximated, after reduction to the source, by a three-cornered shape, and three parameters  $f_{c1}$ ,  $f_{c2}$ , and  $f_{c3}$  were picked from them. The present study employs similar approach. S-wave records were used of horizontal channels of six rock-ground Kamchatka stations obtained from 372 subduction earthquakes of 2011-2014. The range of magnitudes is  $ML = 4-6.8$ , the range of hypocentral distances is 50–250 km. To reduce a recorded spectrum to the source, loss parameters used were estimated beforehand (Gusev and Guseva, 2015). Spectral amplifications of individual stations due to site effects were found empirically using both S-waves and coda. Using site corrections, loss-corrected spectra were reduced: first to the hard-rock PET station; second, to uniform medium. Overall, 1252 S source spectra were analyzed; and corner frequency sets were obtained for each, consisting of  $\{f_{c1}, f_{c2}, f_{c3}\}$  triples; some of these were incomplete. Event estimates for  $f_{c1}$ ,  $f_{c2}$ , and  $f_{c3}$ , obtained independently from data of each station, were then analyzed. About 90% of spectra featured  $f_{c3}$ ; in 10% of cases, no HF cutoff was seen up to 25 Hz. The inter-station scatter among  $\log_{10} f_{c1}$ ,  $\log_{10} f_{c2}$  and  $\log_{10} f_{c3}$  estimates is characterized by standard deviations equal to 0.17, 0.14 and 0.11, correspondingly. This level of accuracy was considered as acceptable. As the mere existence of  $f_{c3}$  is not generally recognized, it was expedient to check the fact in a different way. Pair correlation analysis was employed. We compared  $f_{c3}$  estimates obtained for the same event at two stations. Coefficient of correlation between several such pairs was estimated as 0.62 on the average, with 95% confidence interval equal to (0.58 0.66). This degree of correlation permits us to claim the existence of  $f_{c3}$  as proven for our data. The study was supported by the grant from the Russian Science Foundation (project 14-17-00621), and was carried out at the Kamchatka Branch of the Geophysical Survey of Russian Ac. Sci.