

1 Small-Scale (Flash) Flood Early Warning in the Light of 2 Operational Requirements: Opportunities and Limits with 3 Regard to User Demands, Driving Data, and Hydrologic 4 Modeling Techniques

5 Andy Philipp*, Florian Kerl*, Uwe Büttner* Christine Metzkes*+,
Thomas Singer+, Michael Wagner+, and Niels Schütze+

*Saxon State Office for Environment, Agriculture and Geology, Department of Water, Soil and Waste, 01109 Dresden

+Institute of Hydrology and Meteorology, Technische Universität Dresden, 01069 Dresden

6 Abstract

7 In recent years, the Free State of Saxony (Eastern Germany) was repeatedly hit by both extensive
8 riverine flooding, as well as flash flood events, emerging foremost from convective heavy rainfall. Especially
9 after a couple of small-scale, yet disastrous events in 2010, preconditions, drivers, and methods for deriving
10 flash flood related early warning products are investigated. This is to clarify the feasibility and the limits
11 of envisaged early warning procedures for small scale catchments, hit by flashy heavy rain events. Early
12 warning about potentially flash flood prone situations (i.e., with a suitable lead time with regard to required
13 reaction-time needs of the stakeholders involved in flood risk management) needs to take into account
14 not only hydrological, but also meteorological, as well as communication issues. Therefore, we propose
15 a threefold methodology to identify potential benefits and limitations in a real-world warning/reaction
16 context.

17 First, the user demands (with respect to desired/required preparation times, warning products, etc.) are
18 investigated. Second, focusing on small catchments of some hundred square kilometers, two quantitative
19 precipitation forecasts (including probabilistic QPFs) are spatially and temporally verified. Third, considering
20 the user needs, as well as the input parameter uncertainty (i.e., foremost emerging from an uncertain
21 QPF), a feasible, yet robust hydrological modeling approach is proposed on the basis of pilot studies with
22 deterministic, data-driven, and simple scoring methods. Our contribution delivers a synopsis of the already
23 acquired results for real-world (sub-)mesoscale catchments, comprising investigations of the aforementioned
24 three methodological pillars. An appropriate approach for deriving hydrological forecasts/prognoses relevant
25 for flash flood early warning is concluded from the presented results.

26 1 Introduction

27 For Saxony, considering the last two decades, the hydrologically most intense and most disastrous events
28 occurred in August 2002, August/September 2010, as well as June 2013 (LfULG, 2004, 2013, 2015). Total
29 damage for the aforementioned events sums up to 9 billion Euros (ca. 6.1 in 2002, ca. 0.85 in 2010 and
30 ca. 2.0 in 2013). Especially in August/September 2010, flashy events in small catchments caused large parts of
31 total damages. In this light, the Saxon State Government mandated an independent commission to identify
32 suggestions for improving flood risk management actions (Jeschke et al., 2010). One of the commission's
33 demands was to line out the potentials and limits of small-scale flash flood early warning approaches (i.e.,
34 based on hydrological forecasts).

1 As the authority responsible for operational flood forecasting and warning, the Saxon Flood Center (SFC)
2 drafted a corresponding project with a preferably holistic view on flood risk management procedures, especially,
3 when it comes to small-scale and flashy events. Therefore, a threefold approach is proposed, aiming at (1) the
4 assessment of the demands and requirements of potential users of early warning products; (2) the verification of
5 driving meteorological data for the targeted spatio-temporal scales; (3) checking the usefulness of a preferably
6 broad range of modeling approaches with regard to model skill, robustness, and regional applicability, for small,
7 potentially ungauged basins. The paper at hand provides a short overview of the current state of work and
8 illustrates a way towards an operational early warning system for small catchments in Saxony.

9 **2 Methods**

10 **2.1 User Survey**

11 To investigate the needs and demands of potential users of an envisaged flood early warning system for small,
12 fast-responding catchments in Saxony, a quantitative survey was carried out, based on an online questionnaire.
13 The questionnaire comprised 15 questions, with 12 multiple-choice questions, two questions with gradually-scaled
14 answers, and one question for the submission of verbal comments. Strictly speaking, the survey comprised
15 quantitative and qualitative elements. For the sake of brevity, the full questionnaire is not presented herein but
16 can be found in Philipp et al. (2015).

17 The surveyed sample was selected systematically (i.e., not randomly) and included all legal users (i.e.,
18 according to the Saxon Flood Alarm Bylaw; HWMO, 2014) of SFC products ($n = 578$) who were reachable
19 via email to be invited for participating in the online survey ($n = 491$). The interviewee affiliation spanned
20 administration/authorities at local/district/state level, fire departments and civil protection agencies, as well as
21 the private sector. It has to be stated that the interviewees did not represent lay people since they participate
22 in the official flood management procedures on a legal and regular basis.

23 The survey results were evaluated using descriptive statistics and subgroup analyses by means of contingency
24 tables. Therefore, given answers were investigated in an user-group specific manner, i.e., more than one
25 variable is considered at a time (multivariate approach). A question to address was whether specific user
26 groups answered differently or not. Such an effect can be induced by strongly differing sizes of sub-samples or
27 indicate a truly diverse response behavior. The literature suggests χ^2 -based dependency measures to clarify
28 such questions (Sachs, 1999). For the present study, Cramér's V and χ^2 -based p-values were used.

29 **2.2 Verification of QPFs**

30 The verification of meteorological data comprised two Quantitative Precipitation Forecasts (QPFs) which are
31 operationally used by the SFC. The investigated QPFs are the deterministic Numerical Weather Prediction
32 (NWP) COSMO-DE product (Baldauf et al., 2011) and the probabilistic “Quantile Forecast” (QF) for 16
33 specific areas in Saxony (cf. Figure 1), issued by DWD's Regional Service Center in Leipzig. The two QPFs
34 are compared against a Quantitative Precipitation Estimate (QPE), emerging from rain gauge data, which
35 was spatially interpolated (Ordinary Kriging) to derive areal precipitation estimates. Additionally, weather
36 radar data (DWD's RADOLAN-RW product; Sacher et al., 2011) was employed as another QPE reference. A
37 comprehensive overview of the herein considered QPFs and QPEs is given in Table 1.

38 DWD's Quantile Forecast represents a probabilistic, qualitative expert estimate of areal precipitation for
39 the next 36 hours and consists of three values/quantiles per forecasting time step. Since the QF is issued
40 for 16 specific areas in Saxony (i.e., river catchments with topographic partitioning according to elevation),
41 verification was based on the comparison of areal rainfall for the mentioned 16 regions, and spanned a period
42 from 04/2011 to 06/2014.

Table 1: Overview of the considered QPF and QPE products.

Product	Provider	QPF/ QPE	Type	Temporal resolution	Spatial resolution	Lead time	Update cycle
COSMO-DE	DWD	QPF	Deterministic NWP output (gridded)	1 h	2.8 × 2.8 km	21/27* h	3 h
Quantile Forecast (QF)	DWD-RWB LZ ⁺	QPF	Probabilistic forecast of mean areal precipitation	6/12 h [–]	Forecast regions from ca. 600 to 2.700 km ²	36 h	12 h
Interpolated rain gauge data	DWD	QPE	89 stations for the area of Saxony plus 25 km buffer	1 h	1 × 1 km [~]	—	1 h
RADOLAN-RW	DWD	QPE	Rain gauge adjusted weather radar estimate (gridded)	1 h	1 × 1 km	—	1 h

*27 hours since 01.30.2014 15:00 UTC. ⁺DWD's Regional Service Center in Leipzig. [–]Product comprises two consecutive 6-hour and two further 12-hour intervals. [~]Data gridded via Ordinary Kriging.

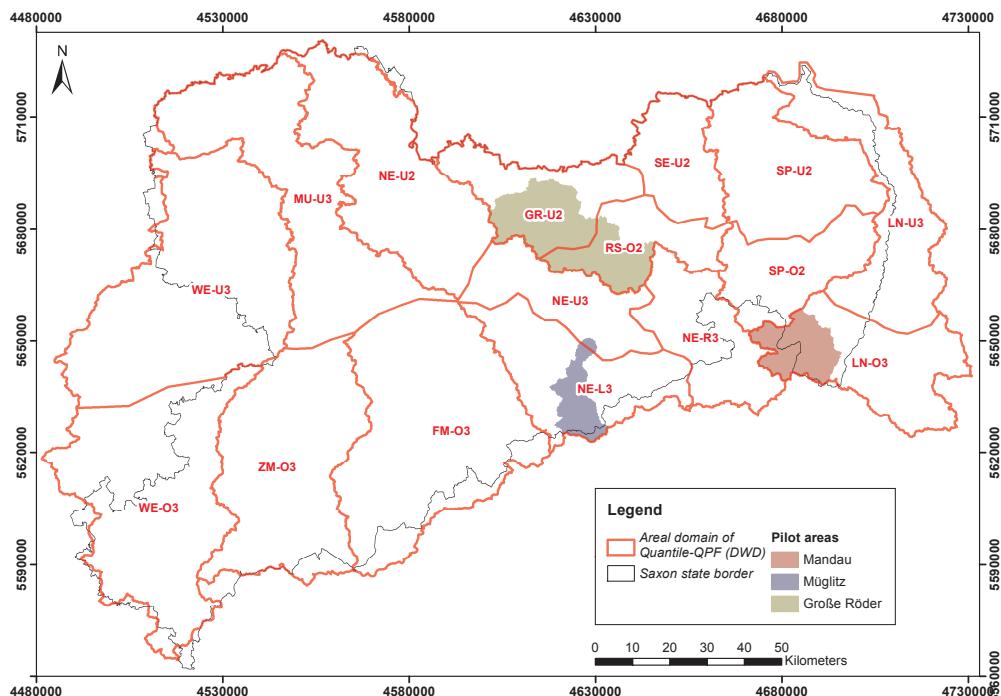


Figure 1: Overview map indicating the areal domain of the DWD Quantile-QPF for Saxony (16 regions; e.g., “FM-O3” indicates the parts of the Freiberger Mulde catchment above 300 m.a.s.l.). The area of the regions ranges between approximately 600 and 2.700 km². Furthermore, the hydrological pilot areas (cf. Section 2.3) are shown. Gauss conformal projection with reference at 12° E (Zone 4).

1 The comparison of areal rainfall was based on 6-hour sums, starting from 06:00 and 18:00 UTC. 6-hour
2 sums were chosen to accommodate the most coarse temporal resolution of the investigated products, given by
3 the Quantile-QPF. The QF features areal rainfall totals (for the 16 forecasting regions) for 0.9, 0.5, and 0.1
4 exceedance probability and two consecutive 6-hour and two further 12-hour intervals. The QF is updated twice
5 a day (at 06:00 and 18:00 UTC) and therefore is a rather general QPF product. However, a main task of the
6 herein presented verification was to evaluate the quality of this product against highly resolved NWP output
7 (i.e., COSMO-DE).

8 A QPF/QPE comparison typically employs a number of tools and methods (Jolliffe and Stephenson, 2012),
9 ranging from simple diagnostic (e.g., time series and totals comparisons, residual and bias analyses, scatter and
10 frequency plots) to integral, quantitative methods. Analyses are often based on threshold-oriented contingency
11 table evaluation and deliver typical verification/skill scores, e.g., FAR, POD (False Alarm Rate, Probability Of
12 Detection) or combined products, e.g., ROC curves (Receiver Operating Characteristic; Fawcett, 2006). More
13 detailed information on the herein employed QPF/QPE verification methodology (as well as concerning the
14 results) can be obtained from Kerl and Philipp (2015).

15 **2.3 Hydrological Modeling Approaches**

16 Three different hydrological modeling techniques were implemented and applied for three pilot areas in Saxony
17 (cf. Figure 1): first, a semi-distributed deterministic model (DeHM), second, a data-driven, neural-network model
18 (DaHM) and, third, a simple classification model, based on the scoring of flood-relevant parameters (ScoHM).
19 Subsequently, the modeling concepts and their application (with regard to calibration, data assimilation, etc.)
20 are briefly described. Only snow-free conditions were regarded for model development and application.

21 **Deterministic hydrological model (DeHM)**

22 DeHM model's topology is based on a nodal representation of sub-catchments. Model calculations are
23 performed for each node, whereas the calculation sequence is determined by the topological order of nodes;
24 each model node holds all relevant parameters. Runoff generation is portrayed by the SCS Curve Number
25 method. Runoff concentration is either modeled via an arbitrarily long cascade of linear reservoirs or via
26 response-function convolution. Channel routing is described with either a time-lag function, a cascade of linear
27 reservoirs, Muskingum method, or a translation-diffusion model. Since there are a number of multi-purpose
28 and flood-retention reservoirs in the pilot areas, flood control was specifically included in the model.

29 Model calibration was based on event-specifically masked hydrograph data and employs a mixed performance
30 criterion after Li et al. (2015). Data assimilation/state updating was realized with a simplified Kalman filter
31 with error variances, following Blöschl et al. (2014). More details on the DeHM model and its application can
32 be found in Schwarze et al. (2015).

33 **Data-driven hydrological model (DaHM)**

34 DaHM is an artificial neural network model, employing a feedforward two-layer perceptron (Hagan et al.,
35 2002). The input vector features flow, rainfall, and cumulative rainfall data with the general 15-element form
36 $I : [Q_{t-[0...3]}; P_{t-[0...3]}; P_{t-[0...6]}^c]$ (with hourly values of flow Q , rainfall P , and cumulative rainfall P^c). Adding
37 to that, and depending on the considered lead time in the forecasting case, inputs that bear a rainfall forecast
38 were included, e.g., for forecasting Q_{t+6} , the input P_{t+6} is added, for Q_{t+12} , $P_{t+6; t+12}$, respectively, whereas
39 the P_{t+x} values portray specific QPF lead times.

40 The Levenberg-Marquardt algorithm was applied for network training, whilst allowing the number of hidden
41 neurons range from 3 to 13. Event-wise masked hydrograph data and hourly areal rainfall were used for training.

Table 2: ScoHM scoring system.

	Description	Parameter bins	Sub-score range
Baseline susceptibility	Mean catchment slope	< 0.02/0.02/0.08/0.14/0.20	0 to 4
	Catchment shape factor*	< 0.20/0.20/0.40/0.60/0.80	0 to 4
	Degree of surface sealing	< 0.05/0.05/0.20/0.35/0.50	0 to 4
Dynamic susceptibility [~]	Proportion of fast runoff components ⁺	< 0.10/0.10/0.23/0.37/0.50	0 to 4
	SPI over the last 30 days ⁻	-3/-2/-1/0/1/2/3	-3 to 3
	Precipitation sum over the last 7 days	Sub-score percentiles ^{&}	0 to 4
	Precipitation sum over 12/24/48 hrs [#]	based on actual data from 01/2010 to 09/2015	0 to 4
Total susceptibility score			-3 to 31

*Catchment being more circular for values near unity. ⁺According to Peschke et al. (1999). ⁻SPI values rounded to integers. [~]In contrast to Collier and Fox (2003), snow-specific dynamic sub-scores were not considered. [#]Only highest sub-score is considered.

[§]Linear reservoir being charged with hourly precipitation. [&]Percentiles: < 75th/75th/90th/95th/99th.

¹ 15 training runs were evaluated for each specific hidden-neuron configuration and the best network was selected.
² Schwarze et al. (2015) give more details on the training and validation of the DaHM model.

³ **Scoring model (ScoHM)**

⁴ The employed scoring model resembles the Flooding Susceptibility Assessment approach proposed by Collier
⁵ and Fox (2003). The method is twofold; first, a baseline susceptibility is derived, based on morphological
⁶ features, e.g., slope, land cover, etc. Second, a time-variant, dynamic susceptibility is calculated, incorporating
⁷ the Standardized Precipitation Index (SPI; Edwards and McKee, 1997), cumulative precipitation measures, and
⁸ the response of a linear reservoir being charged with hourly precipitation.

⁹ The scoring is carried out according to Table 2; baseline sub-scores and the SPI sub-score are mapped linearly,
¹⁰ according to the range of each respective morphological feature. For the remaining dynamic susceptibility
¹¹ sub-scores, frequency analyses were applied to deliver specific percentiles that are in turn connected to specific
¹² sub-score values, e.g., P-sums within the 75th–90th percentile-range of the data result in a sub-score of 1, etc.
¹³ The method requires only one effective parameter, namely the recession constant of the incorporated linear
¹⁴ reservoir, which was manually calibrated to a global value of 8 h.

¹⁵ In contrast to the DeHM and DaHM models, the ScoHM approach does not rely on observed flow data,
¹⁶ neither in the sense of directly including auto-correlative signals, as applies for the data-driven DaHM model (in
¹⁷ form of the Q_{t-x} inputs), nor indirectly via data assimilation/state updating, as applies for the deterministic
¹⁸ DeHM model. Therefore, the ScoHM approach might offer a robustly transferable methodology towards
¹⁹ prediction in small, ungauged basins.

²⁰ **3 Results**

²¹ **3.1 User Survey**

²² Herein, the most important results of the user survey (cf. Section 2.1) are presented in a concise manner;
²³ a more detailed presentation can be found in Philipp et al. (2015). The response rate was 76 % ($n = 373$),
²⁴ which is extraordinarily high (with 69 % or $n = 339$ completely answered questionnaires) and is mainly a result
²⁵ of the systematic sampling (cf. Section 2.1). For 11 out of 15 questions, user-group specific replies were
²⁶ not distinguishable in a statistical sense. The outcomes of the statistical analysis of the survey data can be
²⁷ summarized as follows:

²⁸ **Information and pathways:** (1) The interviewees request selective, event-related information or inform
²⁹ themselves on an event-related basis (rather than on a regular basis). (2) 37 % of all users trust that a more

1 regular and more frequent distribution of warning products will provide increased security for their management
2 decisions, even if the meteorological and hydrological trend remains unchanged. (3) All groups, except the
3 group “private persons”, attach greatest importance to the internet in contrast to other communication channels
4 (e.g., fax, video text, voice mail). The official flood warnings issued by fax or email are also used for information
5 by a majority of users. (4) A high availability of warning services and products is deemed important by a vast
6 majority of users, especially in the case of flooding.

7 **Flood warning products:** (1) A short-termed, but more precise warning is preferred over a long-term
8 estimation, carrying presumably more uncertainty. (2) The majority of users ($> 65\%$) are interested in receiving
9 a possibly reliable forecast of the peak water level. 45 % of users would appreciate being informed about
10 the peak timing. (3) Most popular products for fulfilling early warning purposes are forecasted hydrographs
11 with uncertainty bands (about 50 % of all persons interviewed), as well as a catchment-oriented classification
12 products (“traffic light”, approximately 40 % of all persons interviewed).

13 **Lead time:** (1) The minimum required lead times amount to ≤ 3 h (9 % of users), ≤ 6 h (27 %), ≤ 12 h
14 (50 %), ≤ 24 h (83 %), ≤ 72 h (98 %). (2) A lead time of ≤ 12 h is deemed to be adequate by a slim majority
15 of users in small catchments ($< 200 \text{ km}^2$).

16 **Miscellaneous:** (1) The interviewed user groups vary significantly in terms of the replies given when
17 being asked for the requested updating frequency of flood warnings and their communication via email or fax.
18 (2) Furthermore, the interviewees of various user groups specifically replied to the questions concerning the
19 quality of current SFC products and the quality of the work of the SFC. (3) Moreover, no significant differences
20 in the response behavior of the various user groups could be identified by statistical means.

21 3.2 Verification of QPFs

22 The investigated QPFs (COSMO-DE and QF) were compared against areal precipitation estimates, based on
23 gridded rain gauge data, and, additionally, a radar-based QPE (RADOLAN-RW product). First, threshold
24 exceedance frequencies were derived from the QPEs and QPFs (cf. Figure 2). COSMO-DE delivers exceedance
25 frequencies which are close to the ones derived from rain gauge data. RADOLAN (QPE) underestimates the
26 threshold exceedance frequencies from rain gauge data, whereas the chance of underestimation is higher at lower
27 thresholds, and vice versa. The dashed line at an exceedance frequency of 10 indicates that thresholds greater
28 than 10 mm/6 h should be evaluated with caution, due to limited data sample sizes (i.e., less than 10 events
29 in the investigated period). Due to QF-product related conventions (areal precipitation sum $< 4.5 \text{ mm/6 h}$
30 is set to zero), the exceedance frequencies of the QF remain constant for thresholds $< 4.5 \text{ mm}$. Threshold
31 exceedances drawn from the QF's 50th and 10th percentile are generally more frequent than the observed ones
32 (i.e., from rain gauge data), whereas the 90th percentile underestimates observed frequencies.

33 Second, for a more in-depth view at the regarded QPFs, the contingency-based measures POD and FAR
34 were evaluated (Figures 3 and 4). Again, due to the product-specific convention of the QF, the results for
35 the thresholds $< 4.5 \text{ mm/6 h}$ are identical. Following Winterrath et al. (2012), a minimum of 10 observed or
36 predicted threshold exceedances should be required for the calculation of skill scores. Therefore, POD and FAR
37 were not always evaluated for higher thresholds. Generally, higher precipitation thresholds are connected with
38 lower POD and lower FAR values, and vice versa. Furthermore, the skill variance amongst the forecast areas
39 increases with increasing precipitation thresholds. POD = FAR indicates a boundary threshold for which the
40 considered QPF has no predictive benefit anymore. This boundary is not reached for both QPFs, concerning
41 the investigated thresholds. Finally, for the regarded QPFs, COSMO-DE exhibits the best performance with
42 regard to POD/FAR relations and skill variance.

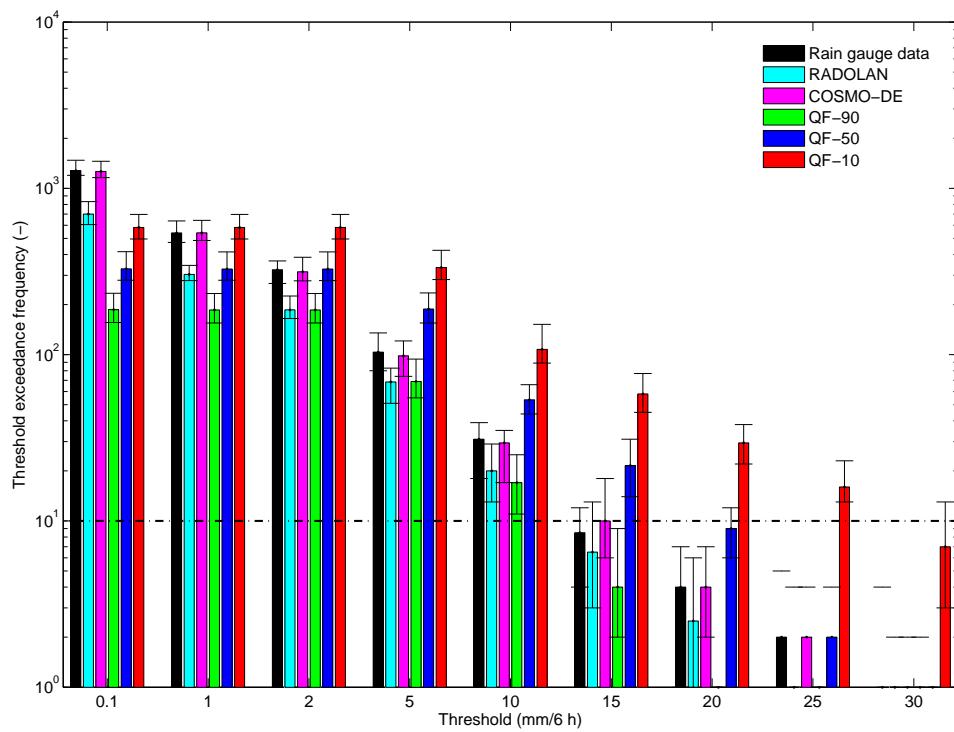


Figure 2: Threshold exceedance frequencies of 6-hourly areal precipitation sums for QPEs (gridded rain gauge data, RADOLAN-RW) and QPFs (COSMO-DE, Quantile Forecast) from 04/2011 to 06/2014. The bars show the median of exceedance frequencies for the respective precipitation products for the 16 forecast areas (cf. Figure 1). The whiskers illustrate the minimum and maximum values; the dashed line depicts an exceedance frequency of 10.

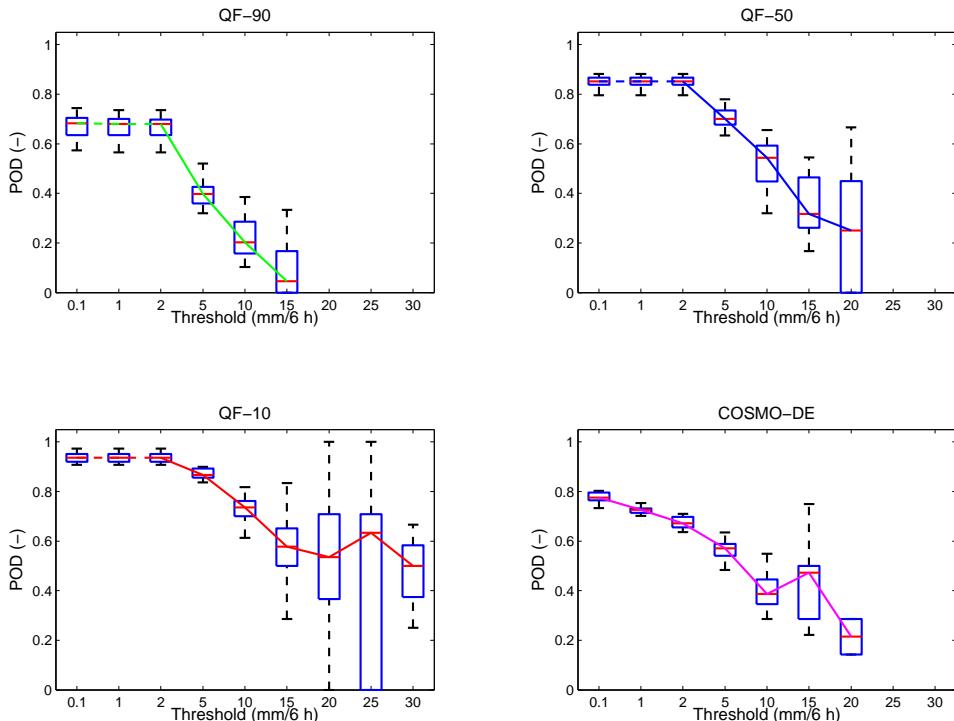


Figure 3: Probability Of Detection (POD) according to thresholds of areal precipitation sums ranging from 0.1 to 30 mm/6 h for the Quantile Forecast and COSMO-DE from 04/2011 to 06/2014. The box plots indicate the spread of POD over the 16 forecast areas.

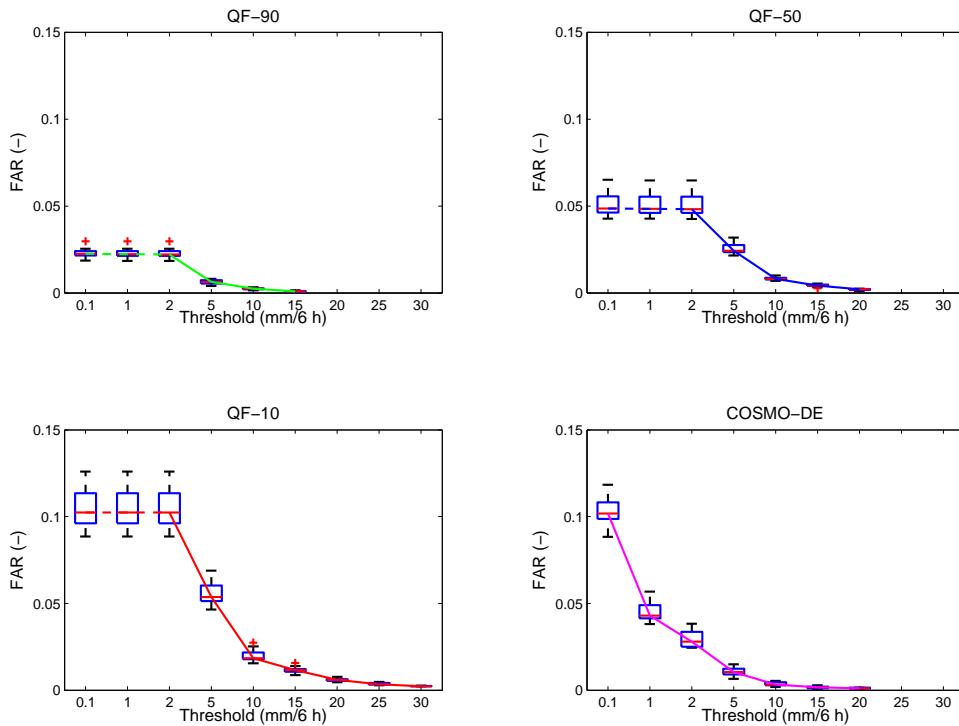


Figure 4: False Alarm Rate (FAR) according to thresholds of areal precipitation sums ranging from 0.1 to 30 mm/6 h for the Quantile Forecast and COSMO-DE from 04/2011 to 06/2014. The box plots indicate the spread of FAR over the 16 forecast areas.

3.3 Hydrological Model Validation

The three presented models (DeHM, DaHM, ScoHM) were applied for the three aforementioned pilot areas (cf. Figure 1). The herein investigated QPEs (gridded rain gauge data and RADOLAN data) and QPFs (COSMO-DE and QF; cf. Sections 2.2 and 3.2) were used as meteorological drivers (for the current state of work, ScoHM was charged with rain gauge data only). Validation for the DeHM and DaHM models is straightforward since modeled hydrographs are simply compared against observed ones. Model evaluation is a bit more delicate for the ScoHM results, since the ScoHM output (i.e., dimensionless scores) does only qualitatively correlate with observed flow values. Therefore, a quantile-mapping procedure (Piani et al., 2009) was applied to relate thresholds of Q with corresponding total-score values.

Model performance was (besides other integral measures as RMSE, NSE, etc.) evaluated on the basis of threshold-oriented contingency table analyses, i.e., it is checked if modeled output matches/exceeds a certain observed flow/alarm level or not. More specifically, the variation of threshold values delivers a set of corresponding skill scores, e.g., POD values with corresponding FARs. These POD/FAR tuples were used to establish catchment-specific ROC curves (Fawcett, 2006). The ROC curves were finally integrated to deliver AUC (Area Under Curve) values, with AUC near unity for a near-perfect model prediction and near 0.5 for no predictive skill. For brevity, results are presented and discussed for the Mandau catchment only, featuring four river gauges.

Generally, different combinations of lead times and update cycles (i.e., the time after which a new forecast is calculated) were investigated; herein, results for an update cycle length of 12 h are presented. Event-specifically masked, hourly hydrograph data and hourly rainfall observations were used during model validation. Data which were used in model calibration/training were not used for validation purposes. Data originated from the period of 2010 to 2015. Figure 5 comprehensively shows the validation results at four gauges within the Mandau pilot region. For the QF-QPF, results for the 50th percentile are exemplarily shown.

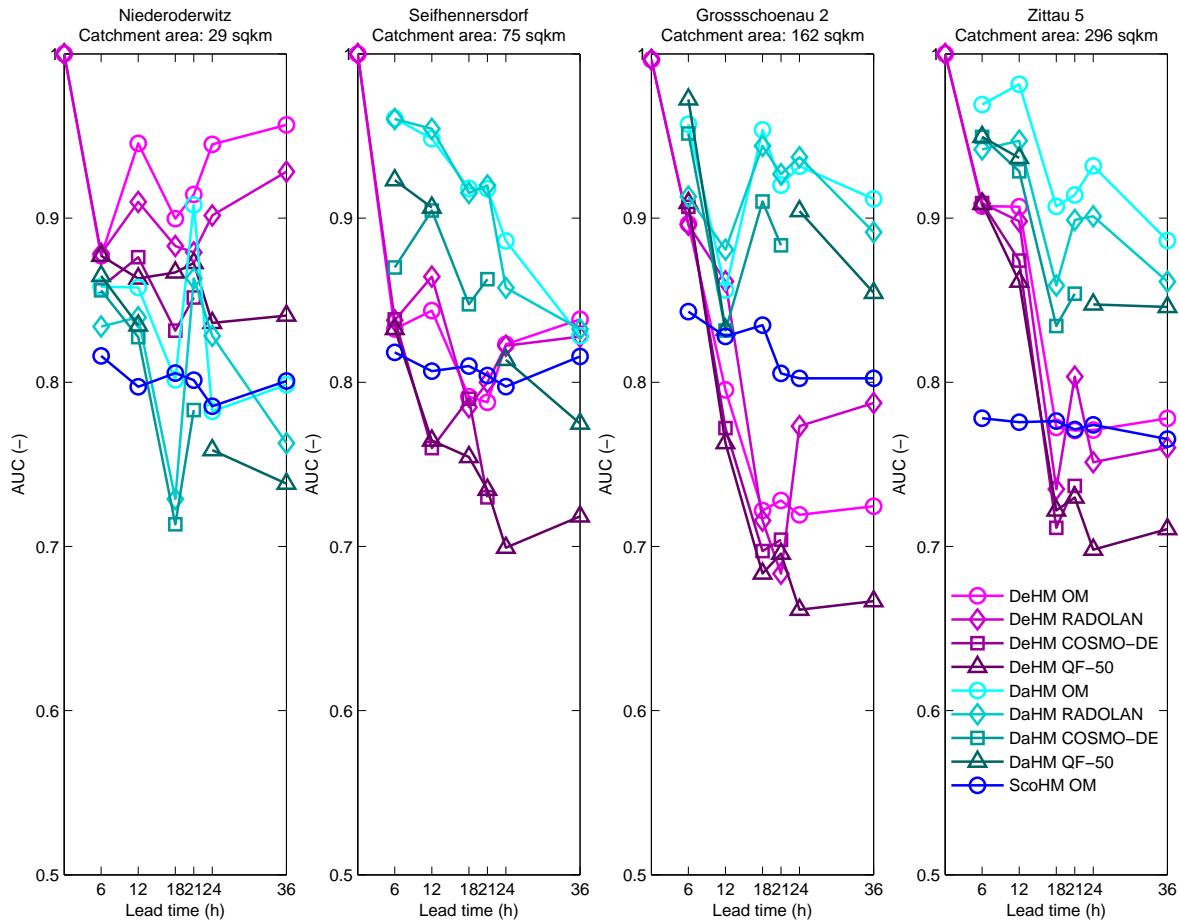


Figure 5: Results of hourly, threshold-oriented model evaluation for DeHM, DaHM and ScoHM output in the Mandau pilot region. Lead times range from 6 to 36 hours, update cycle is 12 hours. OM: ombrometer data (i.e., gridded rain gauge data); RADOLAN: QPE from weather radar scans; QF-50: 50th percentile of Quantile Forecast; COSMO-DE: NWP output. Skill for DeHM at a lead time of zero is based on true model output after assimilation/updating and can be slightly smaller than unity (e.g., apparent for Großschönau 2).

1 For the smallest sub-catchment, Niederoderwitz (29 km^2), DeHM performs best; for the three larger
2 sub-catchments, DaHM features the highest AUC values. However, DeHM and DaHM performance trends to
3 decrease with increasing lead time; ScoHM features a quite constant/robust AUC development. The reason
4 might be that for shorter lead times (6 h) the auto-correlative Q_{t-x} signal, included directly or indirectly in
5 the DeHM and DaHM model, leads to improved performance. This does not apply for the ScoHM results,
6 since the model does not rely on observed flow data. Generally, ScoHM exhibits AUC values around 0.8 which
7 indicates a good overall predictive skill, foremost, when keeping in mind the generality and straightforwardness
8 of the model approach.

9 It can be seen from Figure 5 that QPE data delivers highest predictive skill; incorporating RADOLAN
10 and rain gauge data as precipitation inputs leads to similar skill. Predictive skill under QPF data (QF and
11 COSMO-DE) is generally lower. For different QPFs as drivers, resulting skills do not differ greatly. Apparently,
12 the observed differences in QPF quality (cf. Section 3.2) do not systematically impact hydrological model skill.
13 Finally, it is important to state that validation was carried out on the basis of hourly values; a more general
14 evaluation, e.g., comparing only the highest values within a specific temporal window (e.g., 6 hours), would
15 yield considerably higher skill scores.

4 Conclusions and Outlook

In this study, user demands, driving data, and hydrologic modeling techniques were evaluated within a real-word application context in order to illustrate a way towards a flash flood early warning strategy for (sub-)mesoscale catchments in Saxony. First, the results suggest that the majority of potential users of flood warnings would be satisfied with forecasting lead times of up to 24 hours and that users are foremost interested in predicted peak water/alarm levels (rather than peak timing). Second, on the basis of meteorological verification results, highly resolved NWP data seem to offer the best predictive skill, compared to more general, areally integrated products. Third, differences in the quality of meteorological driving data do not greatly influence hydrological model skill. Fourth, a clear statement on the superiority of one hydrological model over another cannot be made.

In fact, if simple classification models would be sufficient to satisfy warning needs (e.g., providing the information whether or not a specific threshold is likely to be exceeded in the next forecasting interval), results show that such a modeling approach (i.e., ScoHM) performs with favorable skill, compared to more sophisticated modeling techniques, and without introducing cumbersome parameter estimation problems and limited (DeHM) or even nonexistent (DaHM) regional transferability. However, overall forecasting skill always decreases with increasing randomness of driving events and conditions, i.e., the more rare/focused/intense the flood-causing processes and/or the longer the lead time, the smaller the chance of correct detection/warning.

Further research is currently carried out regarding the statewide implementation and comparative evaluation of the herein considered hydrological modeling approaches. Meteorological verification will be carried out for smaller spatio-temporal scales. ScoHM will be validated for QPF inputs. Generally, the set of QPFs will be extended to DWD's 21-member ensemble product, COSMO-DE-EPS. Thus, allowing a comprehensive probabilistic verification and validation. Another goal is evaluating model-specific extrapolation skill to propose a feasible regionalization methodology for deriving threshold-based warnings for ungauged basins.

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